

Evaluating the Validity of Turbulent Kinetic Energy Dissipation Rate derived from Insitu 1D velocity fluctuation measurements

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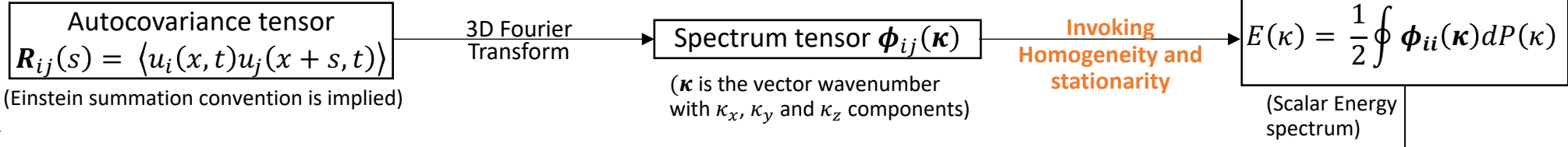
Background

- Turbulent kinetic energy dissipation rate (ϵ) is a fundamental parameter characterizing the structure and intensity of turbulent flows.
- Measurements of Instantaneous turbulent kinetic dissipation rate requires the complete knowledge of 3D velocity field at high spatial and temporal resolution.
- Such measurements require very high sampling cadence ($\sim 10^5$ KHz) and low-noise ($< 10^{-8}$ m²s⁻²/Hz) (or very high SNR).
- Insitu measurements of atmospheric turbulence made using ground based and airborne instruments provide coarsely sampled time-series data of velocity components.
- Typically, only spatially and temporally averaged estimates of ϵ are derived from 1D insitu measurements.
- The 1D measurement data is subject to either a model **second-order structure function fit** or a **model turbulence energy spectrum fit** (in the inertial and/or viscous subrange) based on the hypotheses of Kolmogorov's turbulence theory.

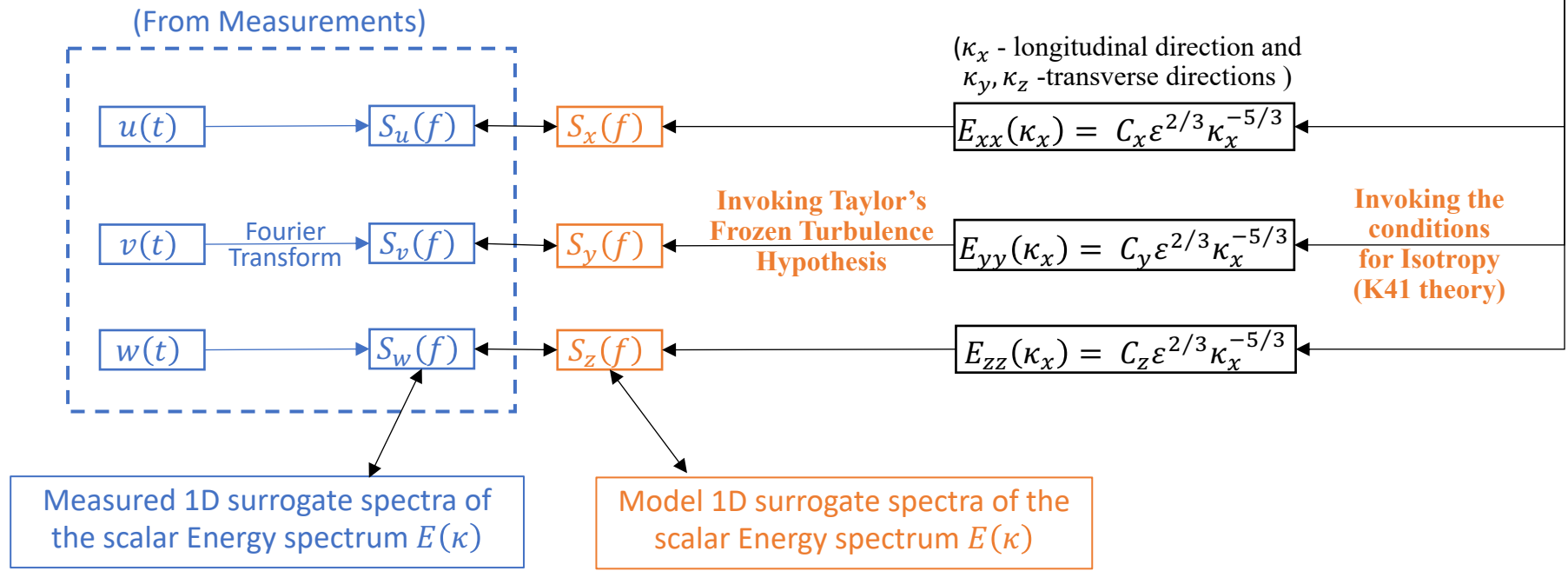


Assumptions in Spectral ε Estimation

Where $\kappa = |\boldsymbol{\kappa}| = \sqrt{\kappa_x^2 + \kappa_y^2 + \kappa_z^2}$

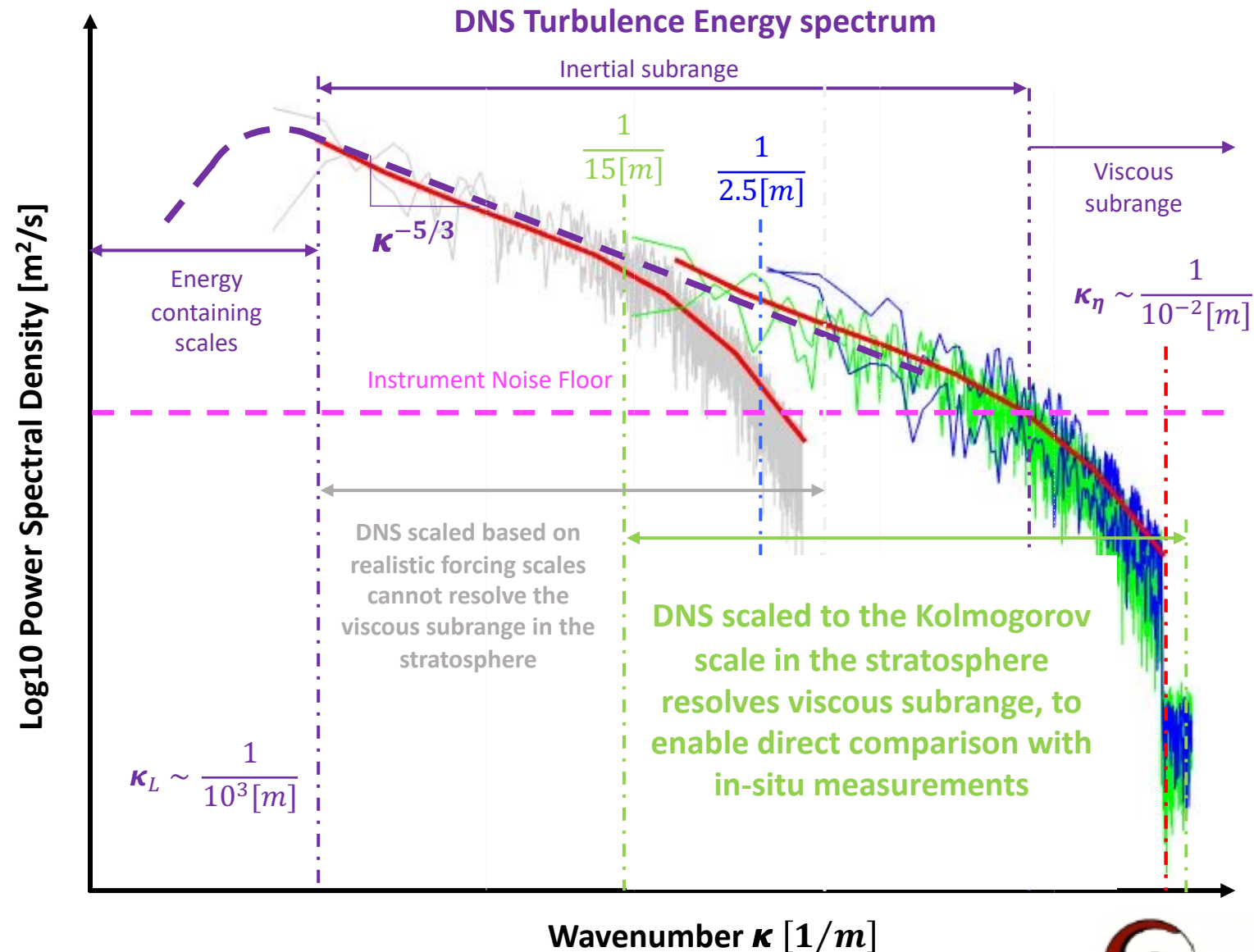


- These assumptions hold true for stationary HIT.
- Insitu observations intercept various states of atmospheric turbulence that likely do not conform to stationary HIT.
- Necessary to evaluate the applicability of the spectral estimation method to atmospheric turbulence measurements.



Synthetic Observations using DNS

- Provides a “truth” data set for conducting synthetic observations to validate the spectral methods used for in-situ measurements of turbulence intensity (TKE dissipation rate).
- Allows assumptions commonly made in the spectral methods to be tested directly.
- Nontraditional DNS scaling used to match Kolmogorov scale and dissipation rate seen in the stratosphere.
- Well-validated spectral shape enables cm-scale turbulence intensity to be indirectly measured/predicted.
- Similar synthetic observations for deriving ϵ from radar observations (1-3 [km]) (using LES), sounding rockets (80-90 [km]) (using DNS), and flying hotwire measurements in the wind tunnel (using DNS) are described in **Lundquist et al 2020, Strelnikov et al 2021, and Schroder et al 2023.**



DNS Setup and Metrics

Model Setup

- Turbulence Reynolds Number $Re \sim \left(\frac{L}{\eta}\right)^{4/3} \sim \mathbf{12,000}$
- η – Kolmogorov length scale, L - DNS box size
- *Forcing*: Energy injected at wavenumbers $\left(\frac{2\pi}{L}\right), \left(\frac{4\pi}{L}\right), \left(\frac{6\pi}{L}\right), \left(\frac{8\pi}{L}\right)$

Scaling

- Non-dimensional DNS results are scaled using HYFLITS measurements of ϵ (TKE dissipation rate) and ν (Kinematic viscosity)

$$\epsilon_M = 2.9e^{-3} [m^3/s^2] \text{ and } \nu_M = 4e^{-4} [m^2/s] \text{ sampled at } 30 \text{ km}$$

- **Kolmogorov length** and **time** scales are constructed using measured ϵ, ν

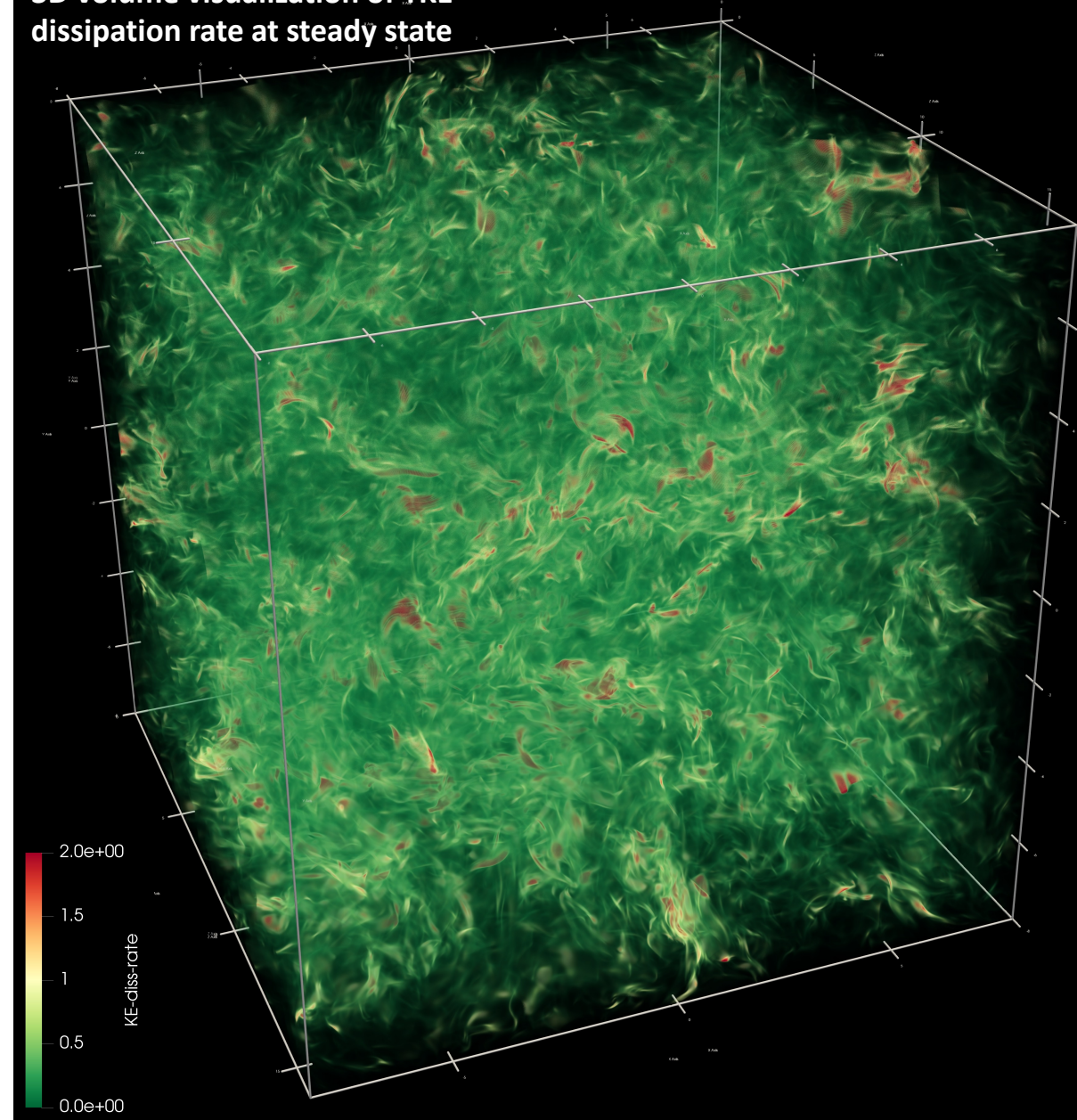
$$\eta_M = \left(\frac{\nu^3}{\epsilon}\right)^{1/4} \text{ and } \tau_M = \left(\frac{\nu}{\epsilon}\right)^{1/2}$$

- $\tilde{\epsilon}_i = \hat{\epsilon}_i C_\epsilon$ $\tilde{u}_i = \hat{u}_i C_u$ $i = 1, 2, 3, \dots N^3$
 [N = number of grid points in X/Y/Z direction; $\hat{}$ **Unscaled DNS**; $\tilde{}$ **scaled DNS**]

- $C_\epsilon = \frac{\epsilon_M}{\langle \hat{\epsilon}_i \rangle}$ $C_\eta = \frac{\eta_M}{\langle \eta_{DNS} \rangle}$ and $C_\tau = \frac{\tau_M}{\tau_{DNS}}$ provide $C_u = \frac{C_\eta}{C_\tau}$

- Scaling results in **box size of 15 [m]** (in X, Y, and Z);
 and **grid resolution of $\delta = 0.02$ [m]**

3D volume visualization of TKE dissipation rate at steady state



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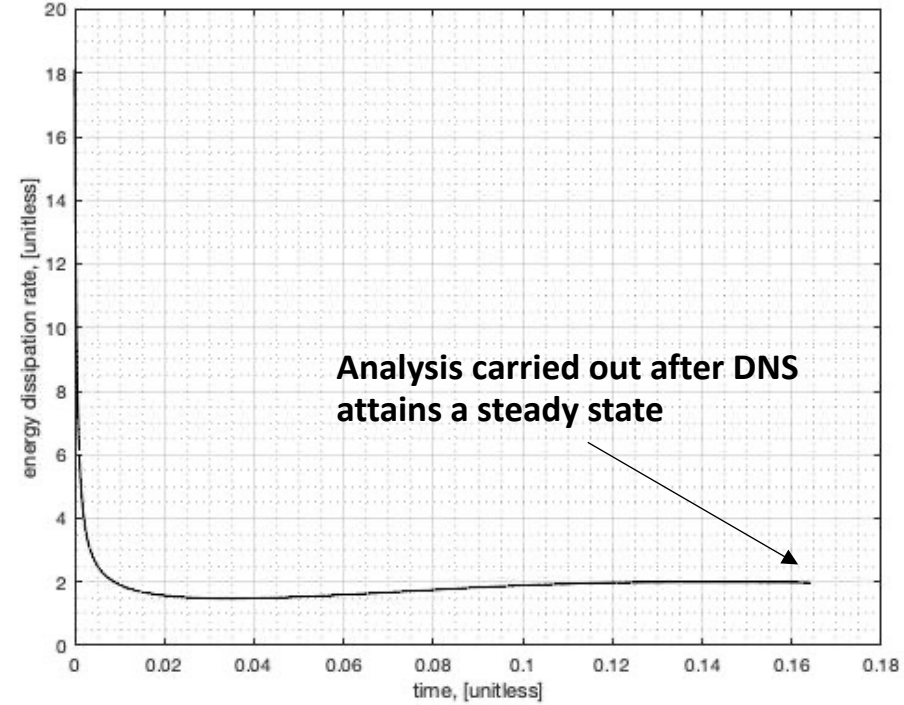
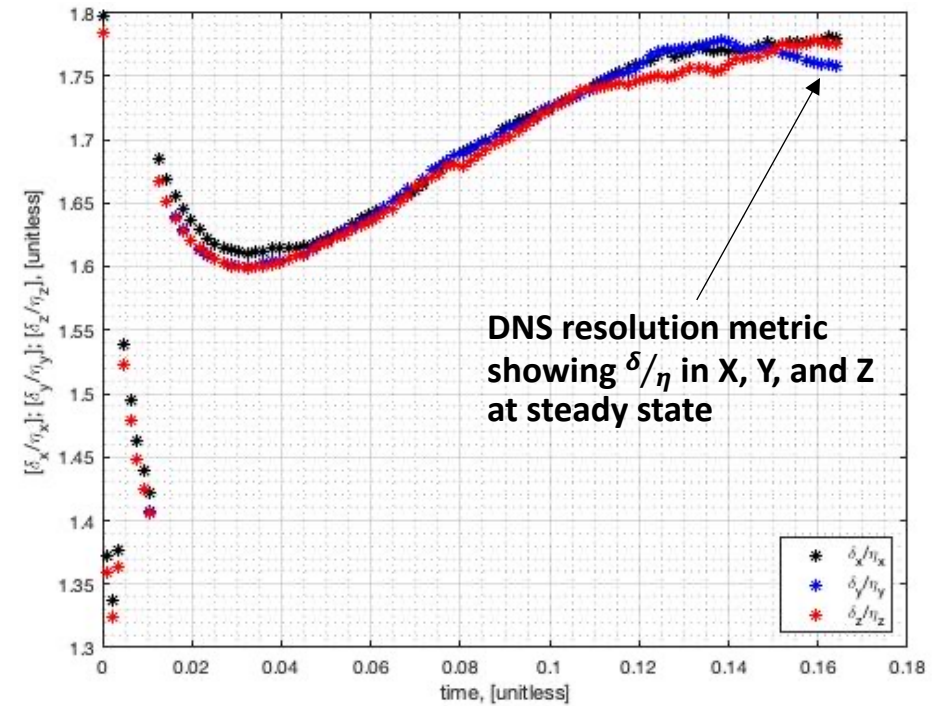
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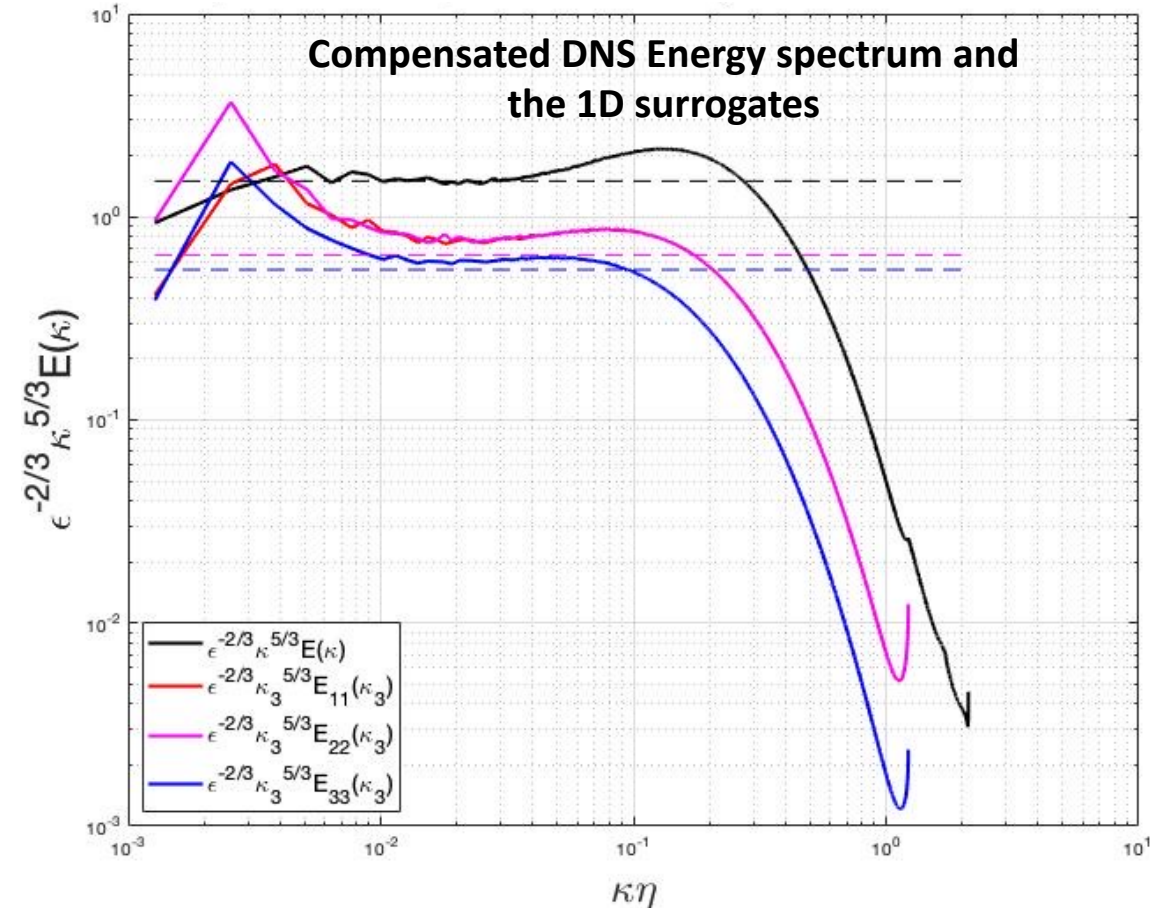
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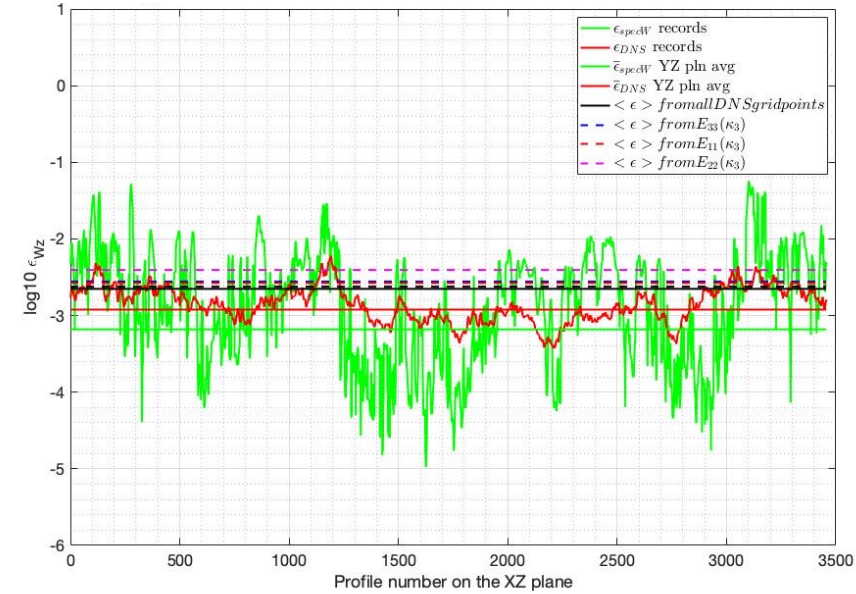
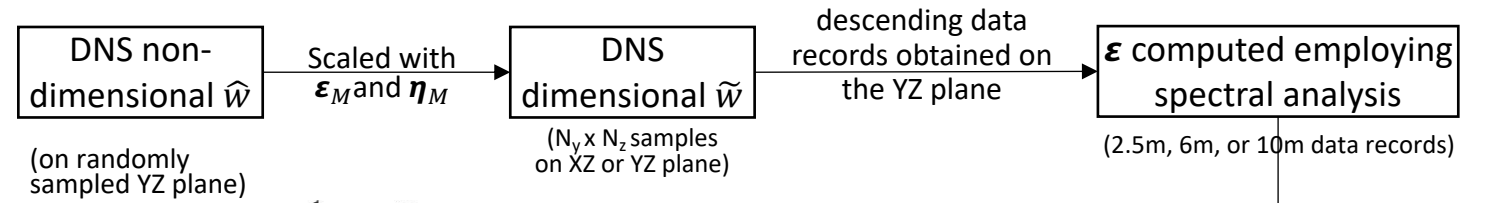
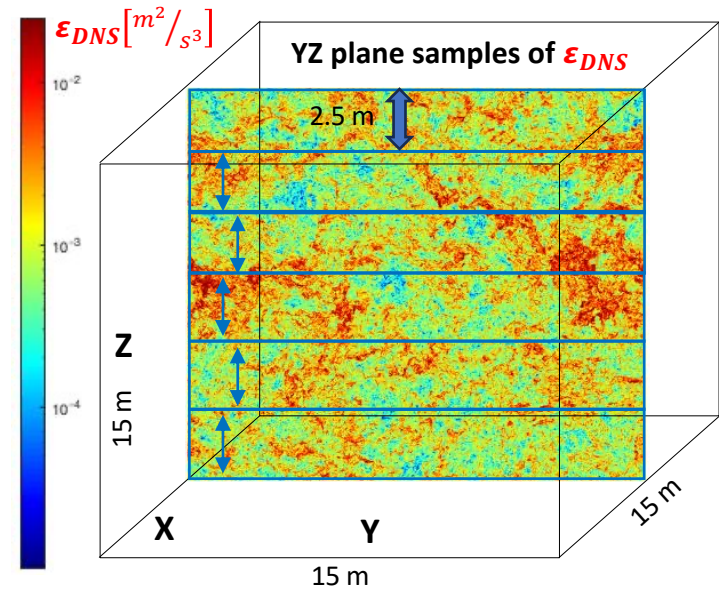
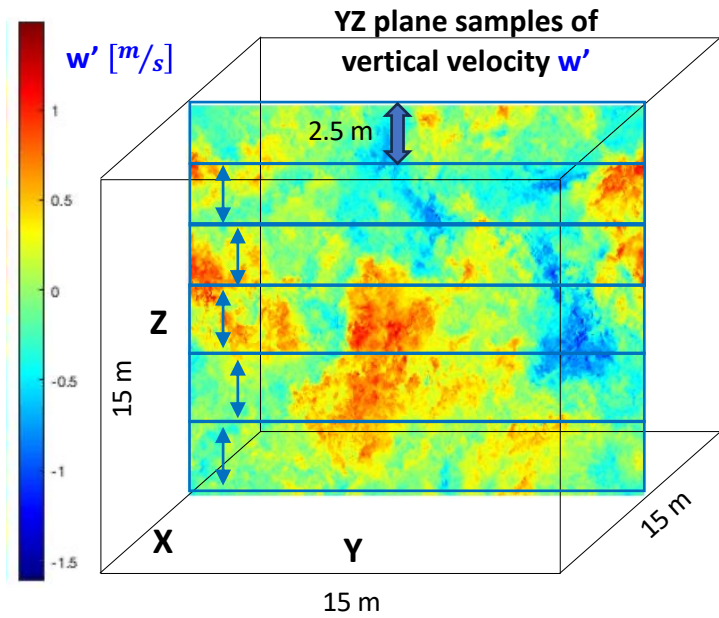


Universal Constants for 1D Energy Spectrum Surrogates

- $E(\boldsymbol{\kappa})$ - Scalar Energy spectrum; $E(\boldsymbol{\kappa}) \sim \alpha \epsilon^{2/3} \boldsymbol{\kappa}^{-5/3}$
- Plotting the 'compensated spectrum' $E(\boldsymbol{\kappa}) \epsilon^{-2/3} \boldsymbol{\kappa}^{5/3}$ against normalized wavenumbers $\boldsymbol{\kappa} \boldsymbol{\eta}$ results in the value for α in the inertial sub-range
- Similar method is adopted to also verify the Kolmogorov universal constants for 1D longitudinal $E_{33}(\boldsymbol{\kappa}_3)$ and transverse $E_{11}(\boldsymbol{\kappa}_3)/E_{22}(\boldsymbol{\kappa}_3)$ surrogate spectra; α_L and α_T respectively (for $\boldsymbol{\kappa}_3$ being the longitudinal direction)
- Experimentally derived values for α , α_L , and α_T reported in literature are 1.5, 0.55, and 0.65 respectively
- Compensated spectra from DNS results in $\alpha = 1.5$, $\alpha_L = 0.59$, and $\alpha_T = 0.7$
- DNS verification of Kolmogorov universal constants show biased estimated for α_L , and α_T

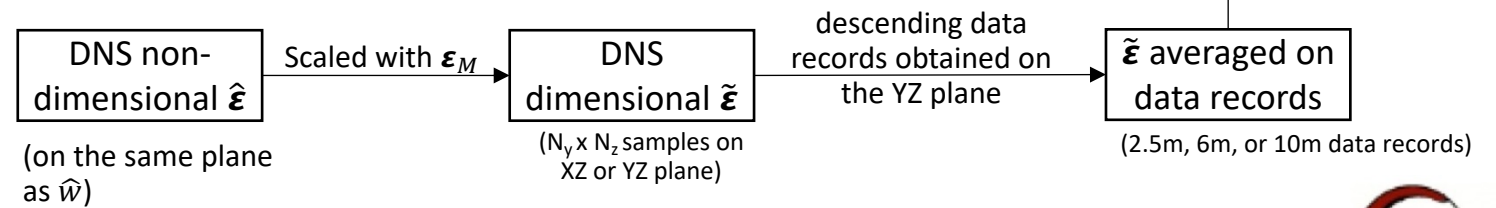


Synthetic Observations - Methodology

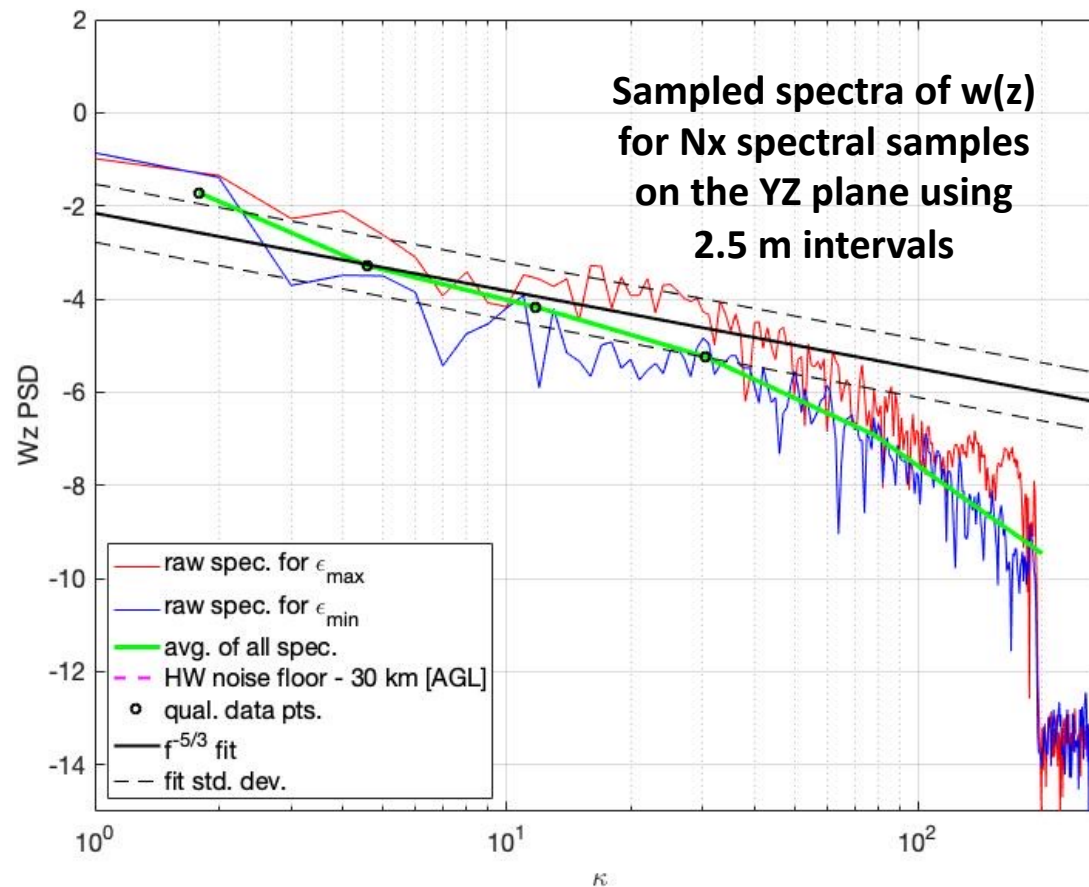
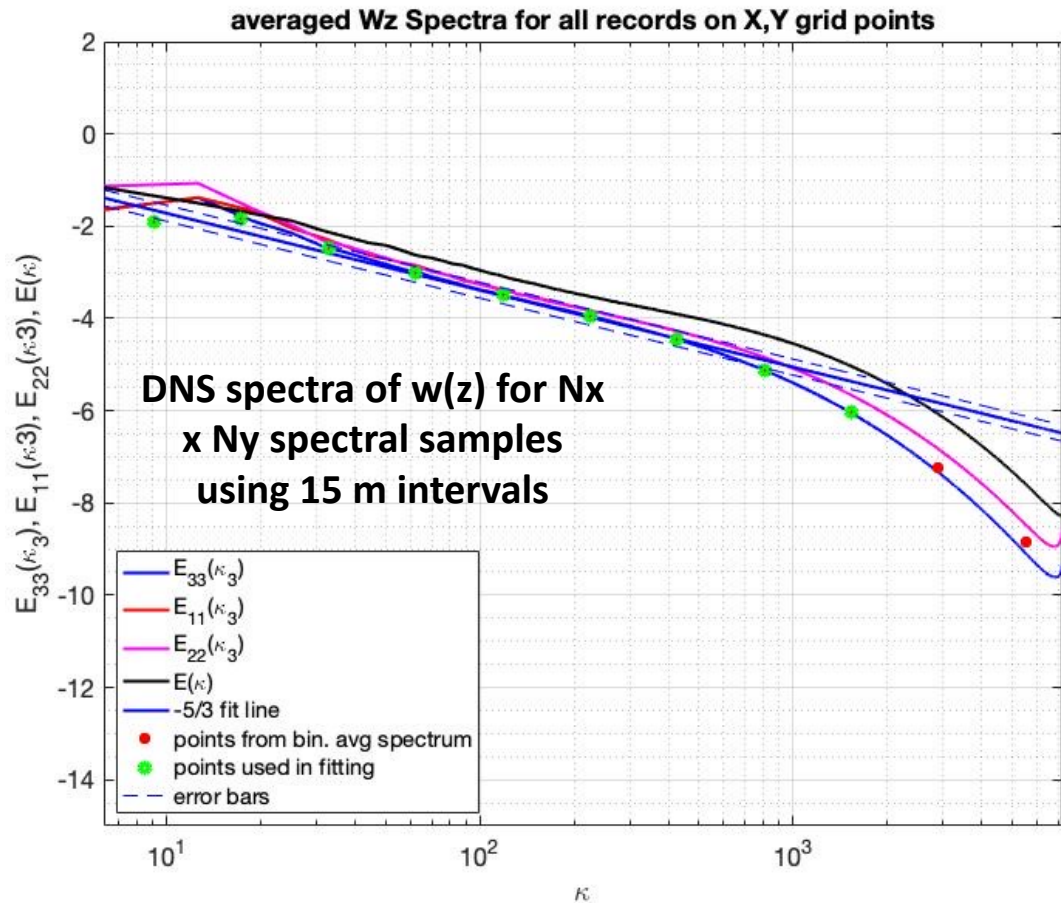


$\tilde{\epsilon}_{spec}$

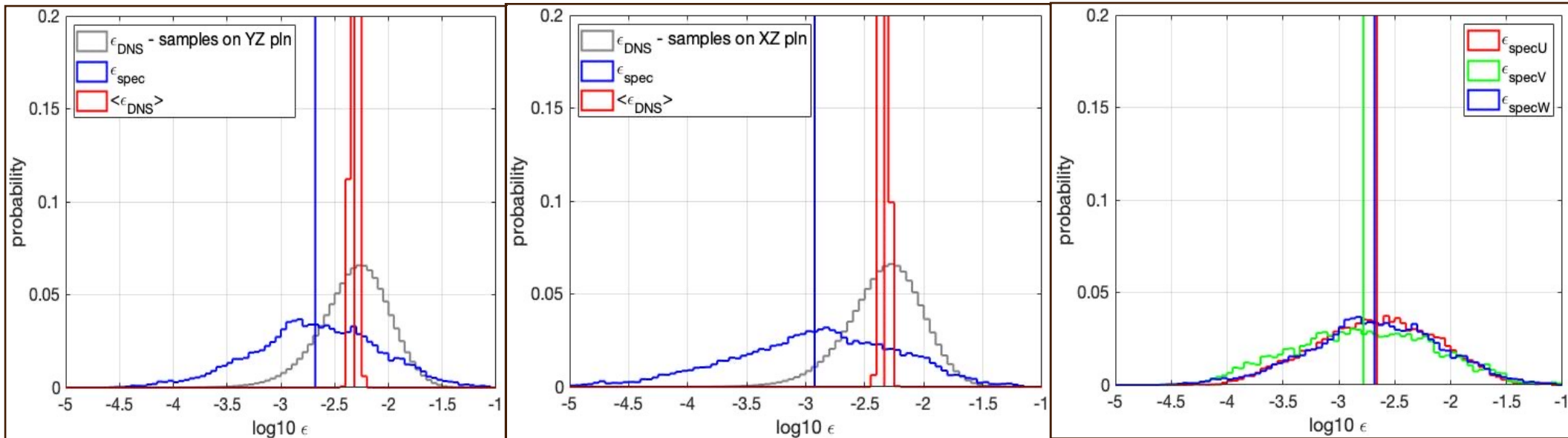
$\langle \tilde{\epsilon}_{DNS} \rangle$



Results

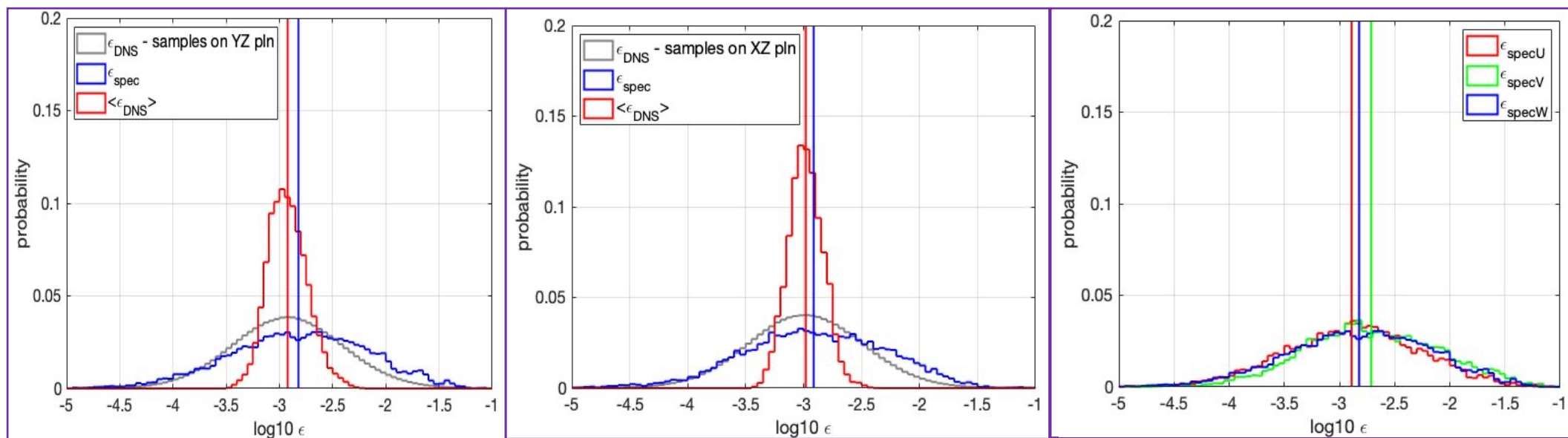


Results



Early-Stage Turbulence

- Higher mean intensity
- Narrow distr.
- Localized
- Spectral method underpredicts mean



Mature Turbulence

- Lower intensity
- Wide distr.
- Spatially more uniform
- Spectral method predicts mean

Preliminary Findings

- **Spectral measurements provide accurate distribution of TKE dissipation rates for mature (stationary homogeneous isotropic) turbulence**
 - Means of $\log_{10} \epsilon_W$ and $\log_{10} \langle \epsilon_{DNS} \rangle$ over 2.5m data records agree to within 0.12 of a decade
 - Variances of $\log_{10} \epsilon_{DNS}$ (not averaged) are similar to variances of $\log_{10} \epsilon_W$ over similar domains
 - Variance of $\log_{10} \langle \epsilon_{DNS} \rangle$ (record average) is significantly smaller than $\log_{10} \epsilon_W$ (spectral average over same record)---thought that this would be the closest comparison!
 - Distributions of ϵ_U , ϵ_V are similar to ϵ_W , indicating absence of significant anisotropy over the spectral data records
 - Averaging of ϵ_U , ϵ_V and ϵ_W reduces variance in spectral estimates only marginally
- **Spectral measurements can underestimate TKE dissipation rates for unsteady (early-stage) turbulence**
 - Mean of $\log_{10} \epsilon_W$ over whole planes shows a bias of 0.5 decade relative to $\log_{10} \epsilon_{DNS}$, with much wider distribution
 - Distributions of ϵ_U , ϵ_V are similar to ϵ_W , indicating absence of significant anisotropy over the spectral data records
 - Bias effect is more pronounced when longer measurement intervals are used (e.g., 15 m records instead of 2 m)
 - Conjecture this is due to spatial averaging of localized strong turbulence (inherent in the spectral method), i.e., lack of homogeneity over spectral data interval length. → Suggests smaller in-situ measurement intervals are better.



Future Work

- Evaluate the influence of Taylor Frozen turbulence hypothesis.
- Compare the distribution of spectrally derived ϵ with second-order structure function derived ϵ .
- The analysis discussed so far neglects to account for the the biases introduced in the ϵ estimates due to observation and instrument errors such as
 - Sensitivity to balloon pendulation effects
 - Sensitivity to instrument/sensor misalignment relative to the longitudinal measurement direction (side-slip)
 - Sensitivity to instrument/sensor noise characteristics.
- Evaluate the sensitivity to wider frequency fitting ranges/time-averaging intervals.
- Investigate the applicability of Thorpe analysis to derive ϵ and comparisons with spectral and second-order structure function methods.

