

**Recent research progress based on the measurements
of high-resolution radiosonde network in China and
the Beijing MST radar observations**

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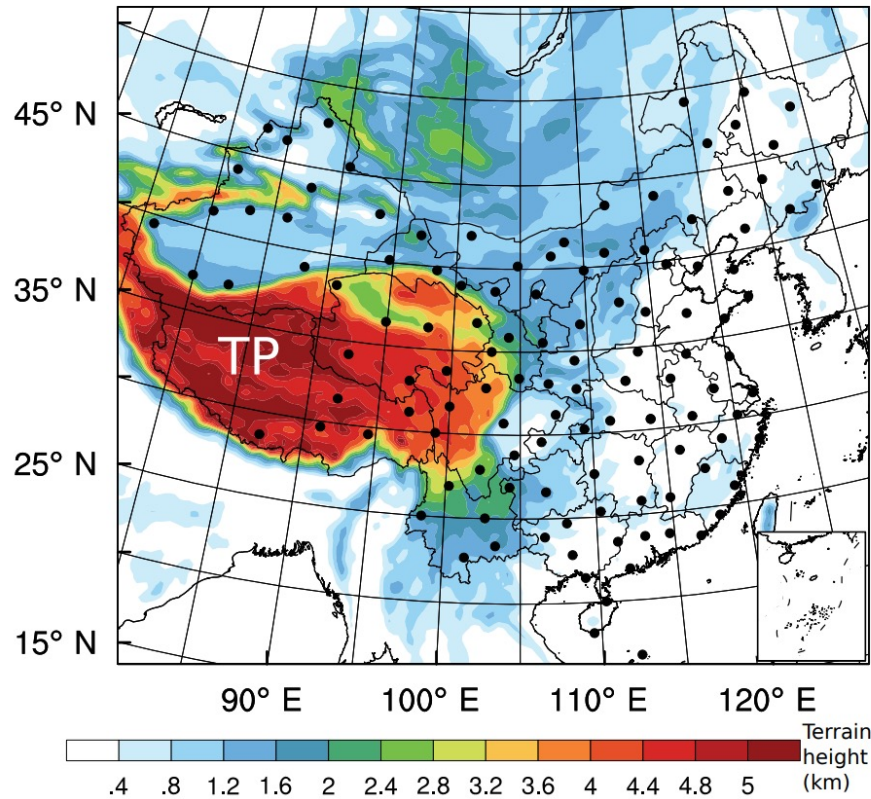
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High-resolution radiosondes in China



Spatial distribution of CMA sounding sites (black dots), overlaid over the terrain height (color shaded) of China

2011, **120** operational radiosonde stations using the L-band sounding systems.

The GTS1 digital electronic radiosonde

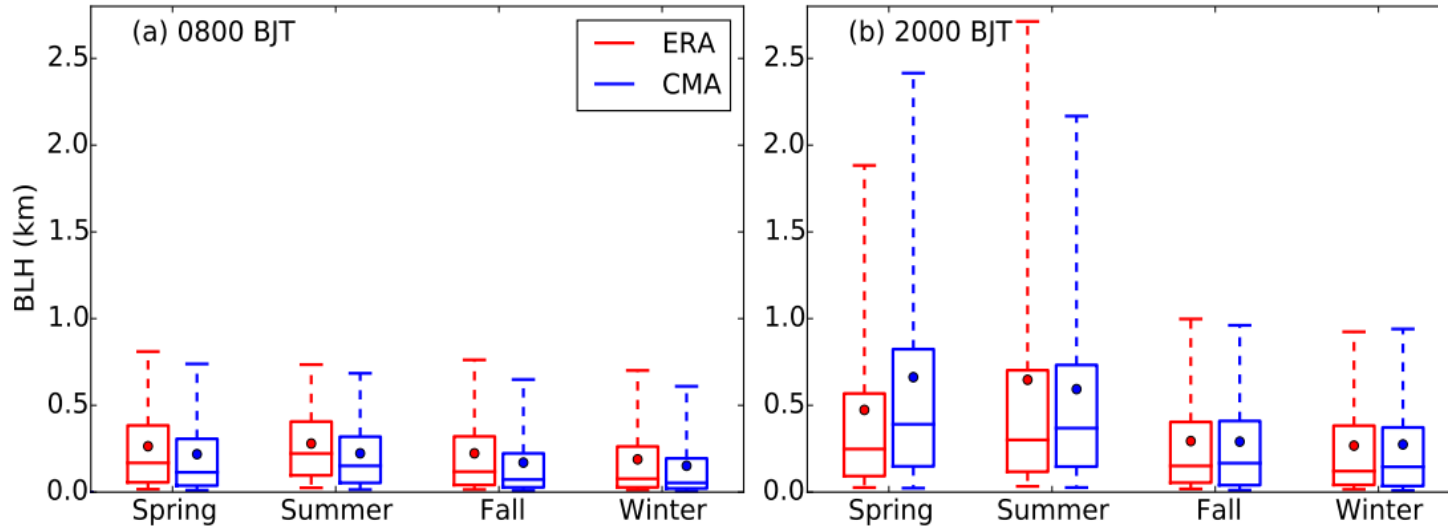
1 January 2011 to 31 July 2015

Time	Profiles
02:00 BJT (summer)	1578
08:00 BJT	190 027
06:00 BJT (summer)	10 313
20:00 BJT	189 634
total	391 552

BLH derived from high-resolution radiosondes: Climatological pattern

- The boundary layer height (BLH) is extracted by the bulk Richardson number (Ri) method. (Vogelezang and Holtslag, 1996; Seidel et al. 2012)

$$Ri(z) = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - u_s)^2 + (bu_*^2)}$$

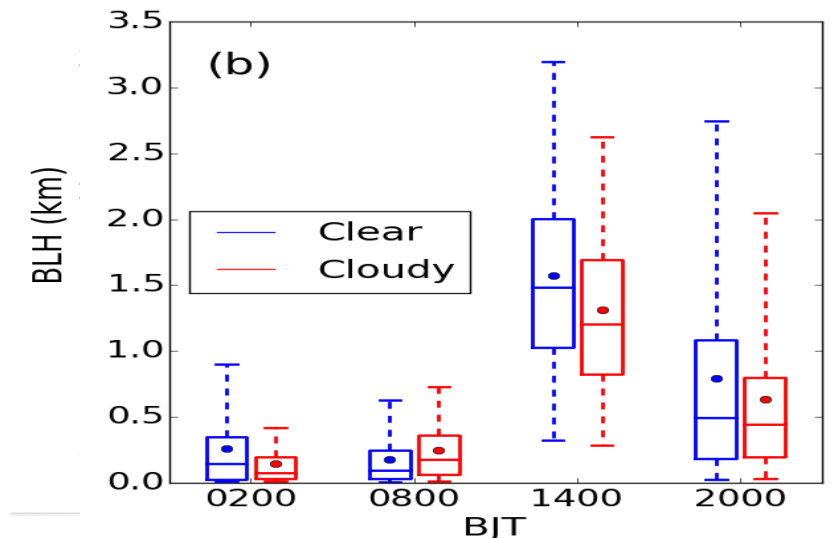


The BLH of **spring and summer** is generally higher.

At **08:00 BJT**, the mean value (dot) of **ERA-BLH > CMA-BLH**

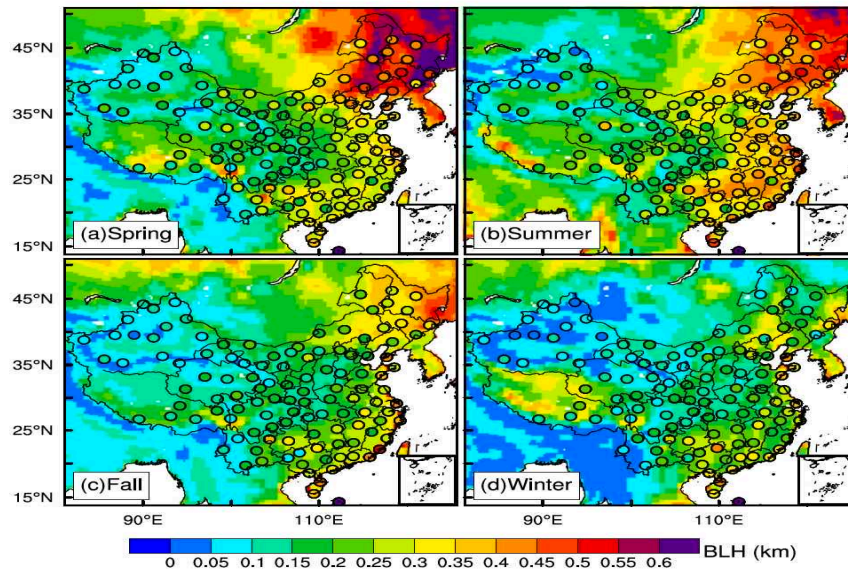
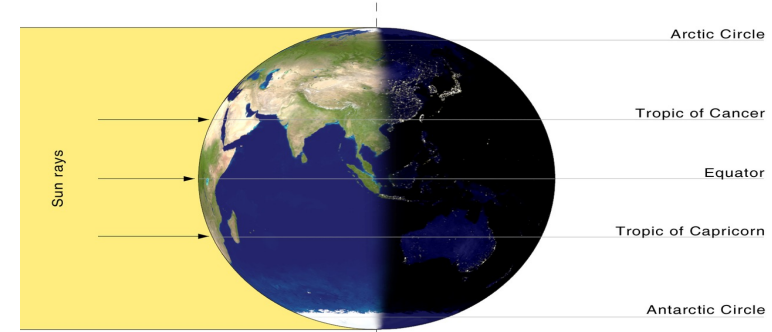
At **14:00 BJT**, the development of PBL is typically **suppressed** due to less solar radiation received at the surface under **cloudy conditions**.

At **20:00 and 02:00 BJT**, **higher** BLHs under **clear** conditions, may be related to the **larger heat storage** of land surface.

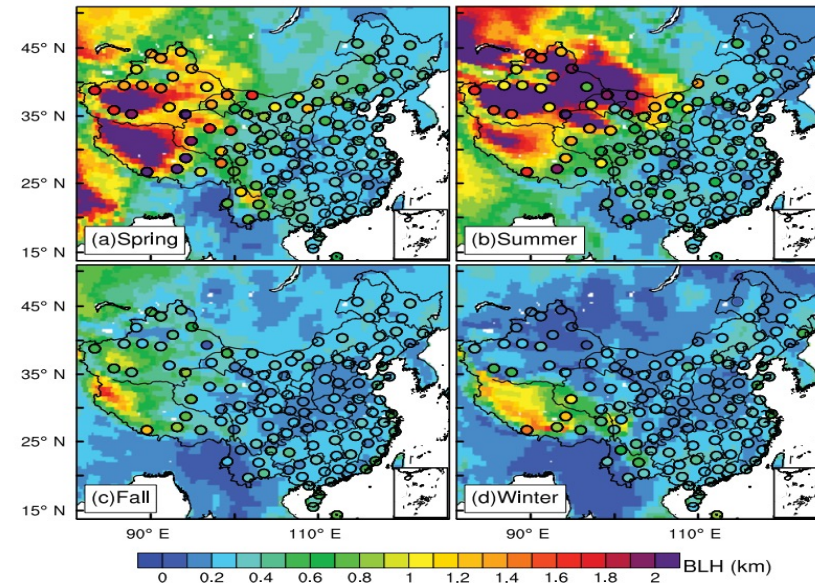


BLH derived from high-resolution radiosondes: Climatological pattern

- The BLH presents the **opposite gradient** spatial distribution at **08:00** and **20:00 BJT**
- These are likely caused by the **differing magnitudes of solar radiation** in the west (at an earlier local sidereal time) and east of China.



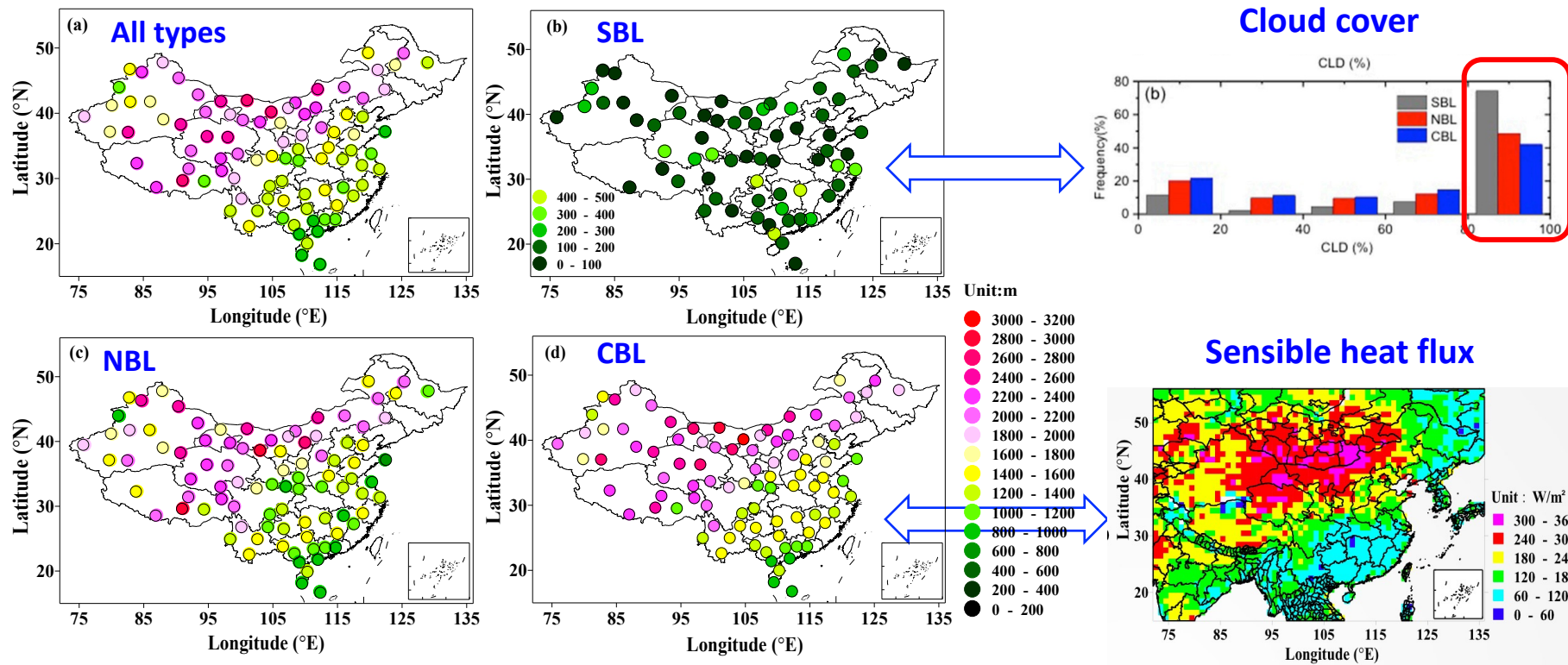
PBL height at 0800 BJT



PBL height at 2000 BJT

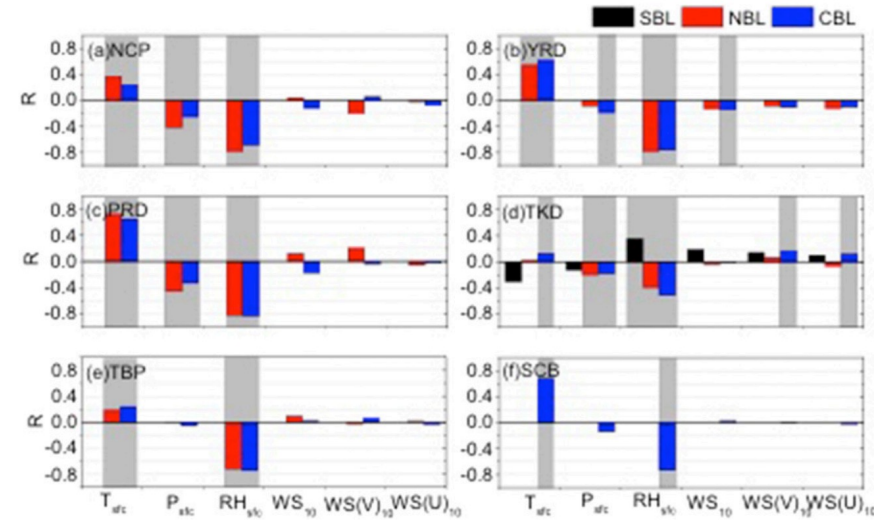
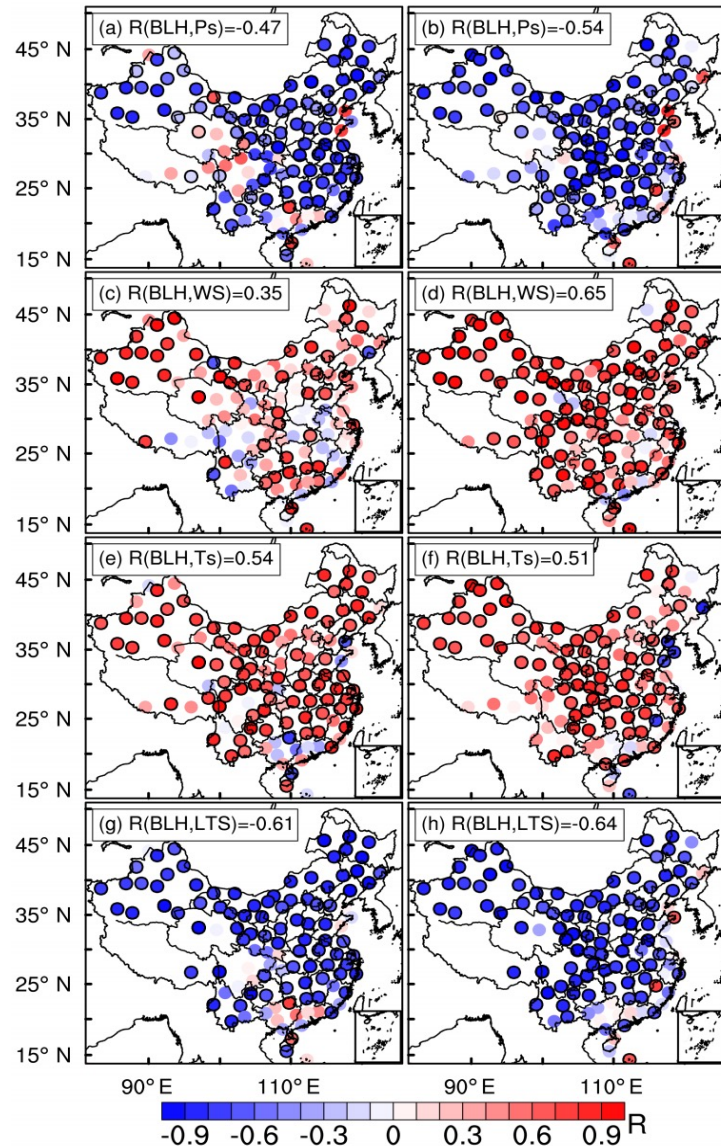
BLH derived from high-resolution radiosondes: Climatological pattern

- Data: Intensive summertime soundings launched at 1400 BJT (2012~2016)
- Method: The PBL with different thermodynamic stability were determined by calculating the near-surface potential temperature difference (PTD).
- A prominent north-south gradient with higher BLH in northwest China.



- The evolution of **SBL** height is mainly determined by **cloud cover**.
- **CBL** height is mainly determined by the observation of **sensible heat flux**.

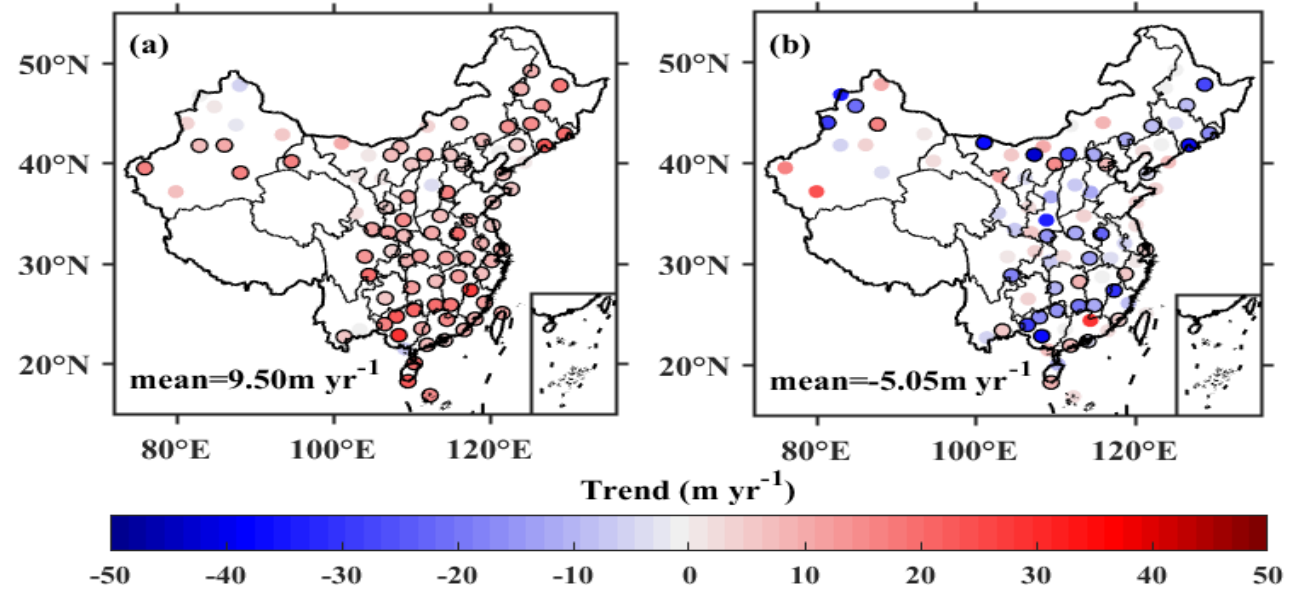
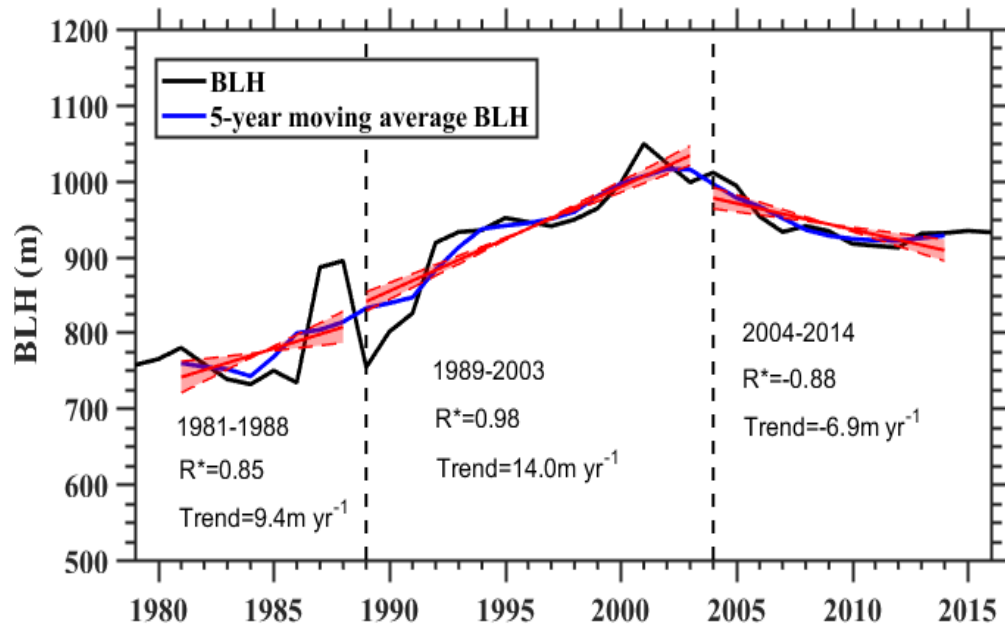
BLH derived from high-resolution radiosondes: Meteorological influence



- This annual cycle of BLH at most sites is found to be **anti-correlated** with the surface pressure, lower tropospheric stability (LTS), and humidity (CBL,NBL).
- And **positively correlated** with the near-surface wind speed and temperature (CBL,NBL), whereas no apparent relationship was found for SBL.

BLH trend derived from radiosondes @14 BJT

- Long-term change trend of BLH in China (1979-2016) is **reversed in the year of 2004**
- BLH increased uniformly at first and decreased inhomogeneously over the latter period



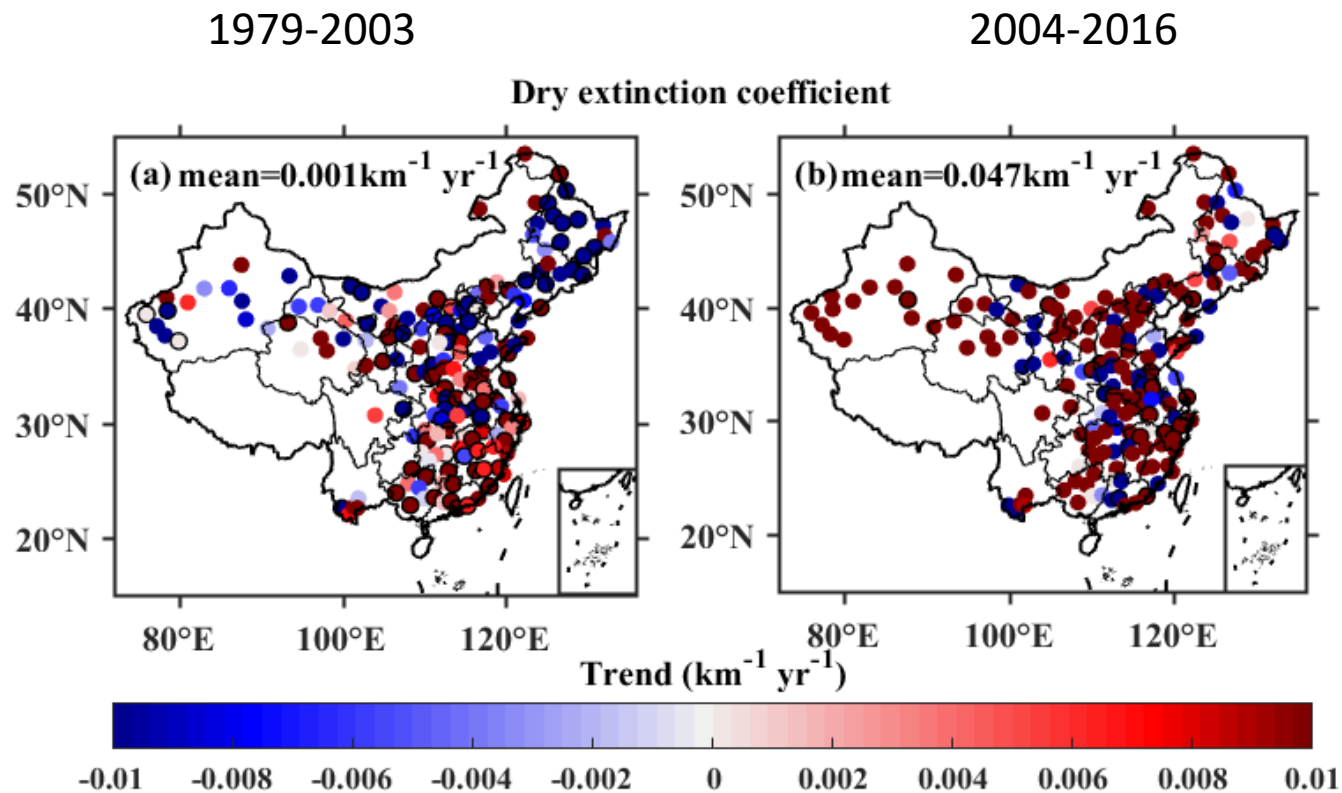
1979-2003

2004-2016

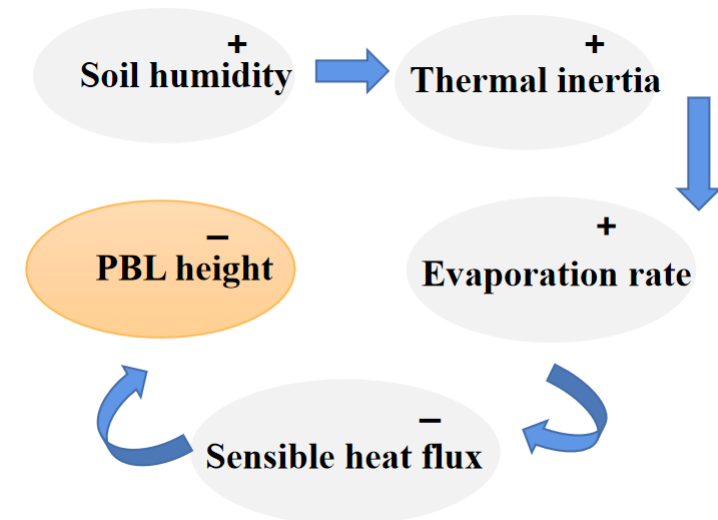
Spatial distr. of BLH trend over China

Potential influential factors for the BLH trend shift

- Aerosol can not explain well the trend shift.
- Soil moisture dominates the trend shift of BLH during the period 1979 - 2016



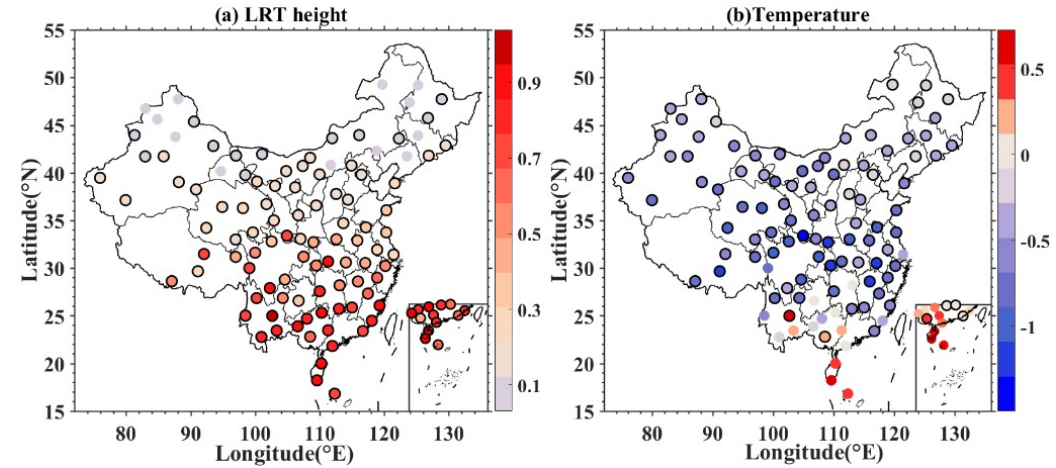
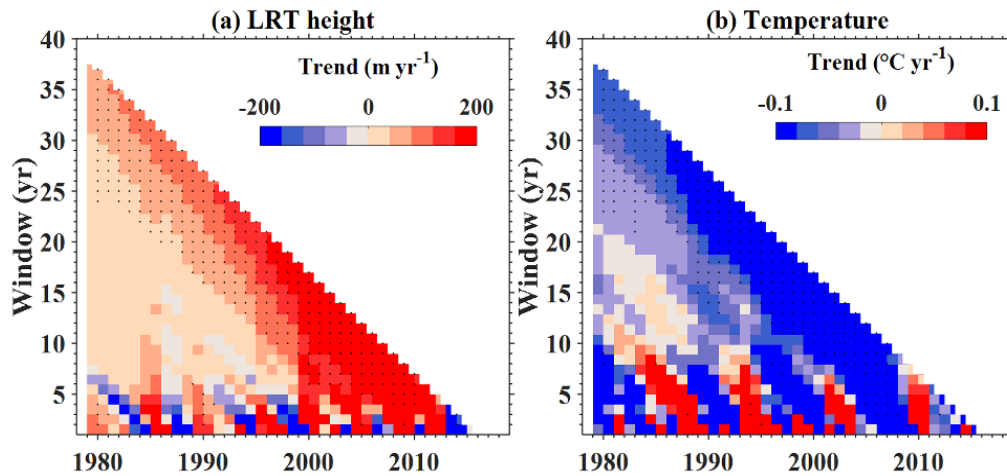
Dry extinction coefficient as derived from visibility was used as surrogate of aerosol loading.



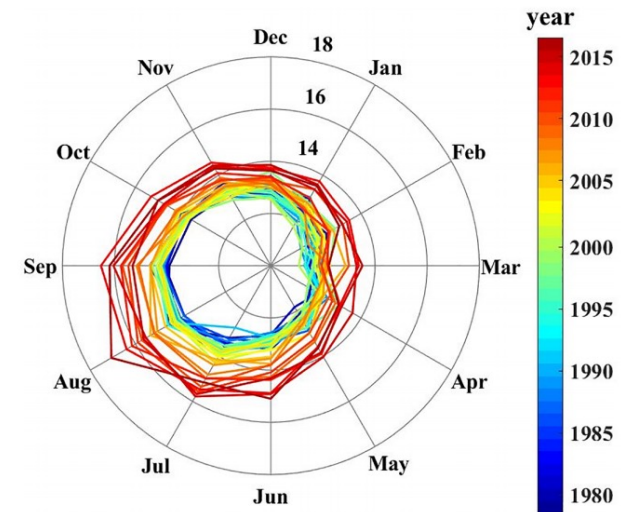
Schematic for the mechanism

Trend of tropopause height in China from radiosondes

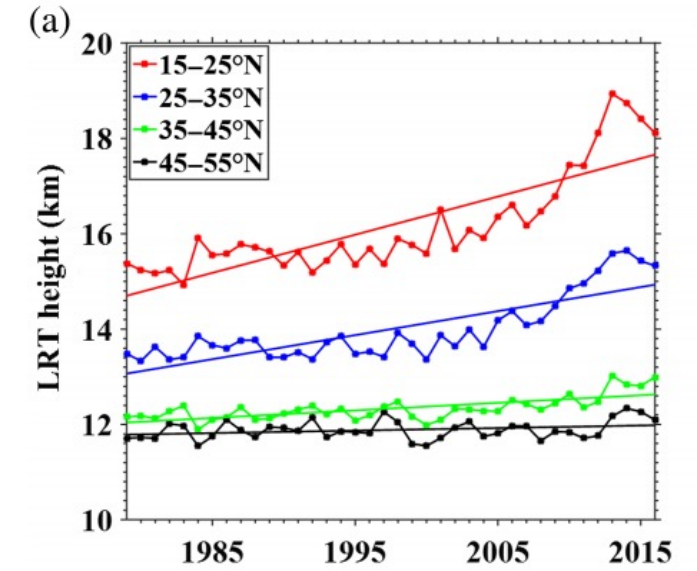
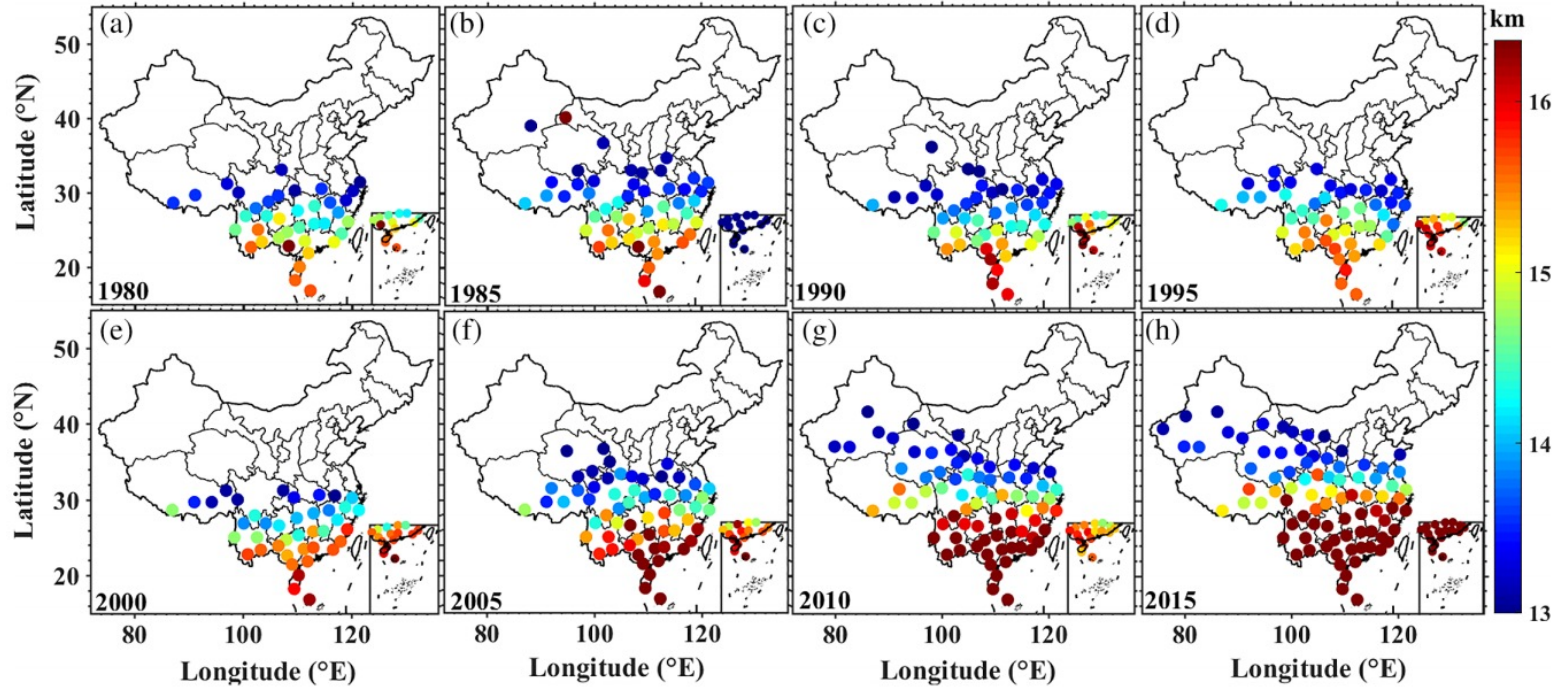
1979–2016 In the background of global climate warming:



- The trend of lapse rate tropopause (LRT) height in most parts of China shows a significant **upwards trend**, with a five-year cycle of fluctuations.
- A clear seasonal variation, growth was **fastest in August**.



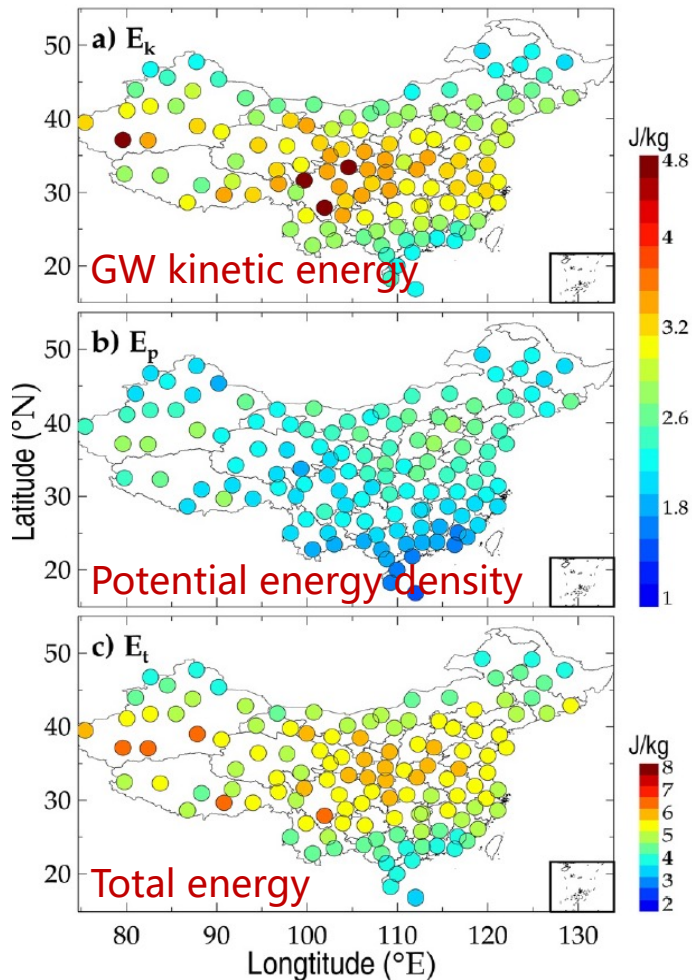
Trend of tropopause height in China from radiosondes



- The LRT height varies rapidly with latitudes, exhibiting a “south high and north low” pattern.
- LRT height over low latitudes is found to be expanding rapidly polewards.

Gravity waves (GW) derived from radiosondes

- Gravity wave (GW) energy is extracted by the **broad-spectral method**. 2016-2019
- Seasonality of E_t is **dependent on latitudes**, triggered by the **shift in the winter subtropical jet**.

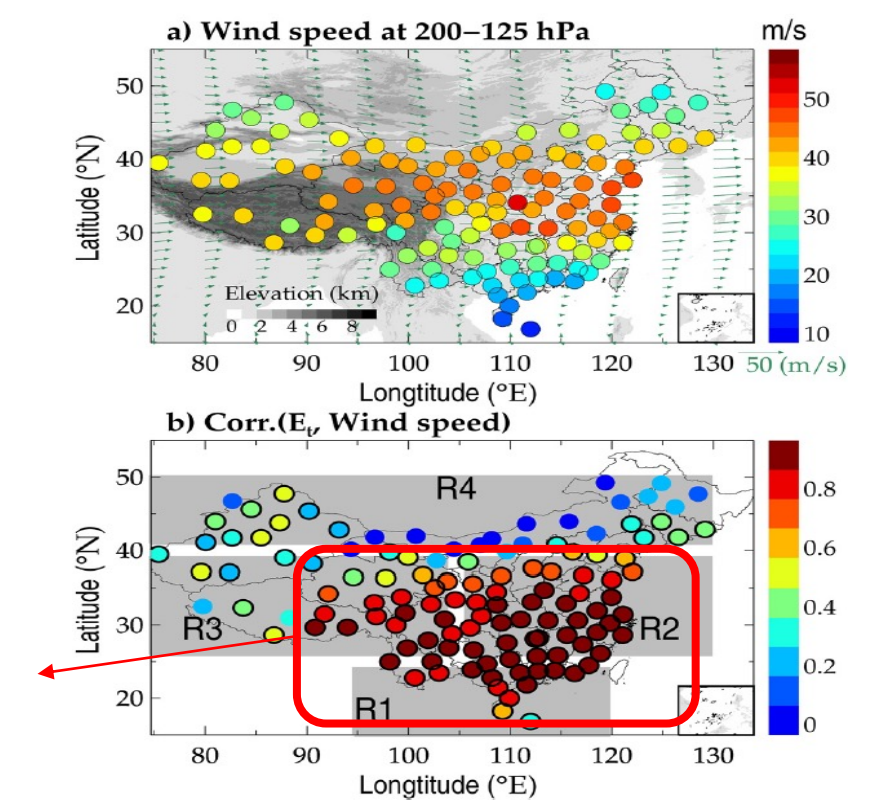


$$E_k = \frac{1}{2} [\overline{u'^2} + \overline{v'^2}]$$

$$E_p = \frac{1}{2} \frac{g^2 \overline{\hat{T}'^2}}{N^2}$$

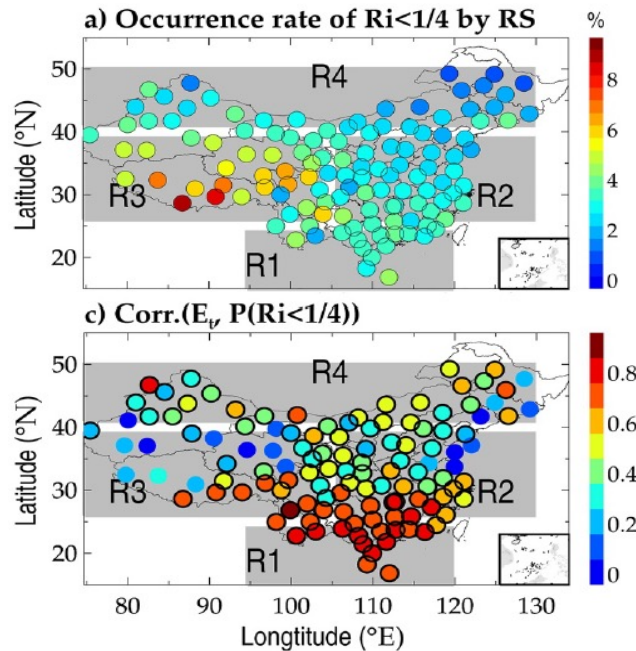
$$E_t = E_k + E_p$$

Highly correlated

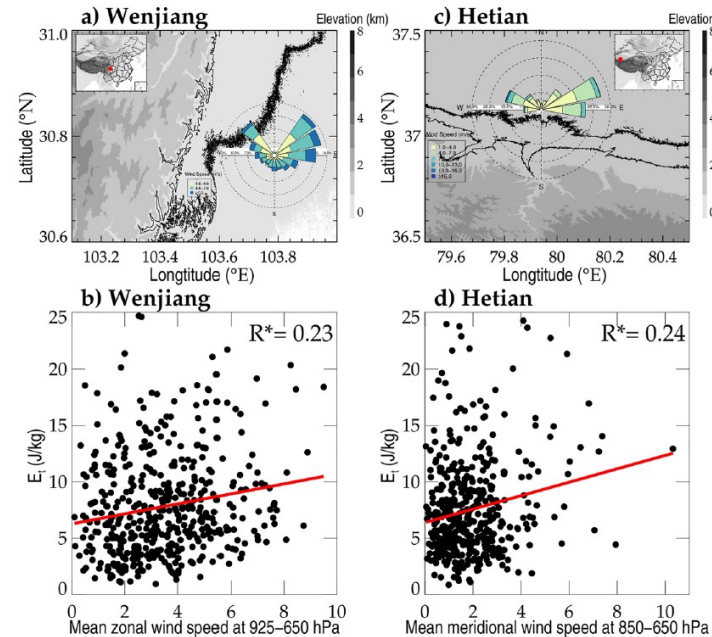


Potential sources for GW generation by random forests regressor

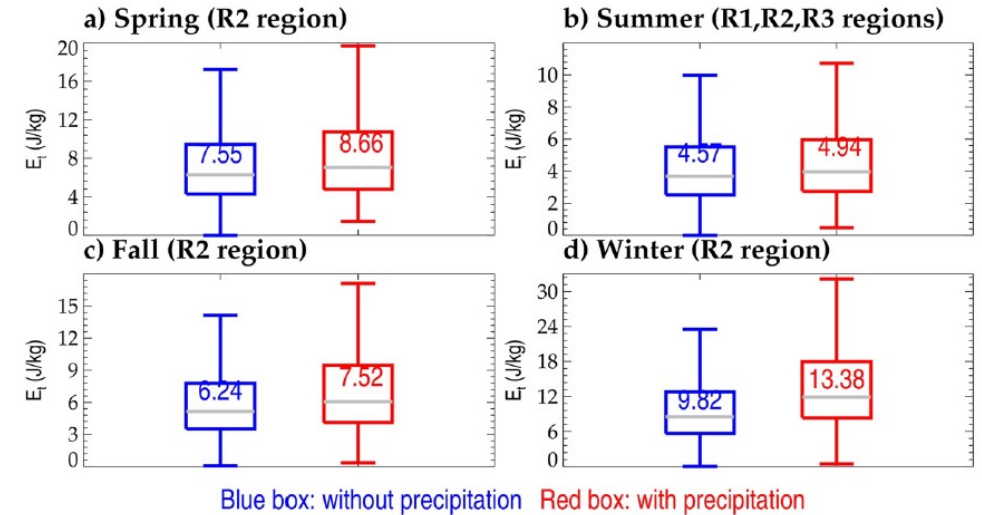
Kelvin-Helmholtz Instability



Terrain-induced flow

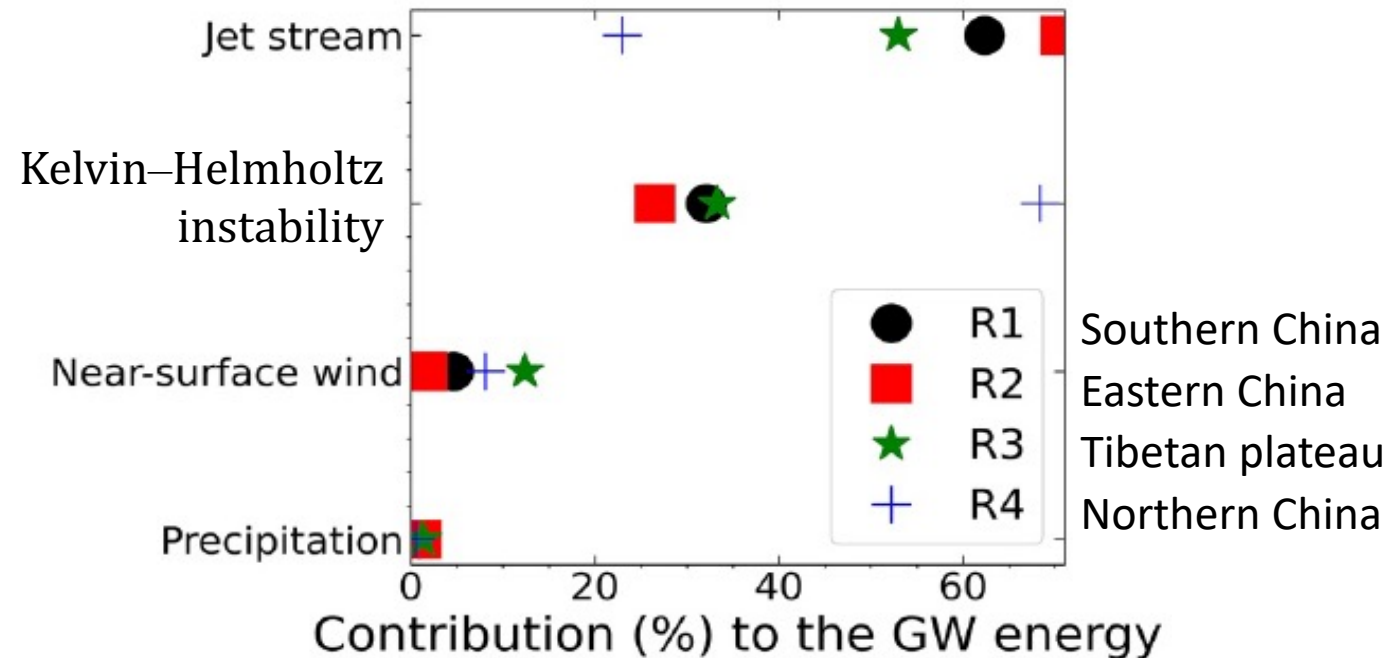


Precipitation



- In addition to the subtropical jet, **Kelvin-Helmholtz instability, terrain induced flow and precipitation** are also sources for GW generation.
- Except for northern China, the contribution to GW enhancement by convective precipitation is $\sim 20\%$ in summer.

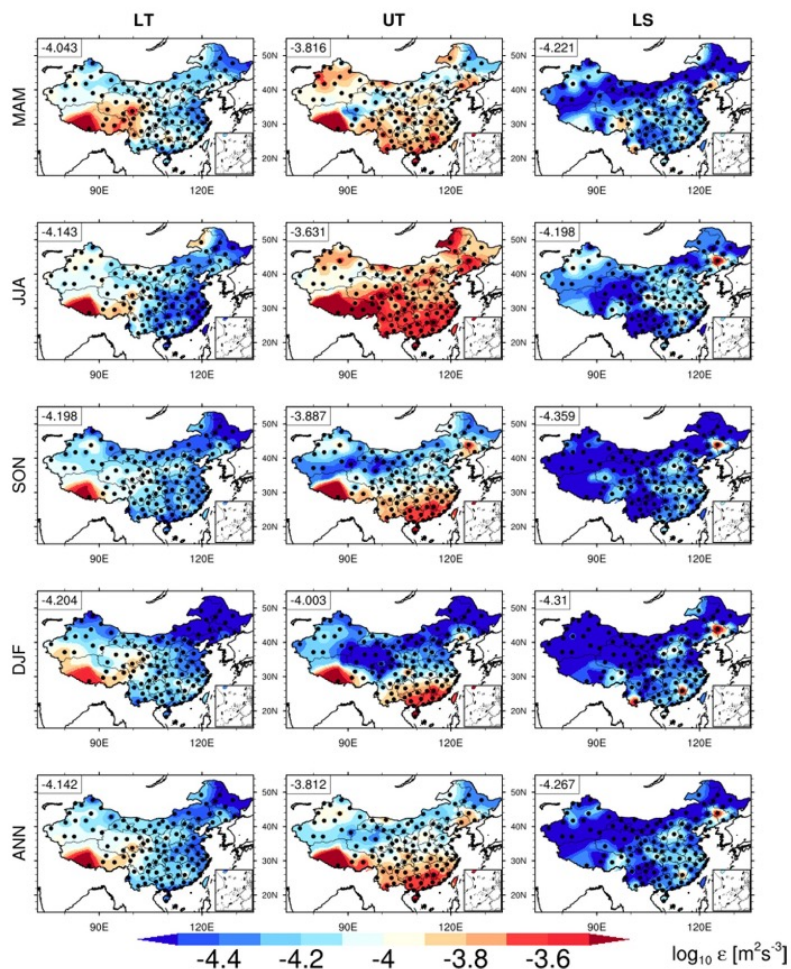
Contributions to GW generation



- **Jet stream** is the **dominant** source for GWs.
- **Enhanced** GW activities favor large **near-surface winds** above **Tibetan plateau**.
- Considerably small contributions of precipitation is due largely to the low frequency of precipitation during nighttime sounding.

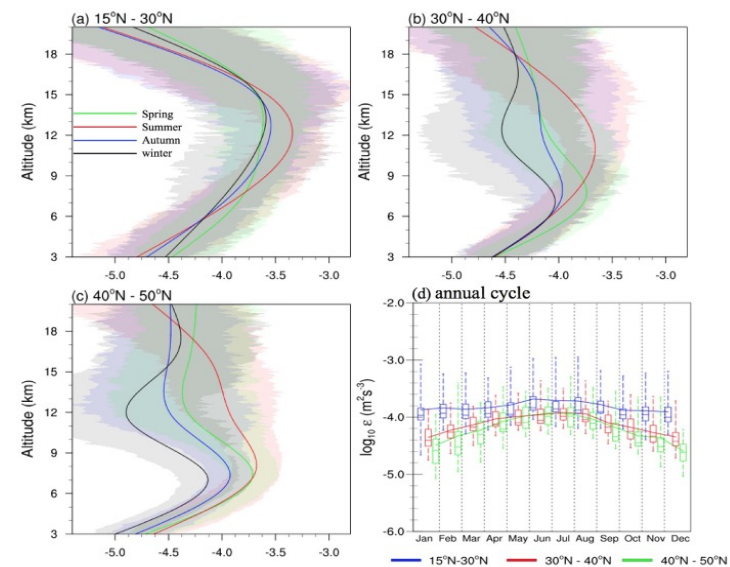
Turbulence from radiosondes: features

- The turbulence dissipation rate(ϵ) was calculated using the Thorpe analysis. **2011–2018**
- The **clear-air ϵ** in the free atmosphere with a ‘**south-high north-low**’ pattern.
- **Large clear-air ϵ** values were observed in both the **LS** and **UT**, especially over the **Tibetan Plateau (TP)**.



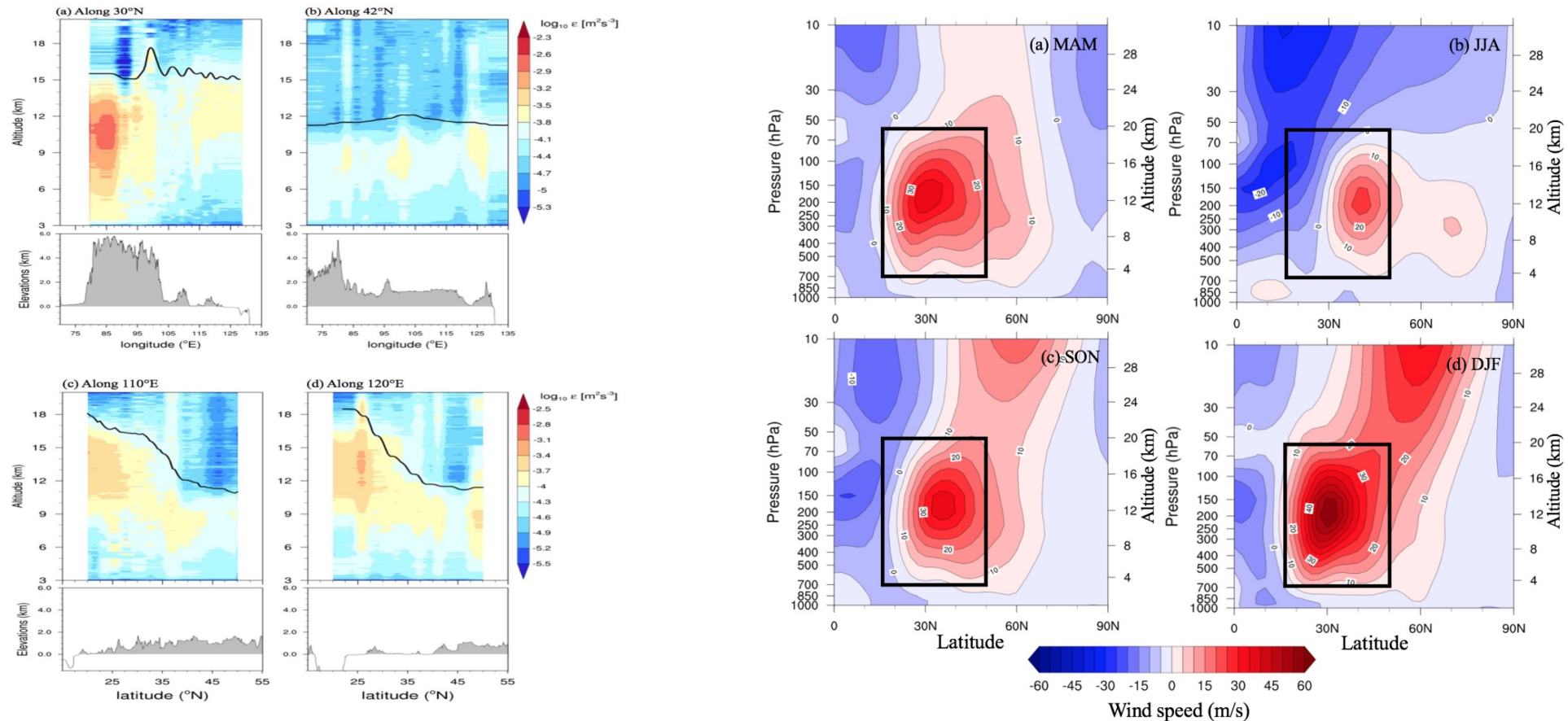
$$\epsilon = C_K L_T^2 N^3$$

$$N = \sqrt{\frac{g}{\theta} \frac{d\theta}{dz}}$$



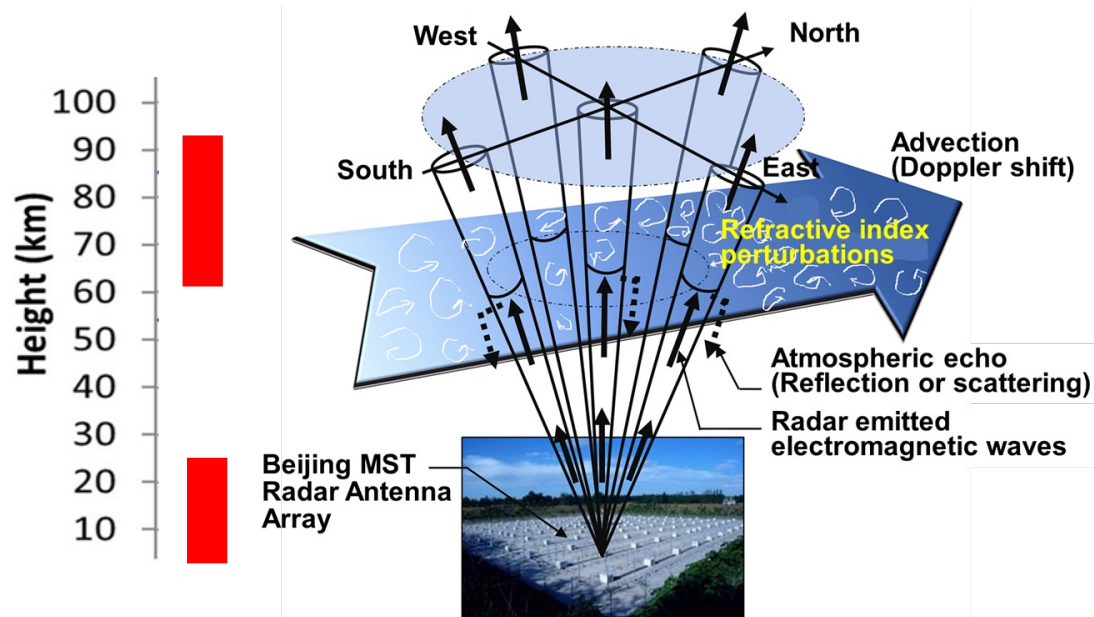
- The seasonality of ϵ was also pronounced, reaching **maxima in summer** and **minima in winter**.

Turbulence from radiosondes: influential factors



- Areas with large undulations in the terrain have strong turbulence.
- In the vertical direction, the altitude of peak clear-air ϵ in the troposphere was found to decrease poleward.
- Strong shear instability around westerly jet streams is an important source of turbulence.

Beijing Mesosphere-Stratosphere-Troposphere (MST) radar



2011 100m×100m

Provide continuous **high time-height resolutions and quasi-simultaneous** observations of the **horizontal wind** and **vertical velocities** of different height ranges.

Powerful tool for investigating **various atmospheric dynamics** (mean flow, tides, planetary waves, gravity waves, turbulence et al.) in the **lower, middle and upper atmosphere**.

Methods used to estimate turbulence parameters

$$\sigma_{tur}^2 = \sigma_{obv}^2 - (\sigma_{shear}^2 + \sigma_{beam}^2)$$

Estimation models of $\sigma_{shear}^2 + \sigma_{beam}^2$

$\sigma_{shear}^2 + \sigma_{beam}^2$ { Horizontal wind
Vertical shear

$\sigma_{tur}^2 < 0$:
Negative turbulent
kinetic energy (N-TKE).

H model Hocking (1983,1985)

$$\sigma_{vb} = \sigma_{f\frac{1}{2}b} \cdot \frac{\lambda}{2} / (\sqrt{2\ln 2}) = (1.0) * 2/\lambda\theta_{1/2}V \cdot \frac{\lambda}{2} / \sqrt{2\ln 2} = (1.0) * \theta_{1/2}V / \sqrt{2\ln 2}$$

$$\sigma_{vs} = \sigma_{v\frac{1}{2}s} / (\sqrt{2\ln 2}) = \frac{1}{2} \cdot \left| \frac{\partial u}{\partial z} \right| \sin(\chi) \Delta r / (\sqrt{2\ln 2})$$

N-2D model Nastrom (1997)

$$\sigma_{shear+beam}^2 = \frac{\theta^2}{3} v^2 \cos^2 \chi - \frac{2\theta^2}{3} \sin^2 \chi \left(v \frac{\partial v}{\partial z} r \cos \chi \right)$$

$$+ \frac{\theta^2}{24} (3 + \cos 4\chi - 4\cos 2\chi) \left(\frac{\partial v}{\partial z} \right)^2 r^2 + \left(\frac{\theta^2}{3} \cos 4\chi + \sin^2 \chi \cos^2 \chi \right) \left(\frac{\partial v}{\partial z} \right)^2 \frac{\Delta r^2}{12}$$

D-H model Dehghan and Hocking (2011)

$$\sigma_{shear+beam}^2 = \frac{\theta^2}{k} v^2 \cos^2 \chi - a_0 \frac{\theta}{k} \sin \chi \left(v \frac{\partial v}{\partial z} \zeta \right) + b_0 \frac{2\sin^2 \chi}{8k} \left(\frac{\partial v}{\partial z} \zeta \right)^2$$

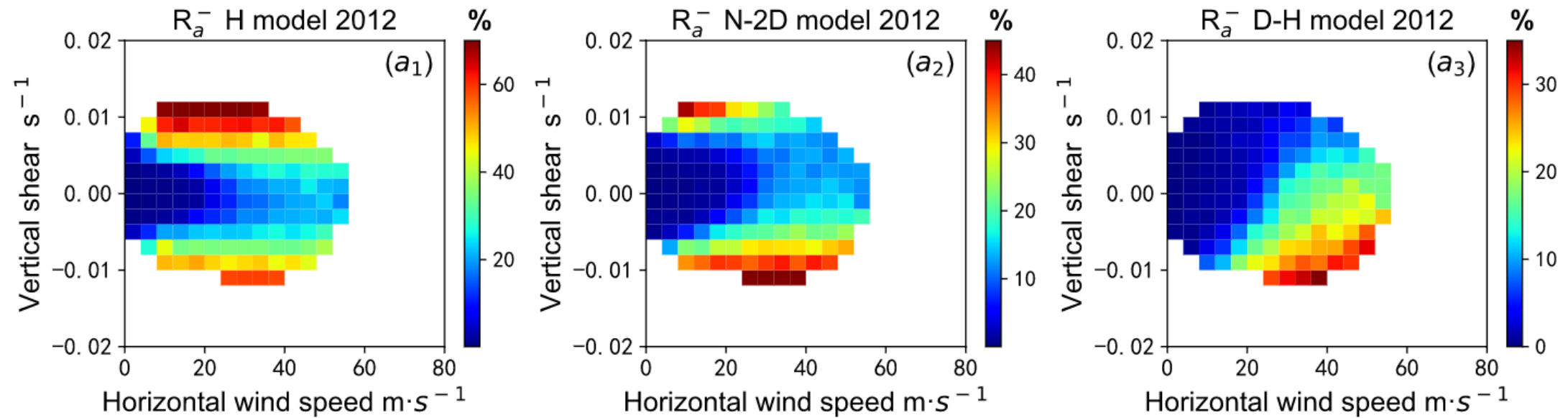
$$+ c_0 (\sin^2 \chi \cos^2 \chi) |v \xi| + d_0 (\sin^2 \chi \cos^2 \chi) \xi^2$$

The turbulent kinetic energy dissipation rate $\varepsilon = A^{-3/2} N \sigma_{tur}^2$

The vertical eddy diffusion coefficient $K_z = 0.15 N^{-1} \sigma_{tur}^2$

The applicability of three models in different wind field conditions

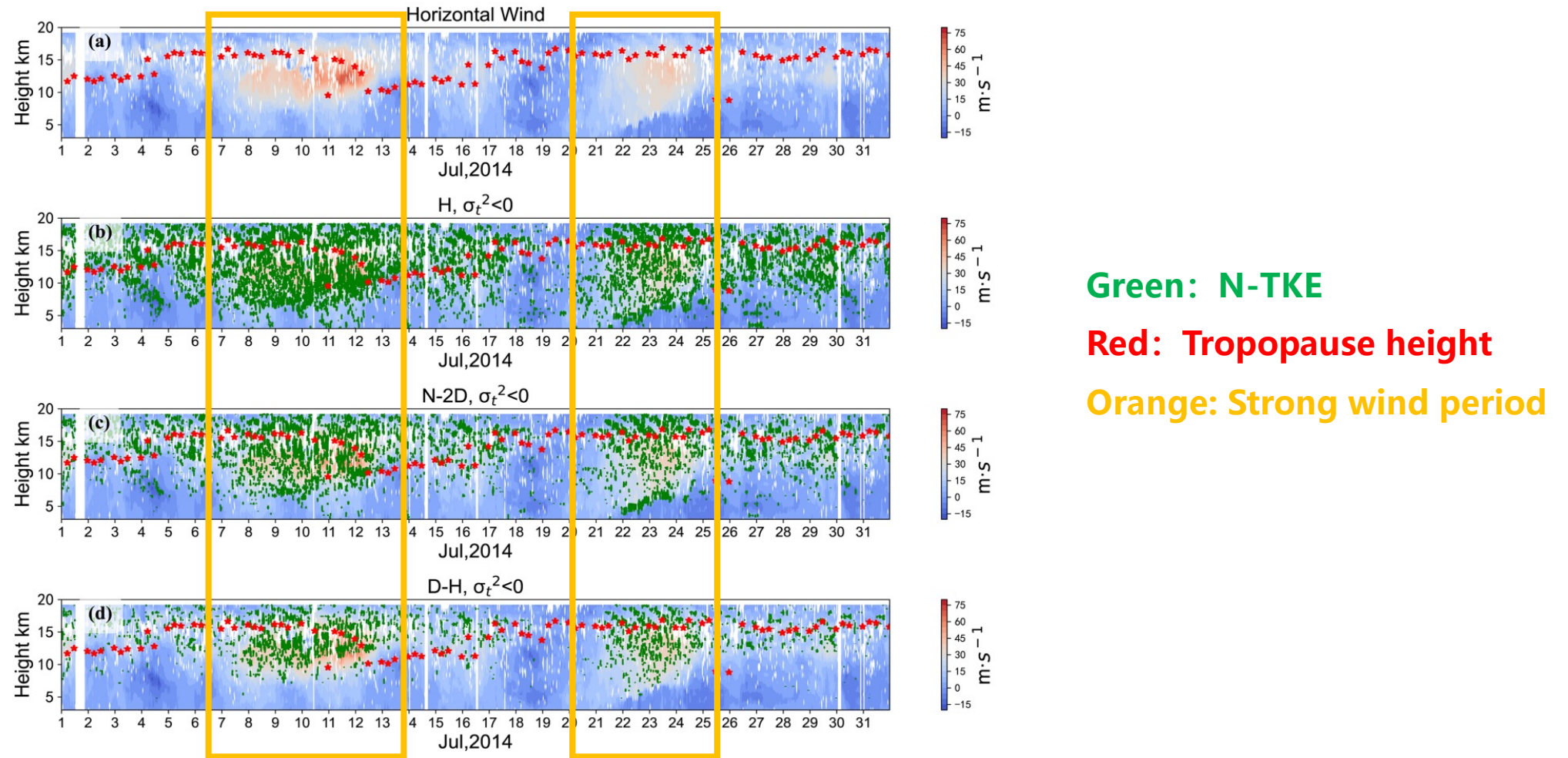
Data: Beijing MST radar 2012-2014 37 000 profiles



The proportion of N-TKE in the H model, N-2D model and D-H model **increases with the horizontal wind speed** and the **vertical shear** of horizontal wind speed.

The maximum values are 80 %, 45 % and 35 %, respectively.

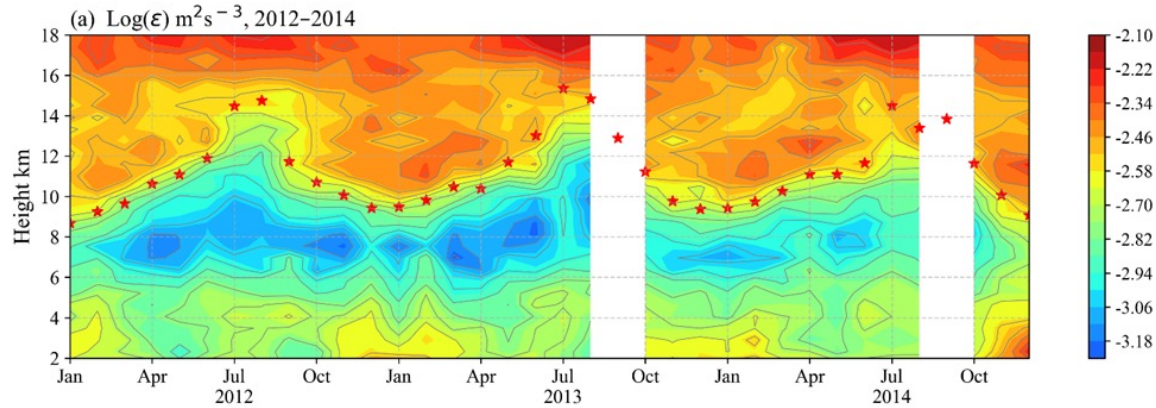
The applicability of three models in different wind field conditions



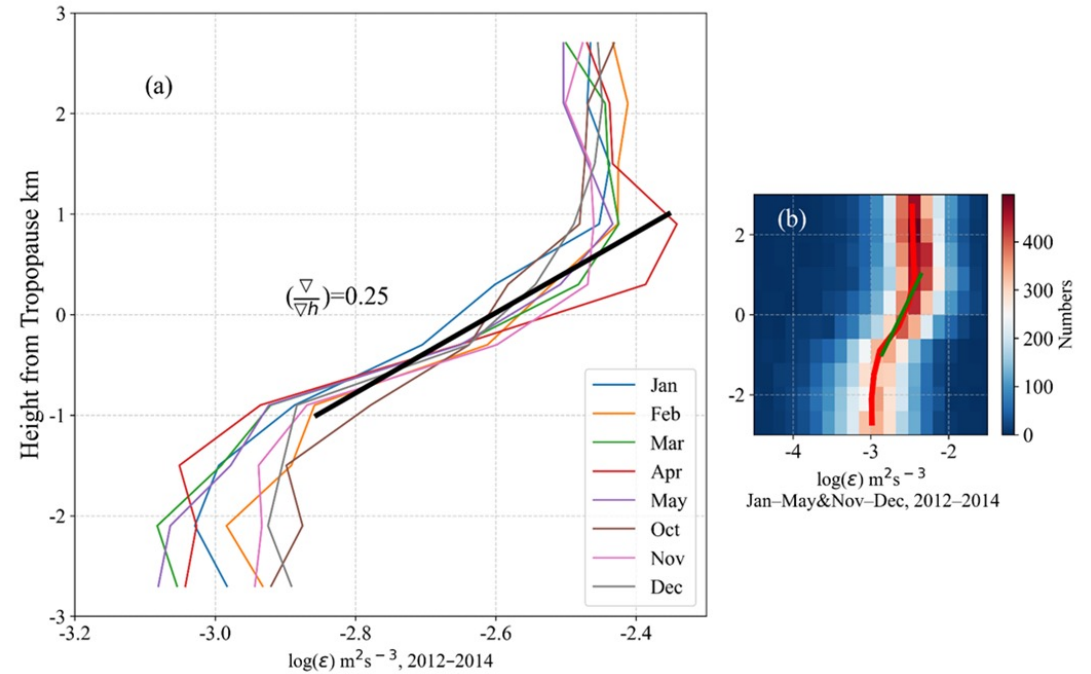
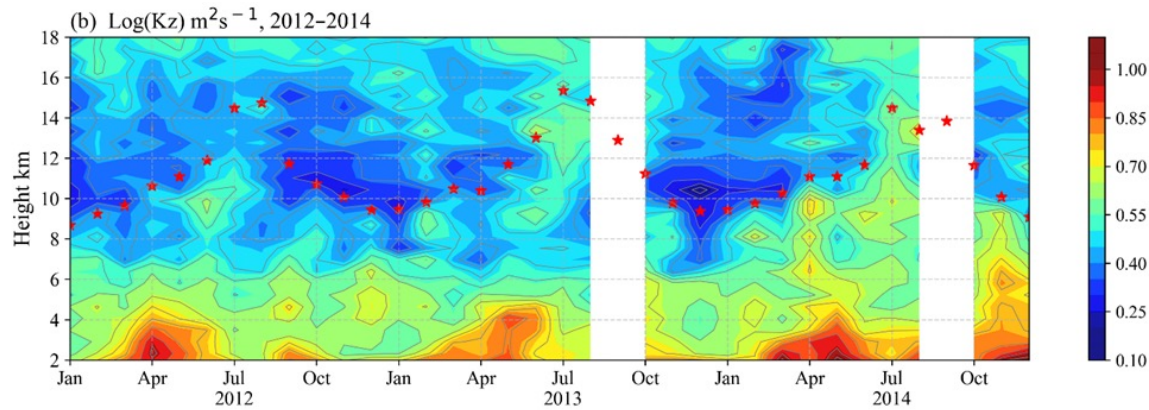
It is still necessary to consider the applicability of the N-2D model and D-H model in some weather processes with strong winds.

Distribution characteristics of the turbulence parameters in the troposphere-lower stratosphere

The turbulent kinetic energy dissipation rate

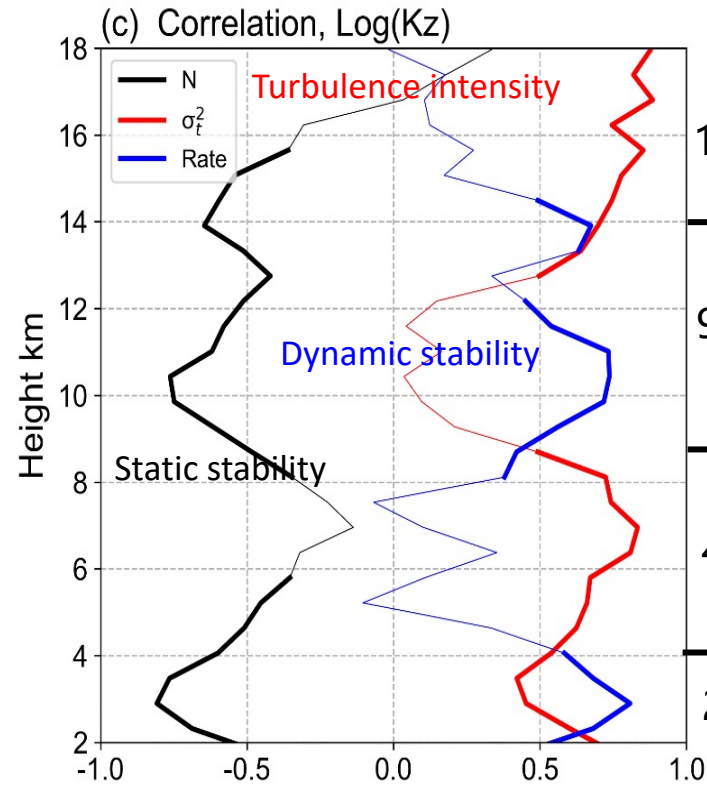
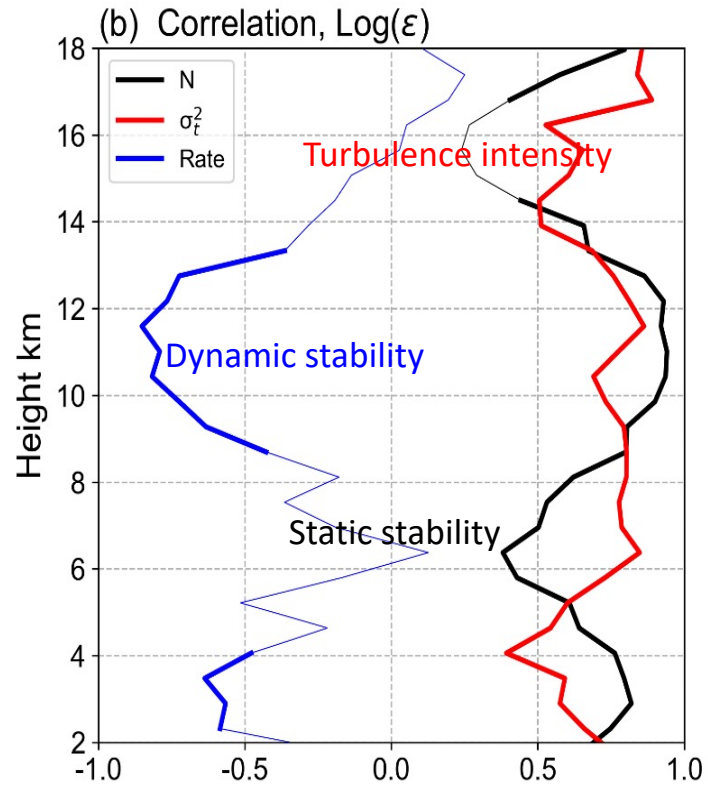


The vertical eddy diffusion coefficient



The **seasonal variation** of turbulence parameters has **noticeable differences** at **different atmospheric layers**.

The atmospheric static/dynamic stability and turbulence intensity influence the distribution of the turbulence parameters



14~18 km: Turbulence intensity

9~14 km } ϵ stability, turbulence intensity
 } K_z stability

4~9 km: Turbulence intensity

2~4 km: Static/dynamic stability

Related papers

- Guo, J. P., Y. C. Miao, Y. Zhang, H. Liu, Z. Q. Li, W. C. Zhang, J. He, et al. "The Climatology of Planetary Boundary Layer Height in China Derived from Radiosonde and Reanalysis Data." *Atmospheric Chemistry and Physics* 16, no. 20 (Oct 2016): 13309-19. <https://doi.org/10.5194/acp-16-13309-2016>.
- Zhang, Wanchun, Jianping Guo, Yucong Miao, Huan Liu, Yu Song, Zhang Fang, Jing He, et al. "On the Summertime Planetary Boundary Layer with Different Thermodynamic Stability in China: A Radiosonde Perspective." *Journal of Climate* 31, no. 4 (Feb 2018): 1451-65. <https://doi.org/10.1175/jcli-d-17-0231.1>.
- Guo, J., Li, Y., Cohen, J. B., Li, J., Chen, D., Xu, H., et al. (2019). Shift in the temporal trend of boundary layer height in China using long-term (1979–2016) radiosonde data. *Geophysical Research Letters*, 46, 6080–6089. <https://doi.org/10.1029/2019GL082666>
- Chen, Xinyan, Jianping Guo, Jinfang Yin, Yong Zhang, Yucong Miao, Yuxing Yun, Lin Liu, et al. "Tropopause Trend across China from 1979 to 2016: A Revisit with Updated Radiosonde Measurements." *International Journal of Climatology* 39, no. 2 (Feb 2019): 1117-27. <https://doi.org/10.1002/joc.5866>.
- Zhang, J., Guo, J.*, Xue, H., Zhang, S., Huang, K., Dong, W., et al. (2022). Tropospheric gravity waves as observed by the high-resolution China Radiosonde Network and their potential sources. *Journal of Geophysical Research: Atmospheres*, 127, e2022JD037174. <https://doi.org/10.1029/2022JD037174>
- Lv, Yanmin, Jianping Guo, Jian Li, Lijuan Cao, Tianmeng Chen, Ding Wang, Dandan Chen, et al. "Spatiotemporal Characteristics of Atmospheric Turbulence over China Estimated Using Operational High-Resolution Soundings." *Environmental Research Letters* 16, no. 5 (May 2021). <https://doi.org/10.1088/1748-9326/abf461>.
- Chen, Z.; Tian, Yufang*; Lü, D. Turbulence Parameters in the Troposphere—Lower Stratosphere Observed by Beijing MST Radar. *Remote Sens.* 2022, 14, 947. <https://doi.org/10.3390/rs14040947>
- Chen, Z., Tian, Yufang*, Wang, Y., Bi, Y., Wu, X., Huo, J., Pan, L., Wang, Y., and Lü, D.: Turbulence parameters measured by the Beijing mesosphere–stratosphere–troposphere radar in the troposphere and lower stratosphere with three models: comparison and analyses, *Atmos. Meas. Tech.*, 15, 4785–4800, <https://doi.org/10.5194/amt-15-4785-2022>, 2022.

**Thank you very much
for your attention!**

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