

# The electron temperature distribution up to $10 R_s$ as derived by the DYN model

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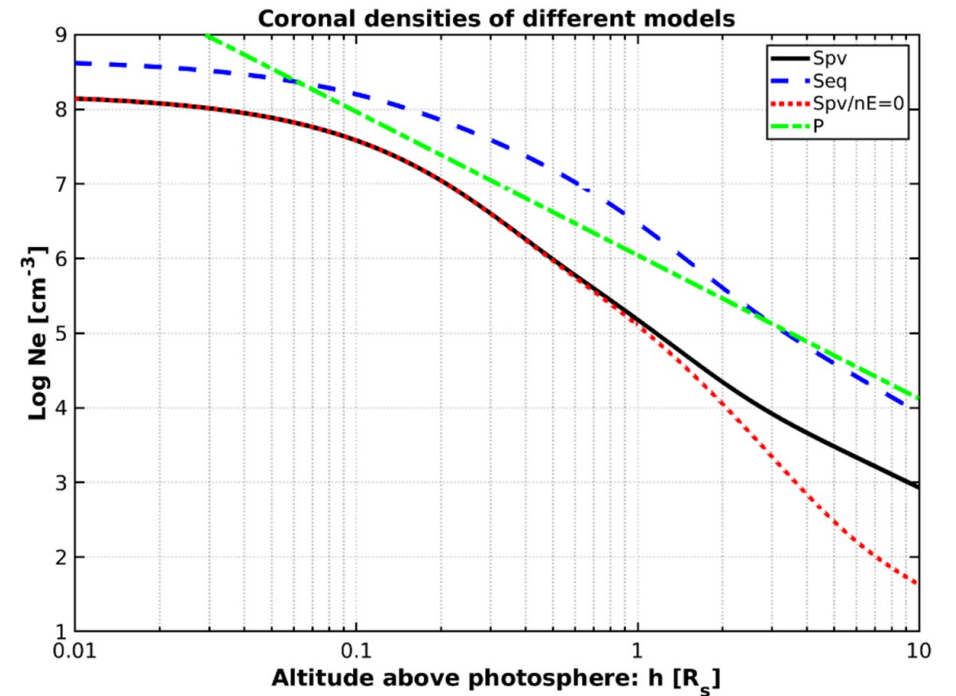
# The DYN model

- Lemaire & Stegen (2016) added the  $m_H n_e(r) \frac{du}{dt}$  factor to Parker's (1958) model:

$$dT(r)/dr + [d(\ln n_e)/dr]^{-1} + m_H * g * R_S^2 / (k r^2) + m_H n_e(r) \frac{du}{dt} = 0$$

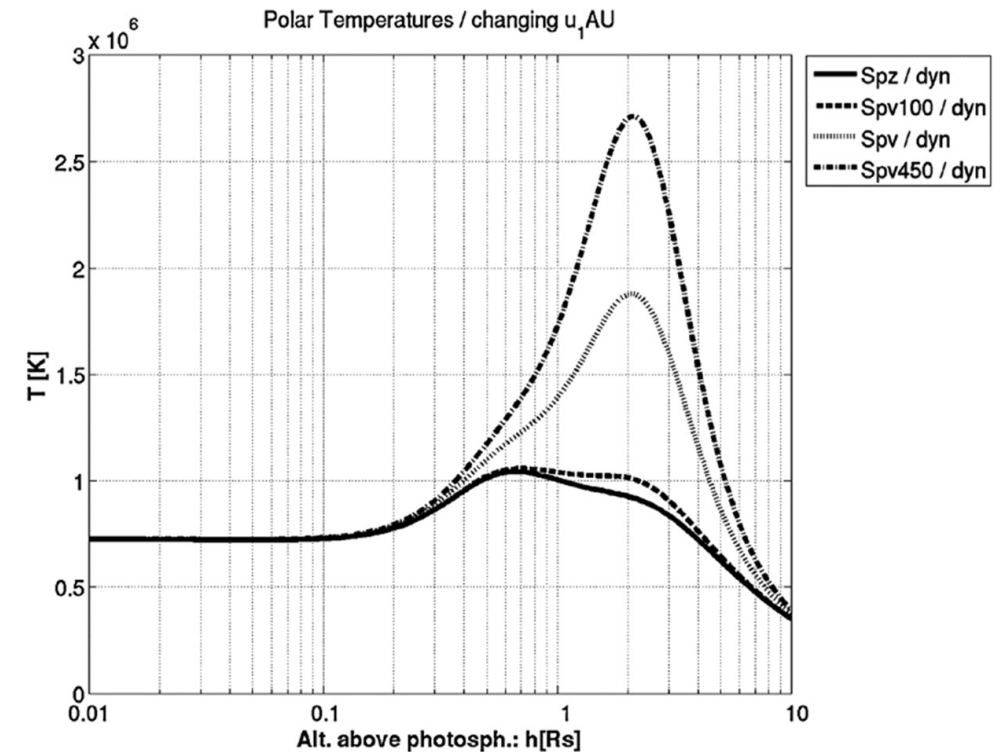
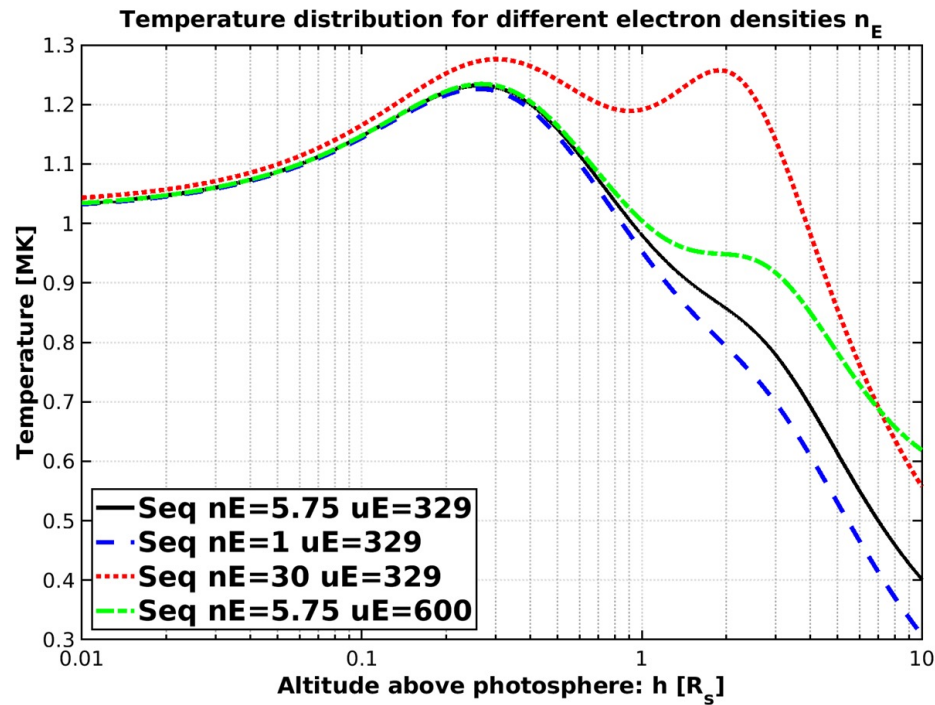
- Pottach (1960), Brandt et al (1965) and Gibson et al (1999) used Parker's hydro-dynamical model with  $T(\infty)=0$  and  $P(\infty)=0$ .
- Lemaire & Stegen (2016) also added a term to Saito (1970)'s fit. This was to correct for  $n_e(1AU)$ . Saito (1970)'s fit up to  $4 R_S$ :

$$n_e(r) = 10^8 [3.09 r^{-16} (1 - 0.5 \sin(\phi)) + 1.58 r^6 (1 - 0.95 \sin(\phi)) + 0.0251 r^{-2.5} (1 - v \sin(\phi))] + n_e(1AU) (215/r)^2$$



# Temperature distributions / Model limitations

$$T(r) = \frac{T^*}{n(r)} \int_r^\infty \frac{n(x)}{x^2} dx, \text{ where } T^* = m_H * g * R_S^2 / (k r^2)$$



# PUNCH's role

- Wind speed measurements at such low radial distances

$$[u(r) = u_E \frac{A_E}{A(r)} \frac{n_E}{n(r)}]$$

- DYN is a global model that depends on the large-scale flow
- DYN is bimodal (equatorial= slow, polar= fast)