

Mechanisms of North American Drought as diagnosed with models and the 20th Century Reanalysis

Richard Seager

Lamont Doherty Earth Observatory

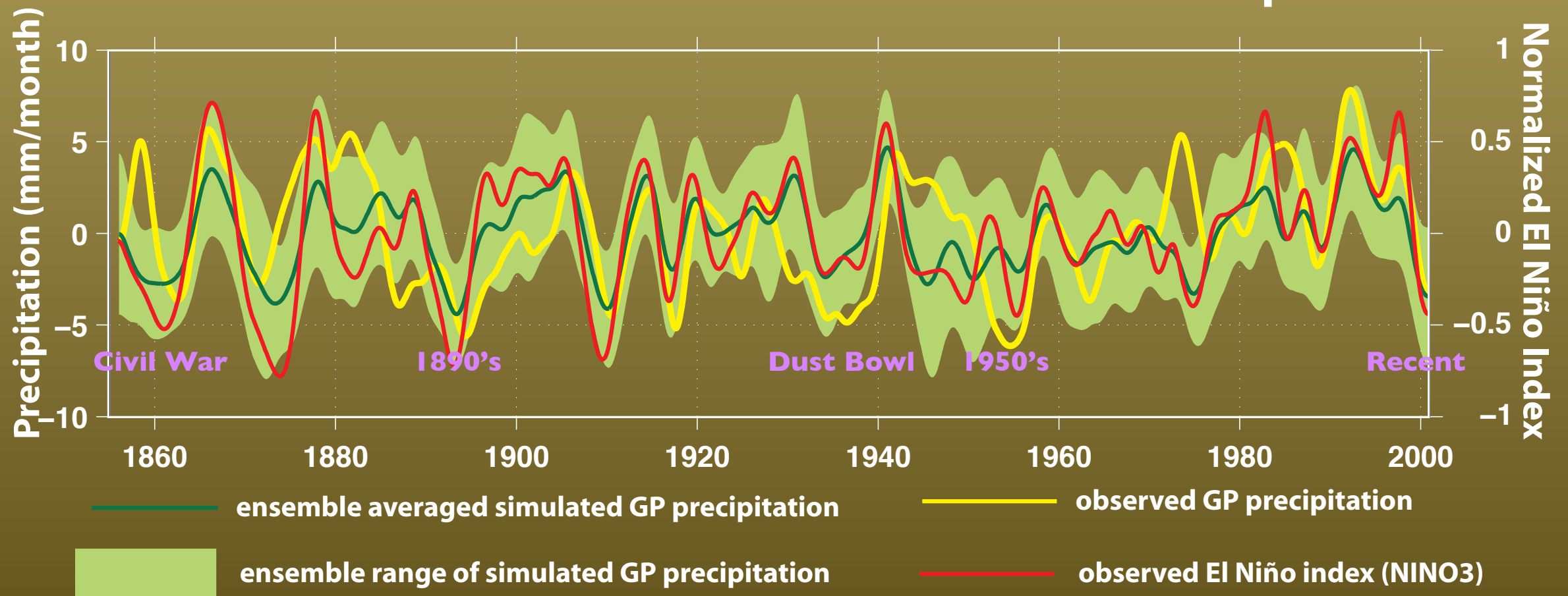
Ben Cook and Ron Miller

NASA Goddard Institute for Space Studies

Models have established oceanic causes of N. American drought: e.g. Great Plains precipitation variability since 1856

Seager et al. (2005)

Climate Model Simulation¹ of U.S. Great Plains² Precipitation

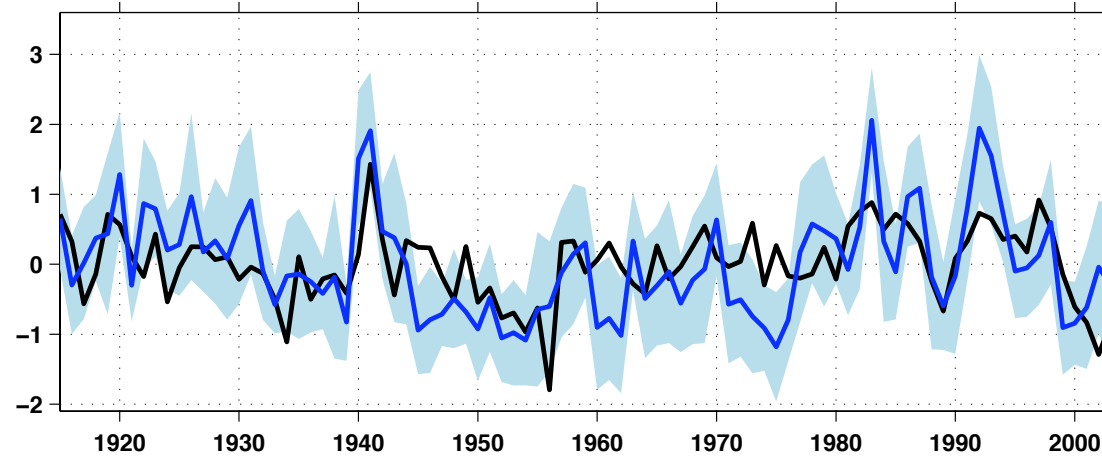


1. An ensemble of 16 model runs with observed SST prescribed in the equatorial Pacific Ocean (20°S-20°N) and calculated elsewhere, using a two-layer slab ocean model
2. Great Plains are defined as the area between 110° W– 90° W and 30°N–50°N

VIC and CCM3 simulations of soil moisture

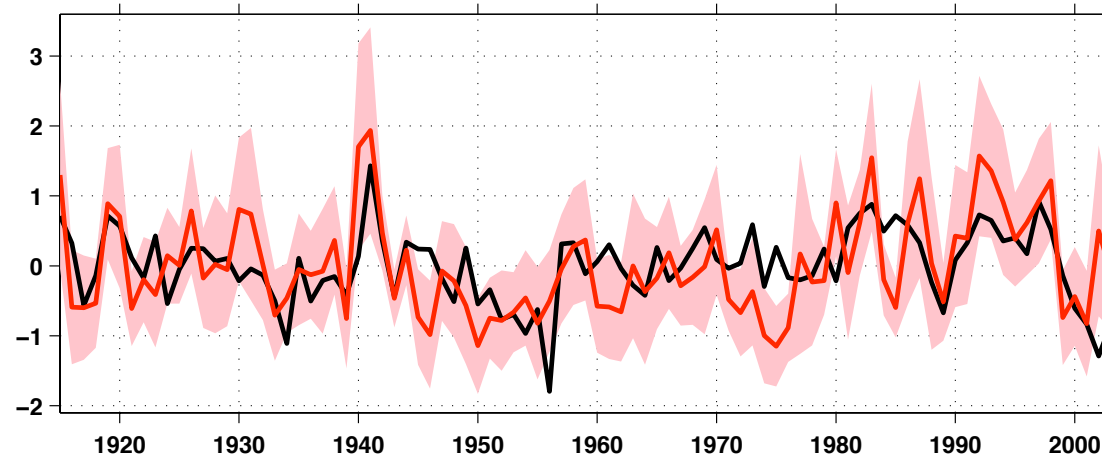
Tropical Pacific SSTs are dominant, as seen in simulation of last century of soil moisture in SWNA.

Std Annual VIC (black), GOGA Mean (blue), +/- 2 STD (lt blue) SM 25N-40N, 95W-125W US



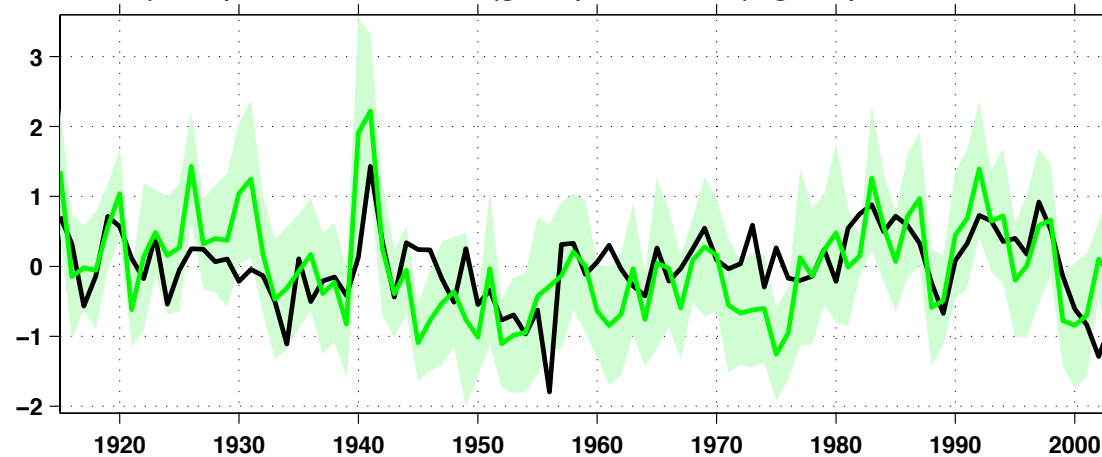
global SST forcing

Std Annual VIC (black), POGA-ML Mean (red), +/- 2 STD (pink) SM 25N-40N, 95W-125W US



tropical Pacific SST forcing

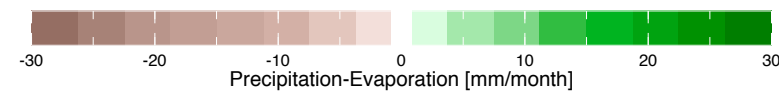
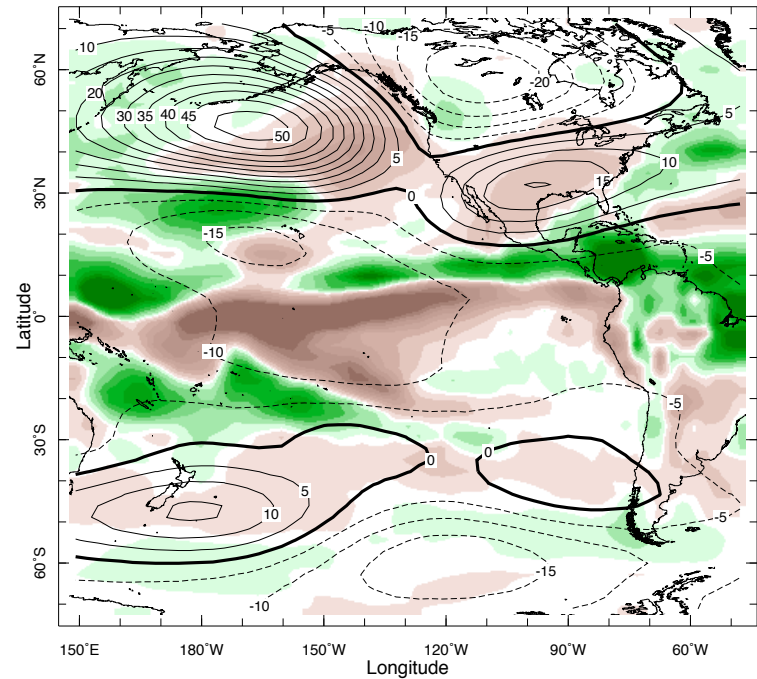
Std Annual VIC (black), GOGA VTG Mean (green), +/- 2 STD (lt green) SM 25N-40N, 95W-125W US



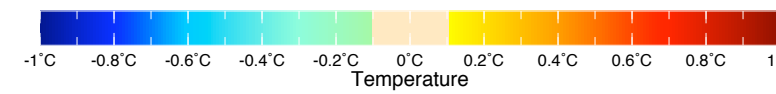
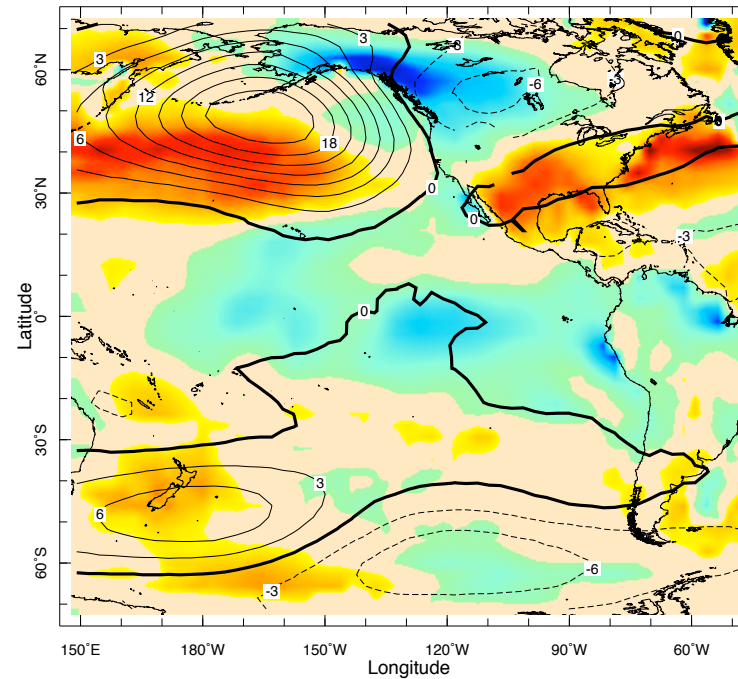
global SST forcing plus variable trace gases

GOGA Oct-Mar 1948-1957 Average

P-E (color), 250 mb Height (contours)

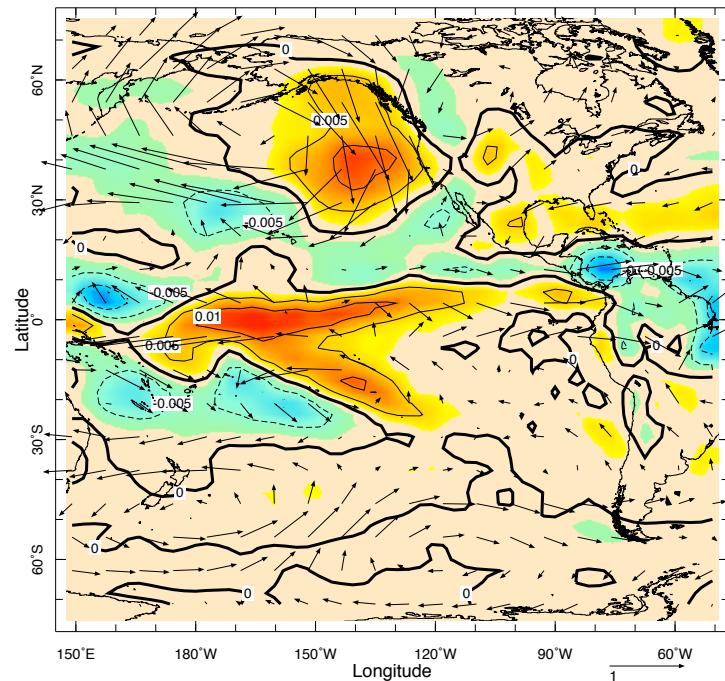


Sfc T (color), 850 mb Height (contour)



winter half years of
1950s drought

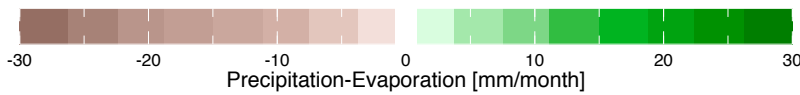
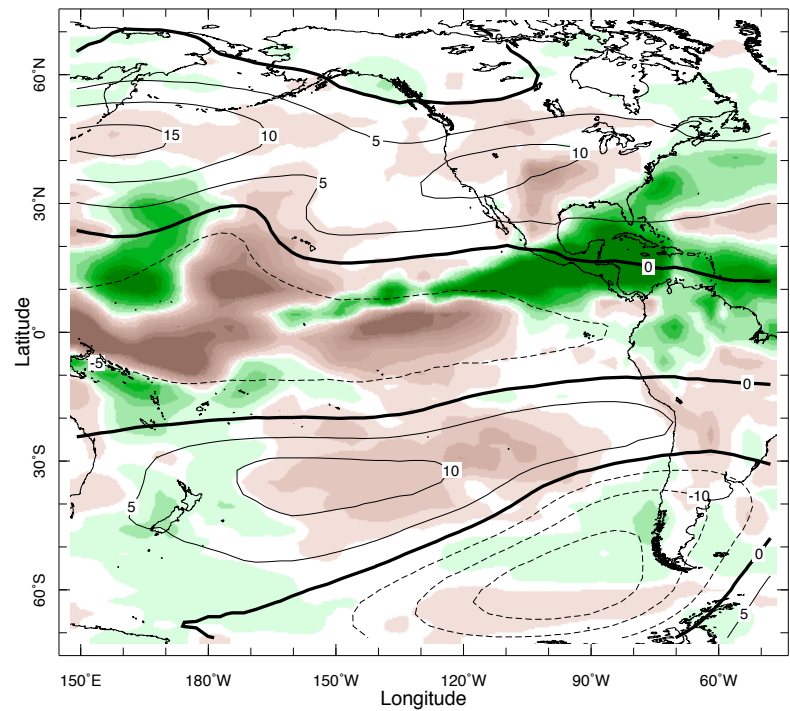
500 mb Vert Vel, 850 mb Winds



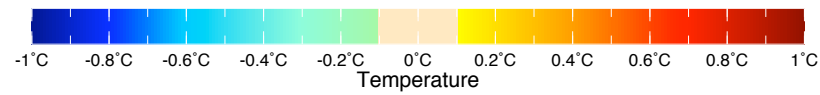
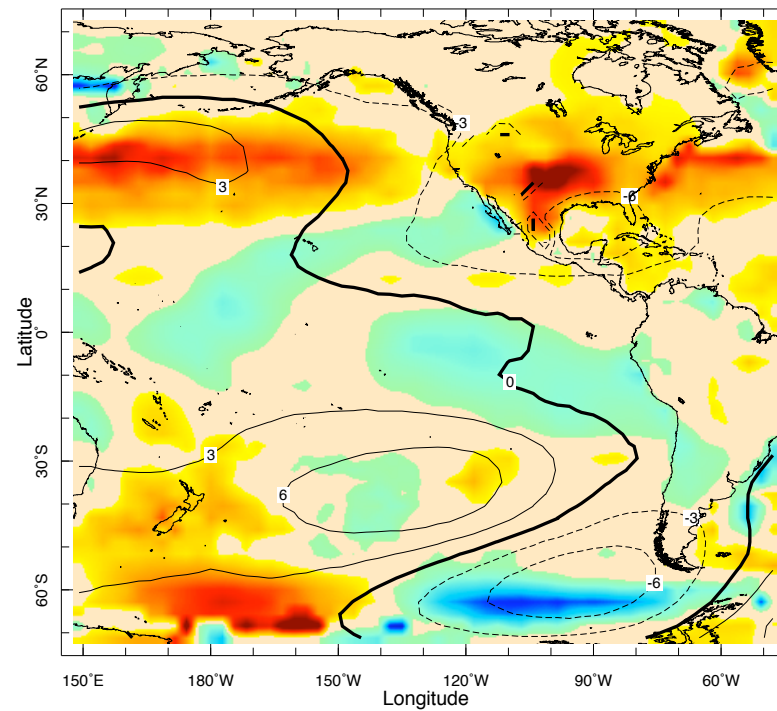
Mechanisms for drought
have been proposed based
on SST-forced models.
Forced-mid-latitude ridge
and descent, northward
shifted storm track.

GOGA Apr-Sep 1948-1957 Average

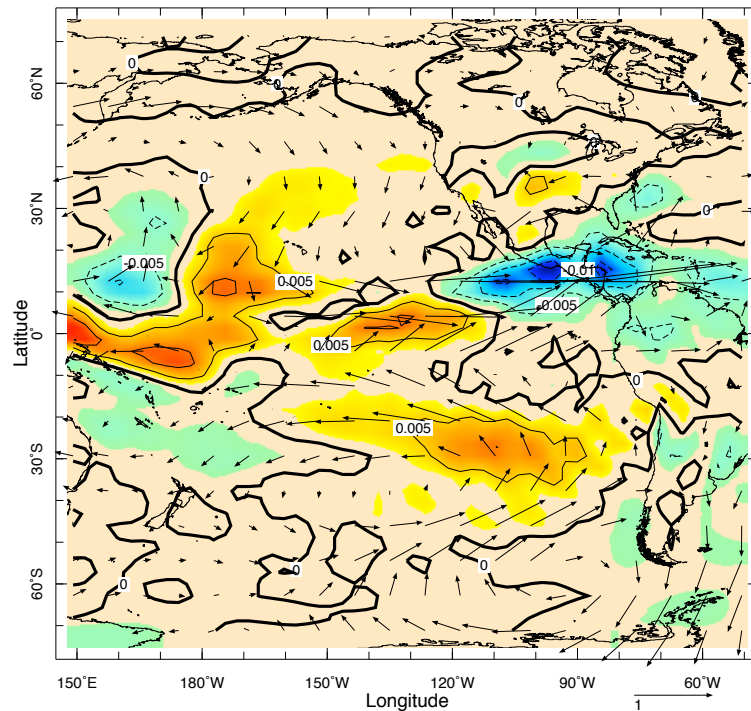
P-E (color), 250 mb Height (contours)



Sfc T (color), 850 mb Height (contour)



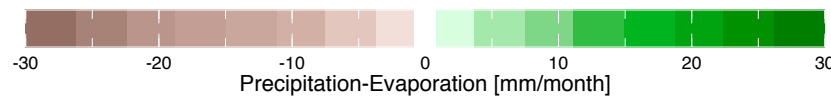
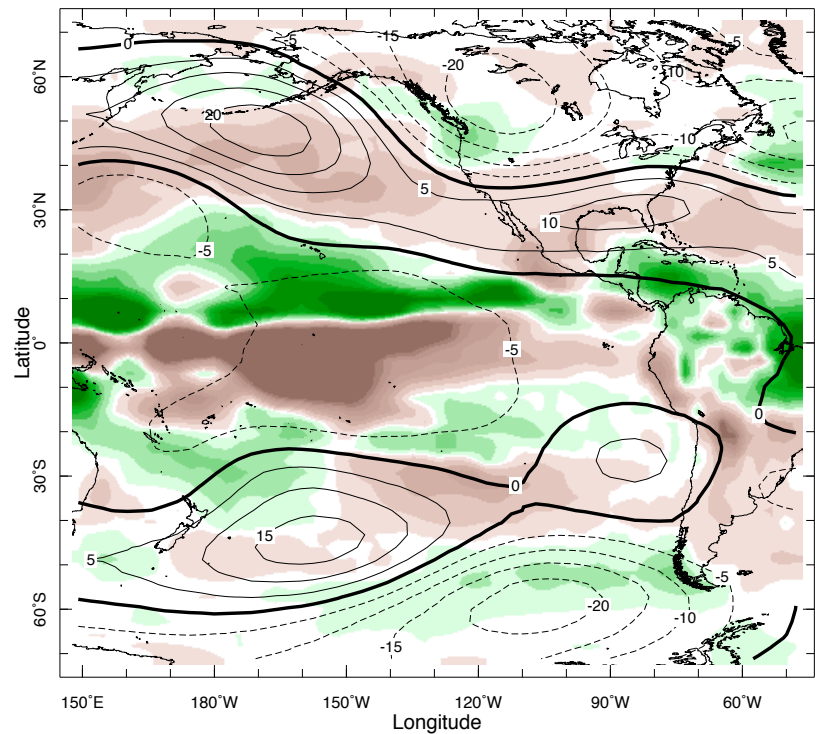
500 mb Vert Vel, 850 mb Winds



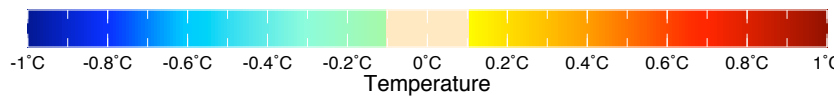
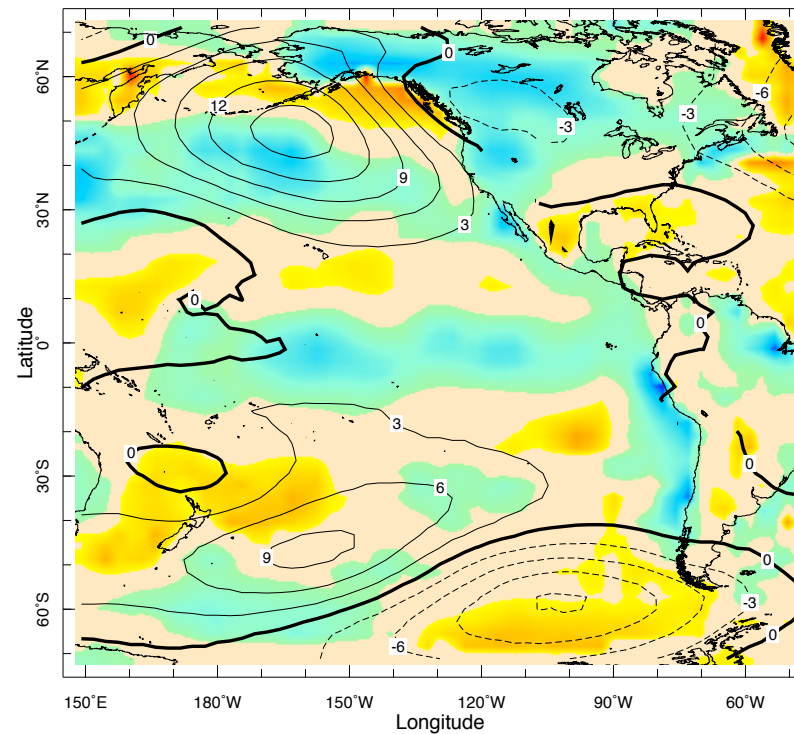
Summer half years of the 1950s drought. SST-forced mid-latitude ridge and also evidence of N. Atlantic forcing of low level low.

GOGA Oct-Mar 1932-1939 Average

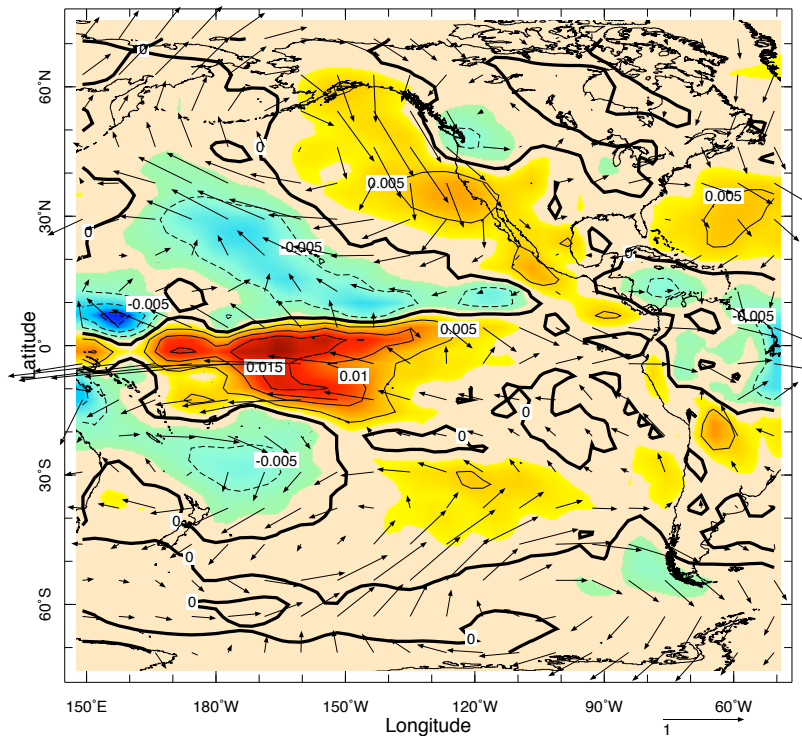
P-E (color), 250 mb Height (contours)



Sfc T (color), 850 mb Height (contour)



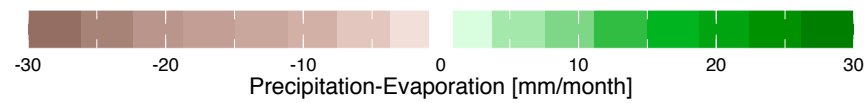
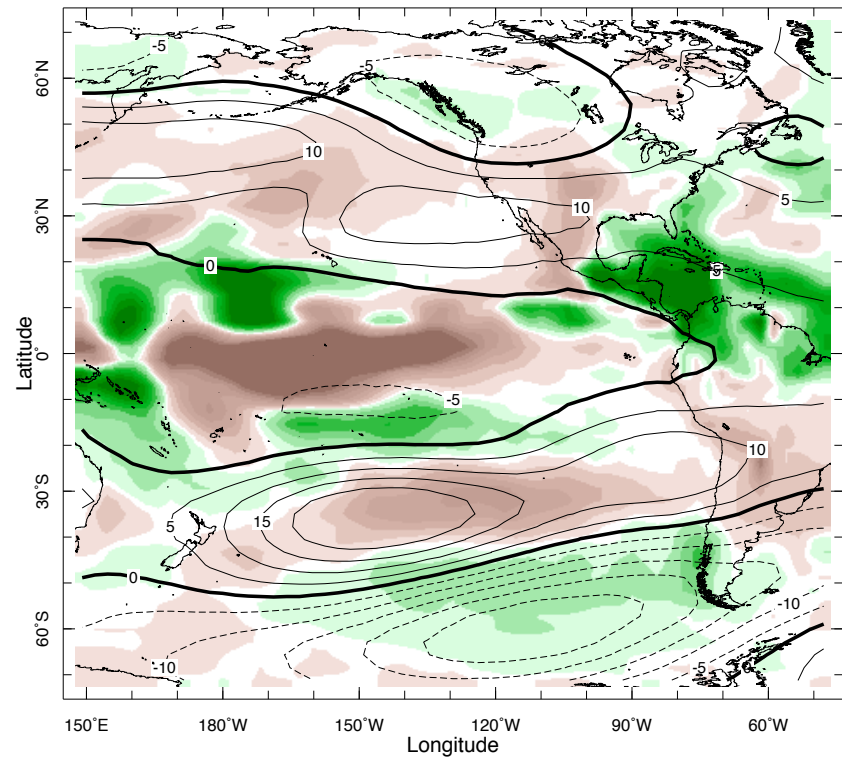
500 mb Vert Vel, 850 mb Winds



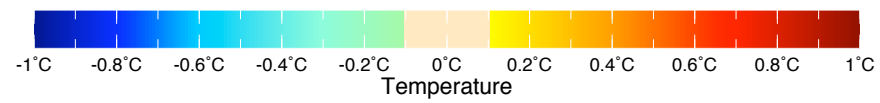
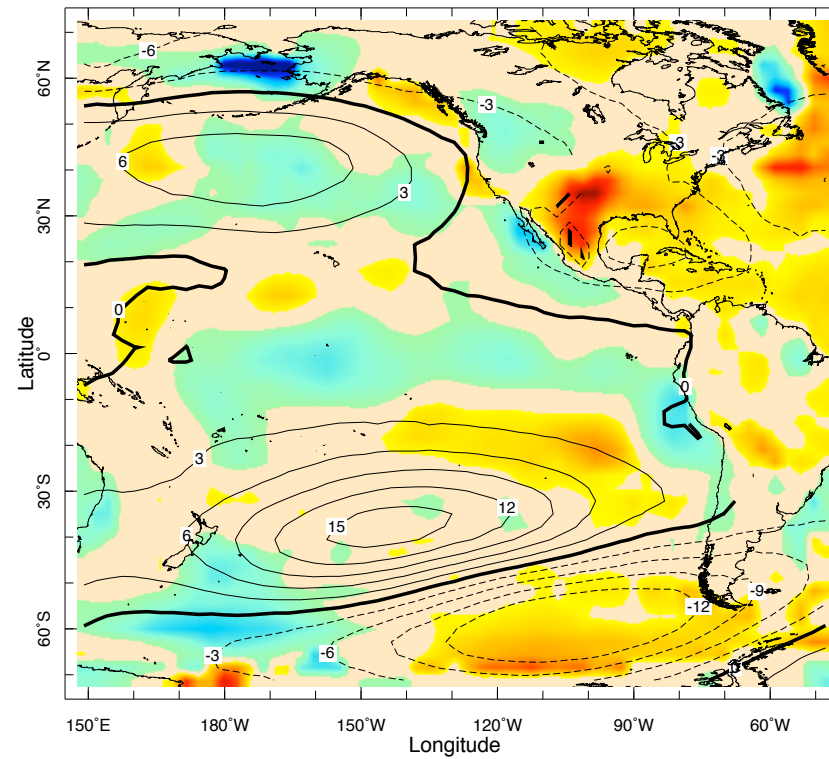
The modeled circulation and P-E anomalies for the 1930s drought look much like those for the 1950s drought 'cos of La Nina-warm N. Atlantic for both.

GOGA Apr-Sep 1932-1939 Average

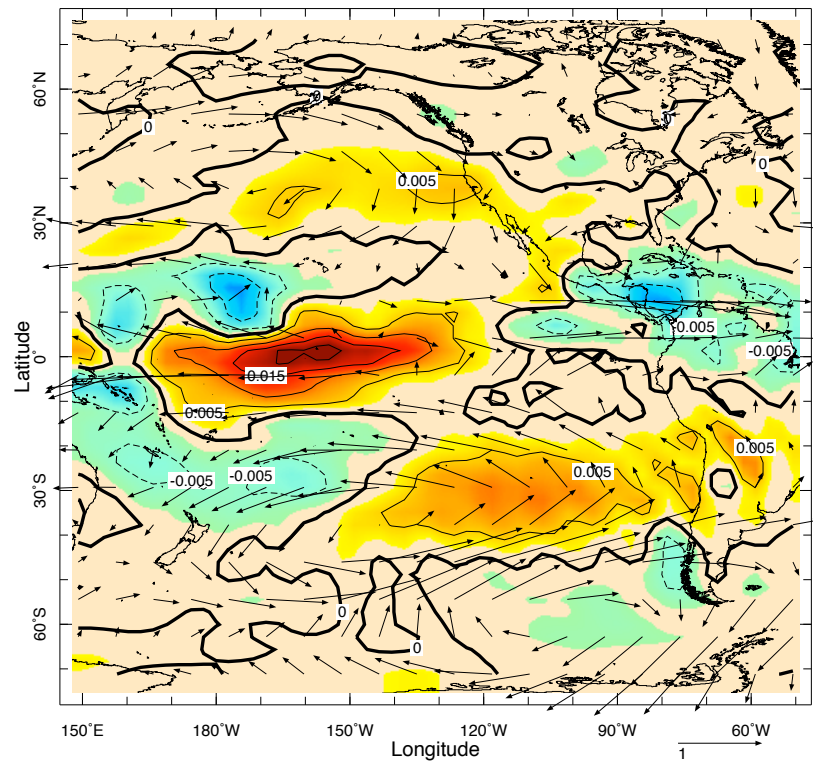
P-E (color), 250 mb Height (contours)



Sfc T (color), 850 mb Height (contour)



500 mb Vert Vel, 850 mb Winds

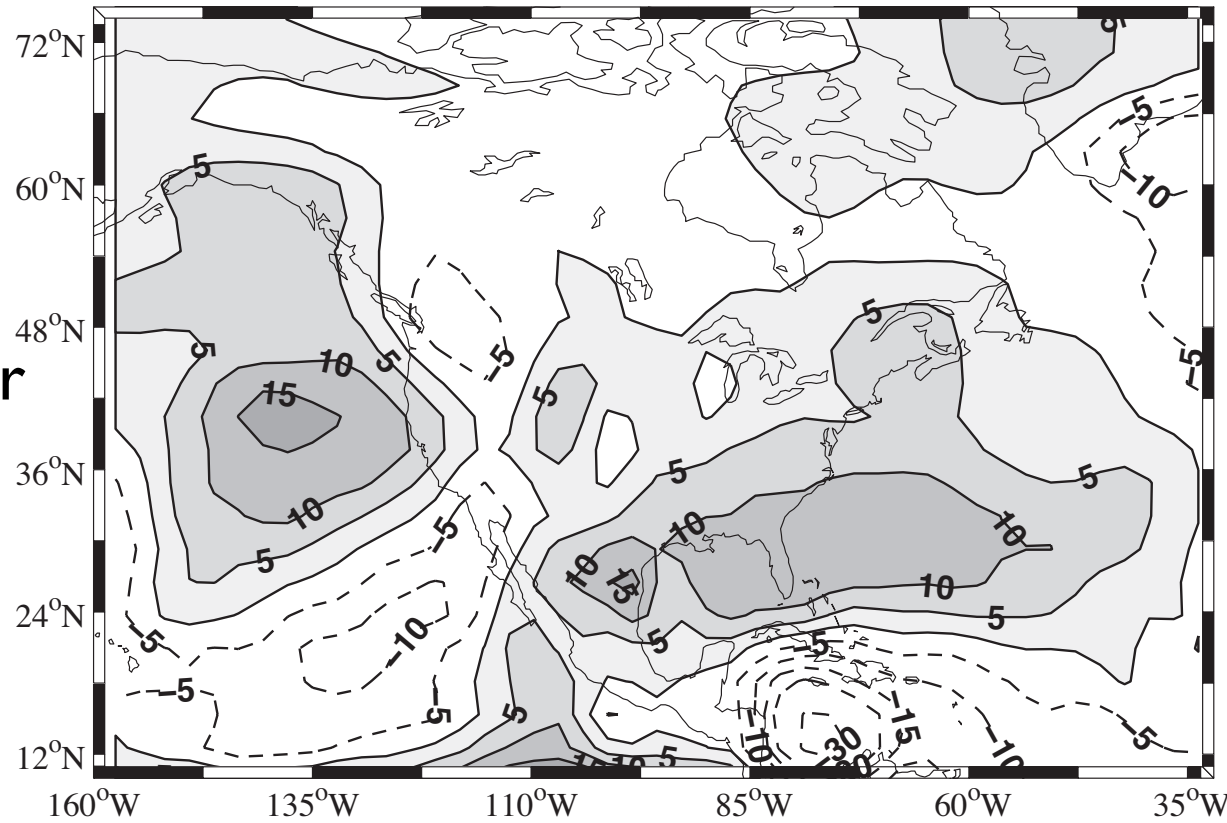


1930s and 1950s
similarity is in summer
half years as well

Despite the modeling progress the lack of long atmospheric data sets made validation of the mechanisms of drought difficult. In particular how similar really were the Dust Bowl and 1950s droughts given the more northward center of the Dust Bowl?

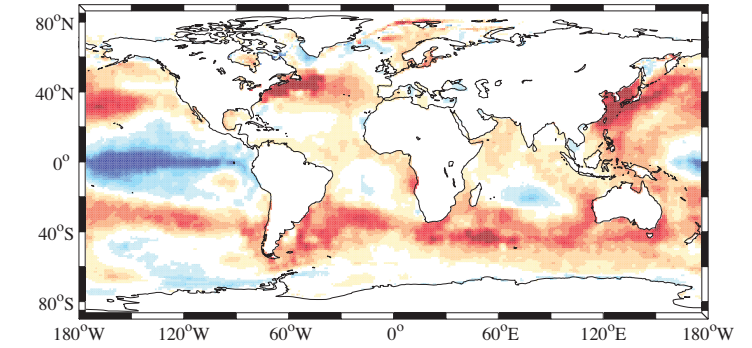
Enter: The 20th Century Reanalysis!
And the US CLIVAR Drought WG experiments.

ONDJFM PAC-Cold/ATL-Warm Omega Anomaly: 500 hPa

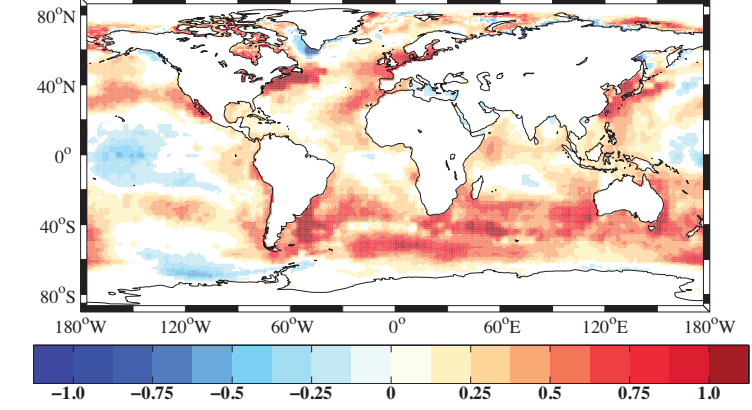


winter

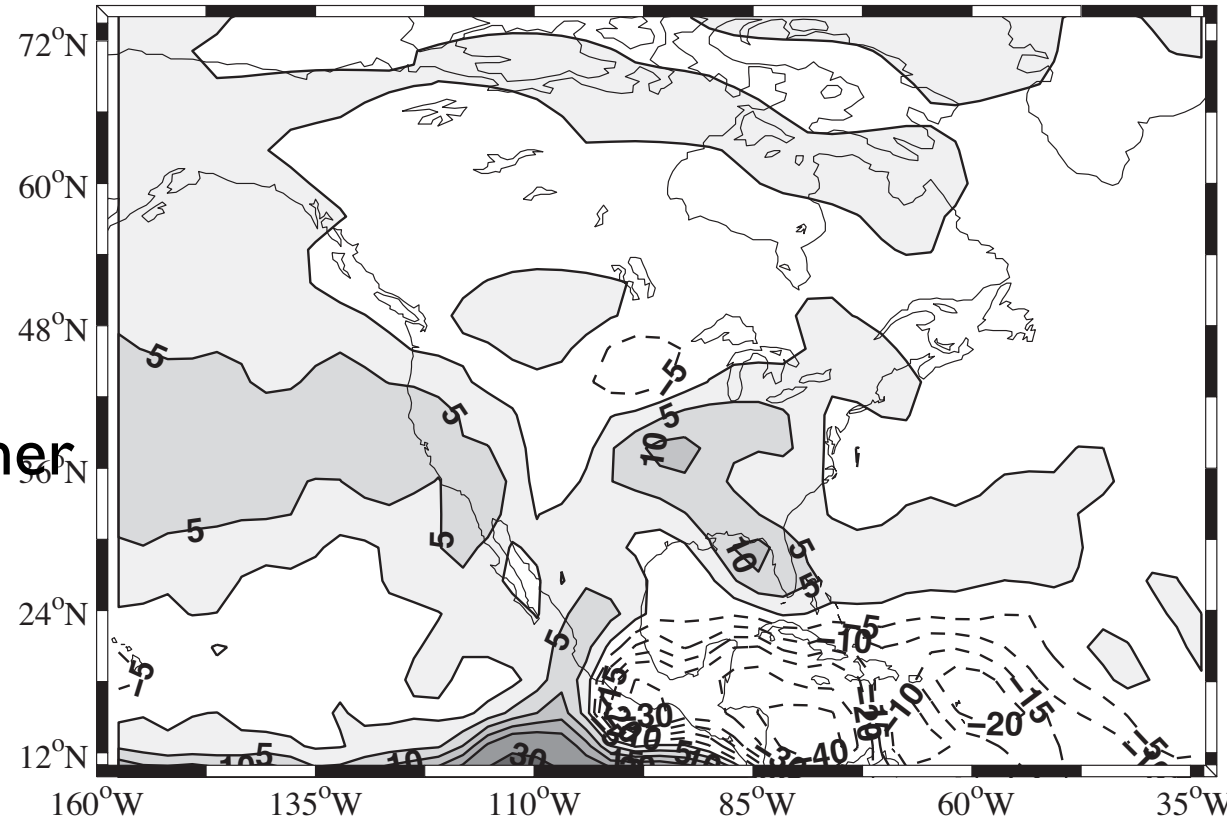
PCAW ONDJFM SST Anomaly



PCAW AMJJAS SST Anomaly



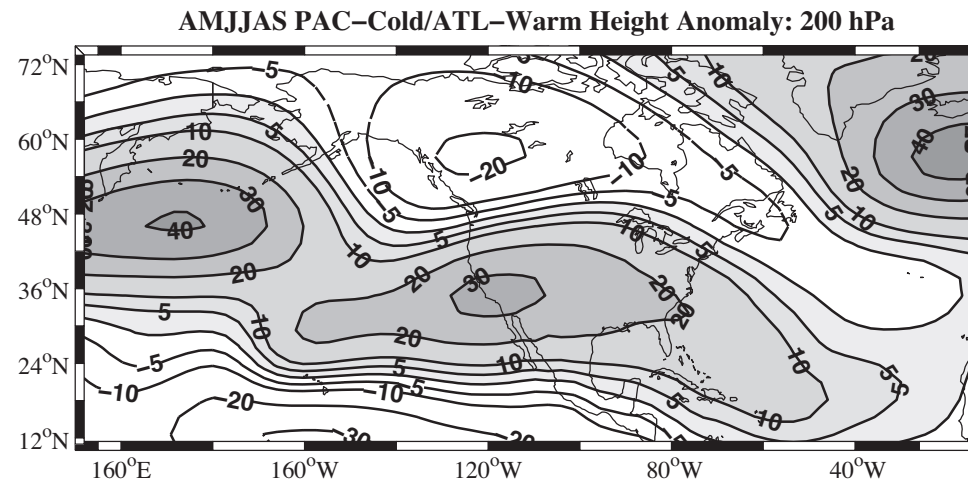
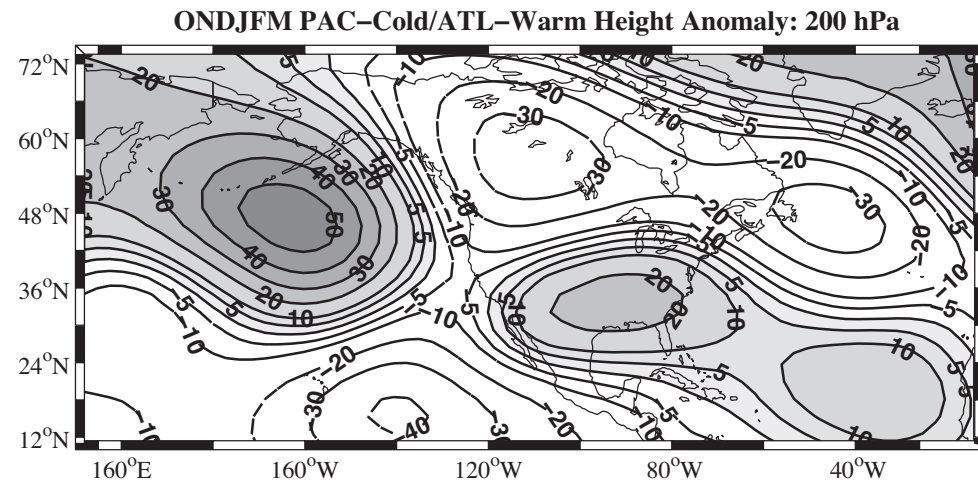
AMJJAS PAC-Cold/ATL-Warm Omega Anomaly: 500 hPa



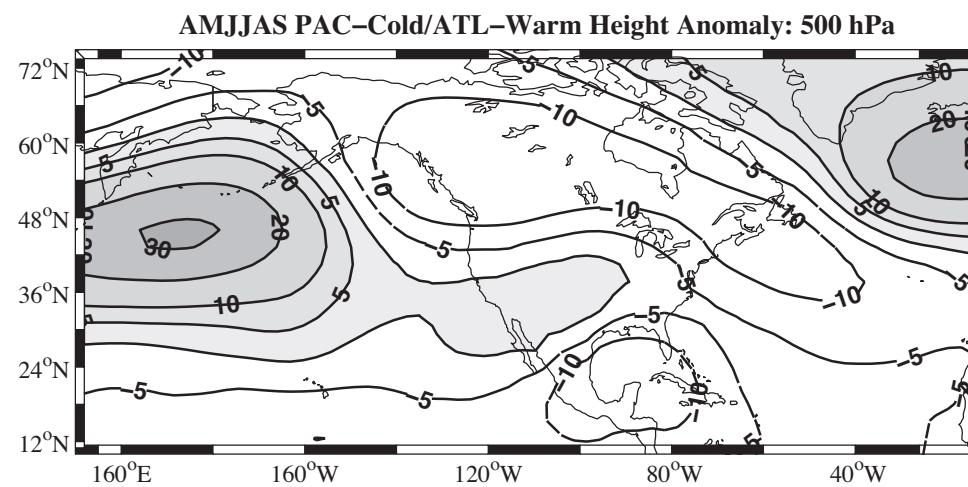
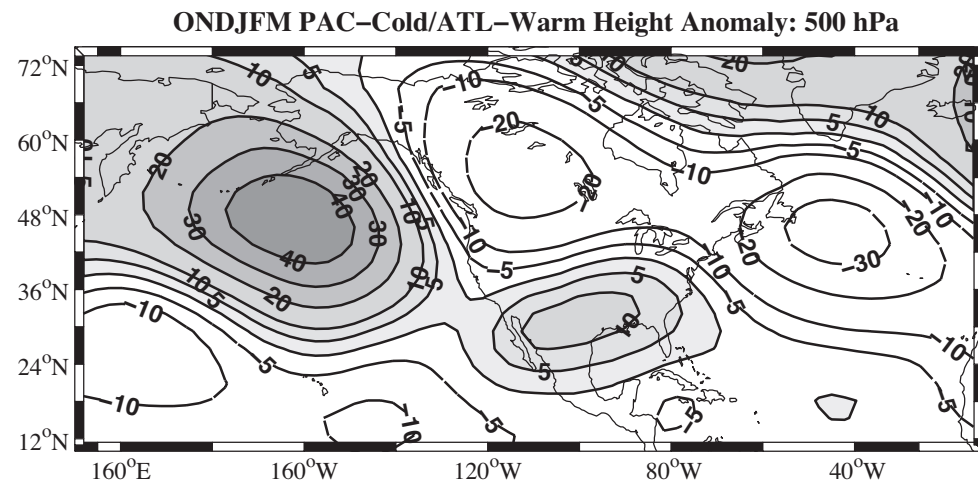
summer

The multimodel mean 500mb
omega for cold Pacific-warm
N.Atlantic SSTA.
Subsidence off W. Coast and
over Gulf of Mexico,
Florida ...

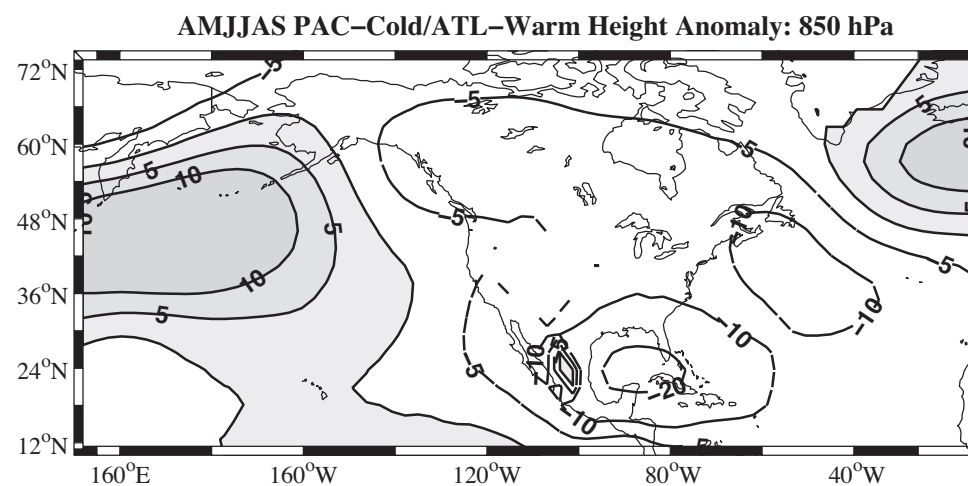
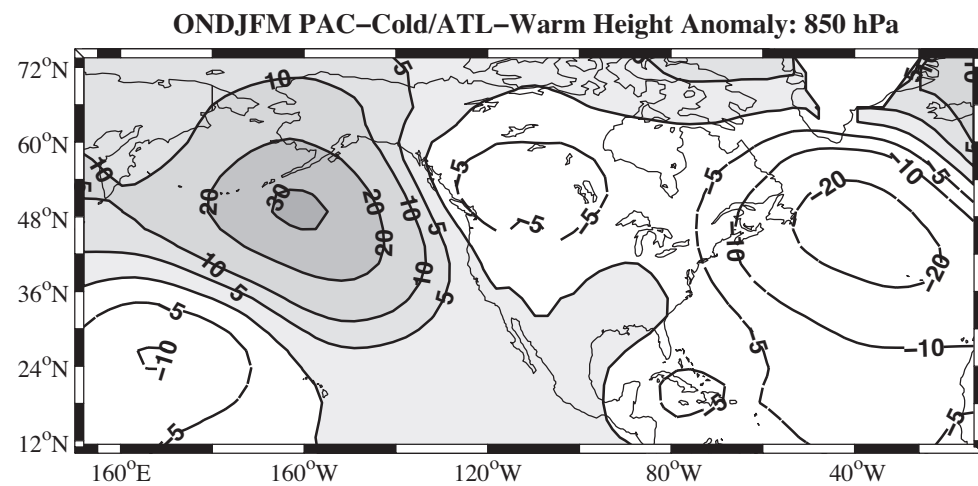
Cold Pacific-warm Atlantic height anomalies



200mb



500mb



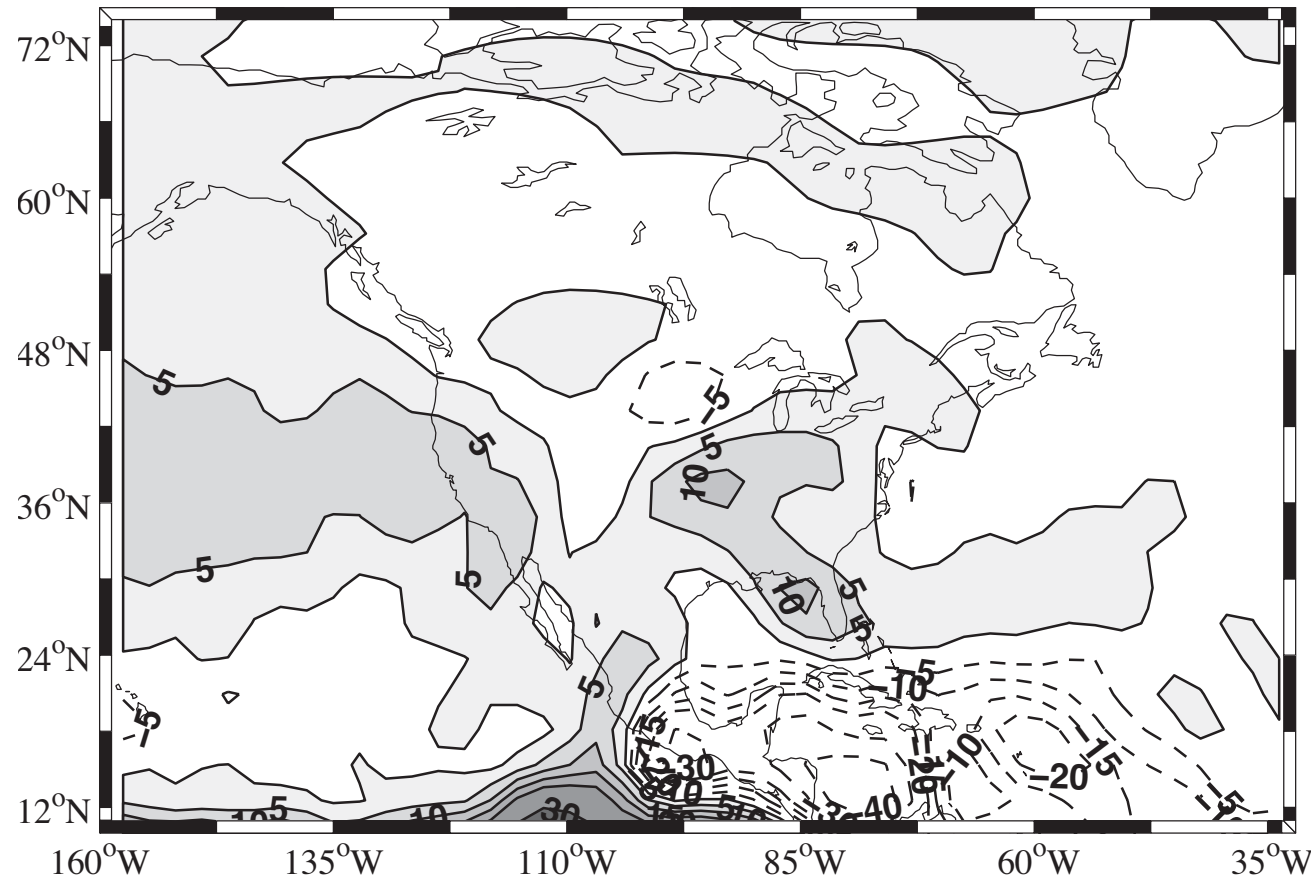
850mb

Winters

Summers

*note low level summer
Gulf of Mexico low*

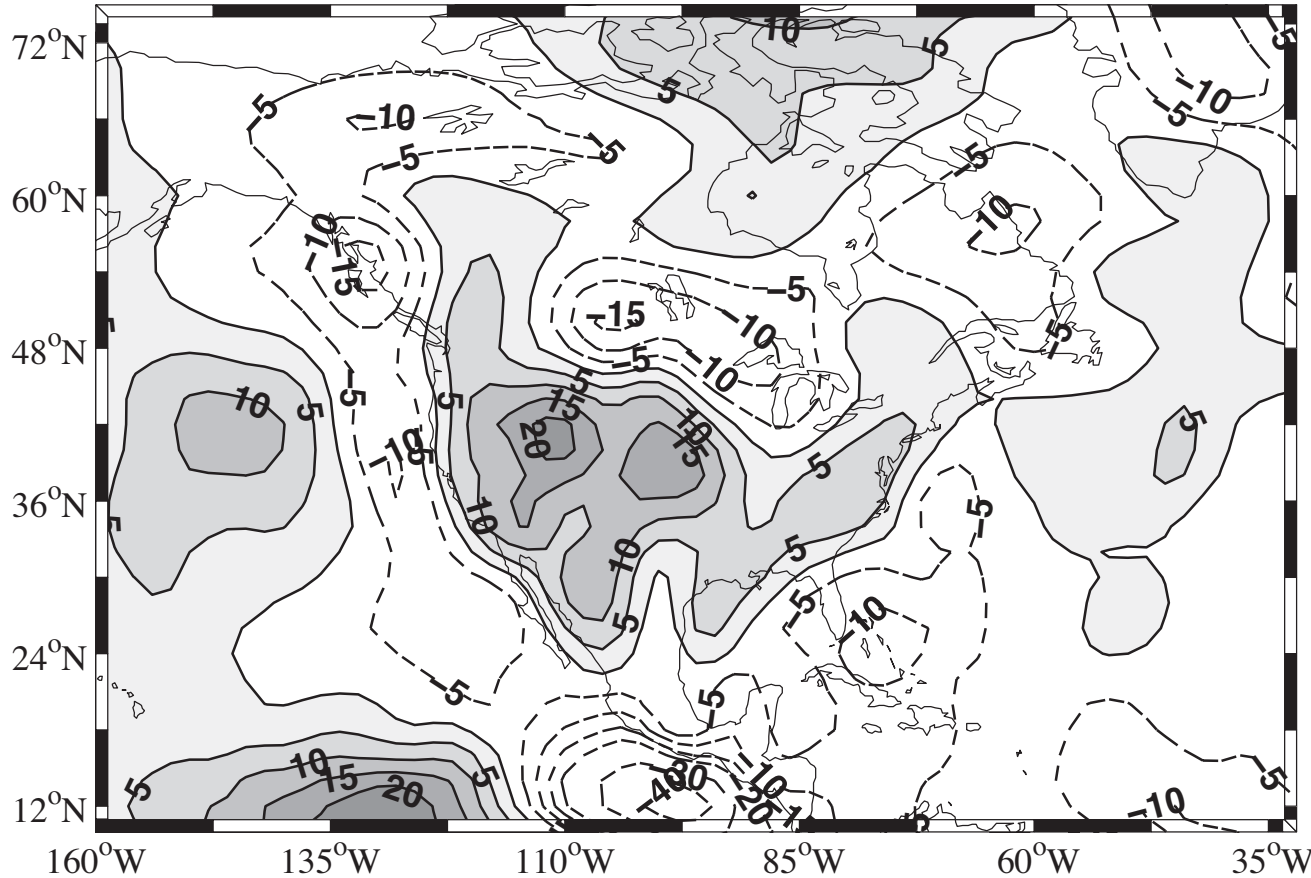
AMJJAS PAC-Cold/ATL-Warm Omega Anomaly: 500 hPa



SST-forced
'expected' 500mb
summer omega

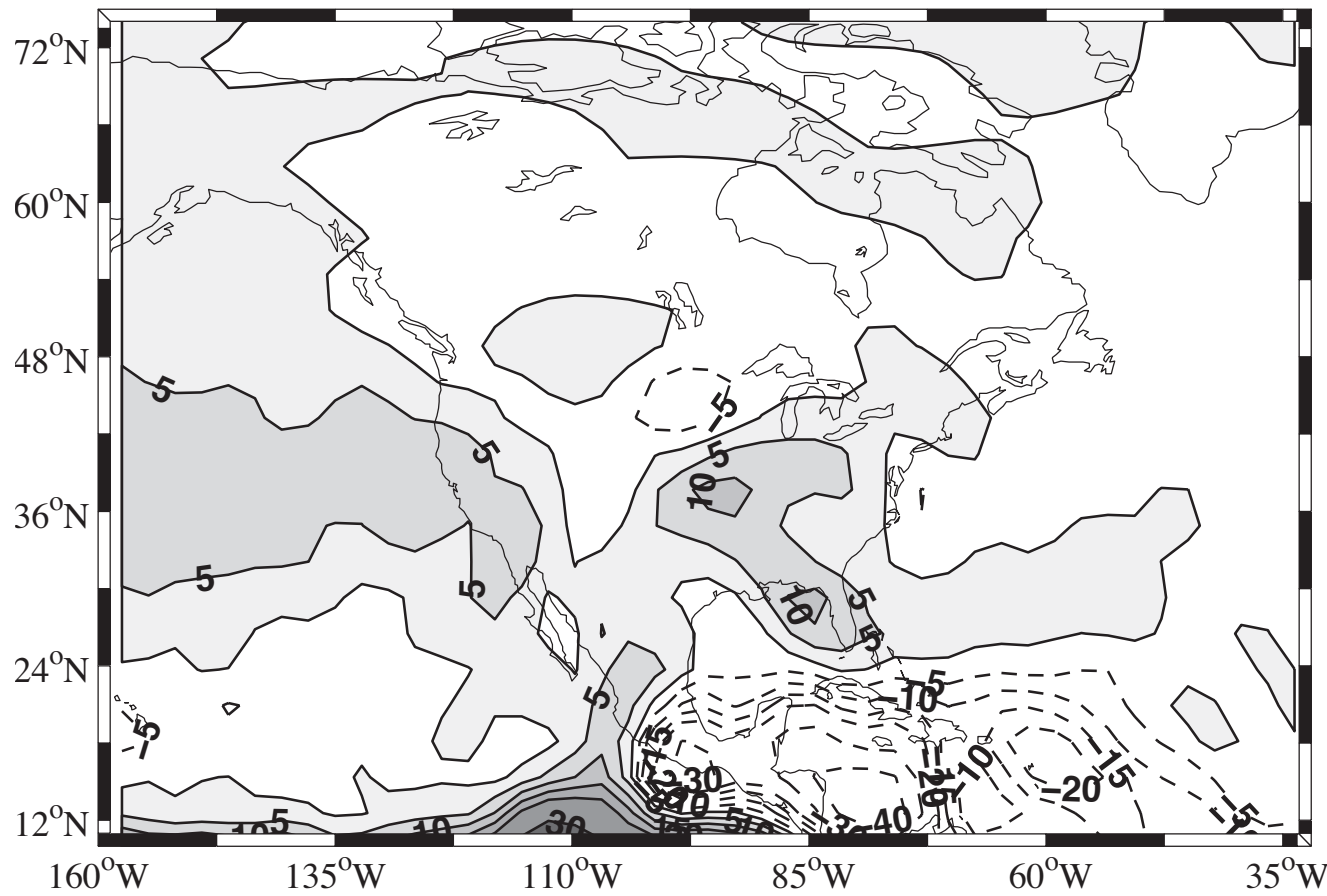
During the Dust Bowl,
strong continental-
centered subsidence at
variance with that
expected from 1930s SST
forcing

AMJJAS Omega Anomaly: 500 hPa, 1932-1939



'actual' 20CR
500mb omega in
1930s summers.

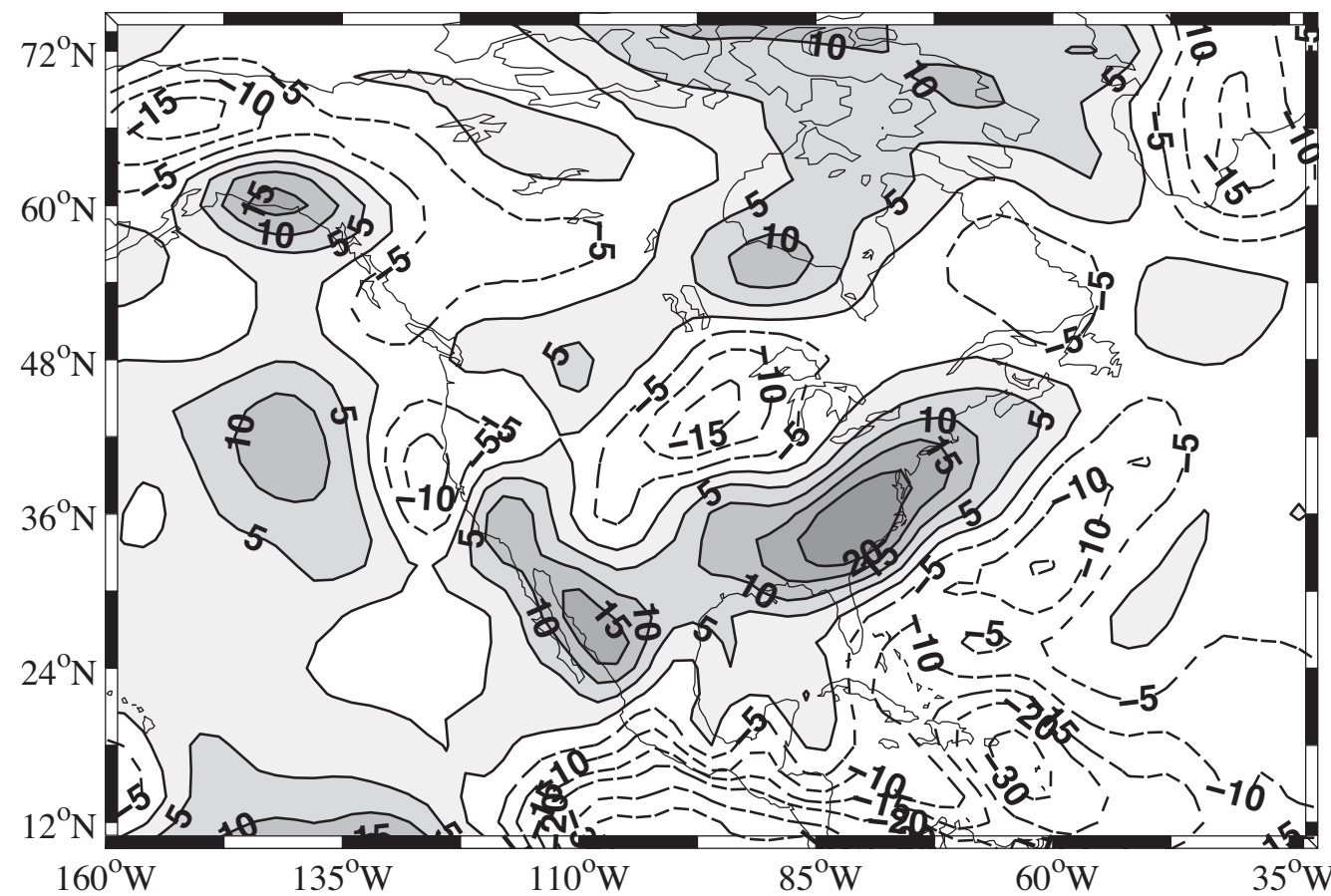
AMJJAS PAC-Cold/ATL-Warm Omega Anomaly: 500 hPa



SST-forced
'expected' 500mb
summer omega

Better agreement between
expected and actual
500mb omega during the
1950s drought.

AMJJAS Omega Anomaly: 500 hPa, 1948-1957



'actual' 20CR
500mb omega in
1950s summers.

What explains Dust Bowl departure from normal pattern?

Hoerling et al. (2009) appeal to internal atmospheric variability.

But the Dust Bowl drought was unique in that wind erosion, caused by poor land use practices, created vast dust storms

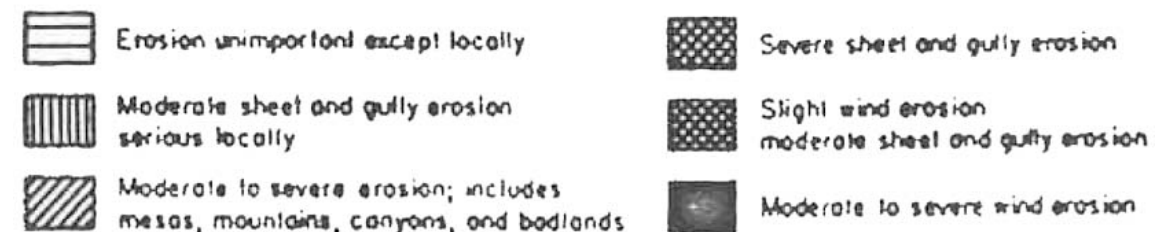
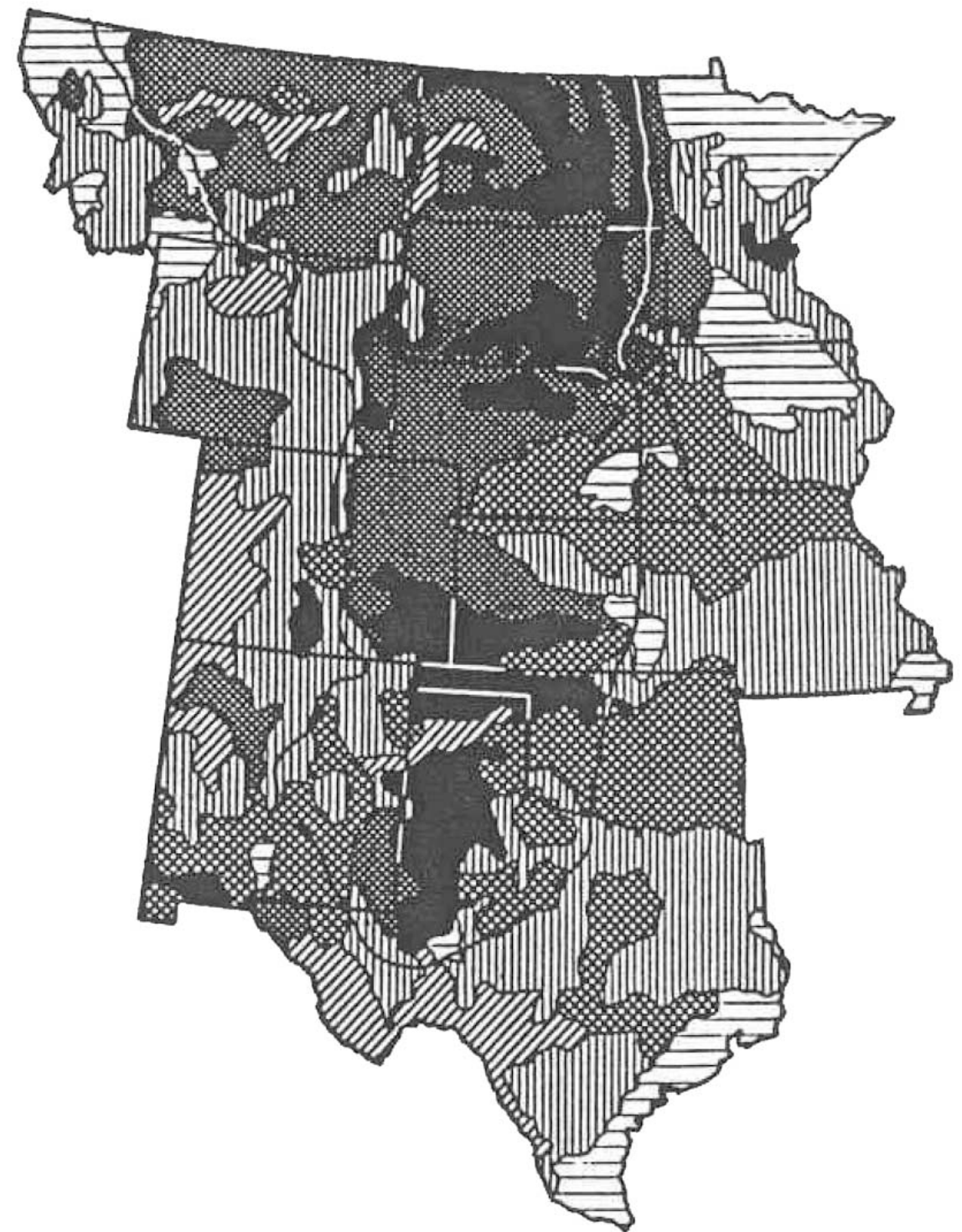
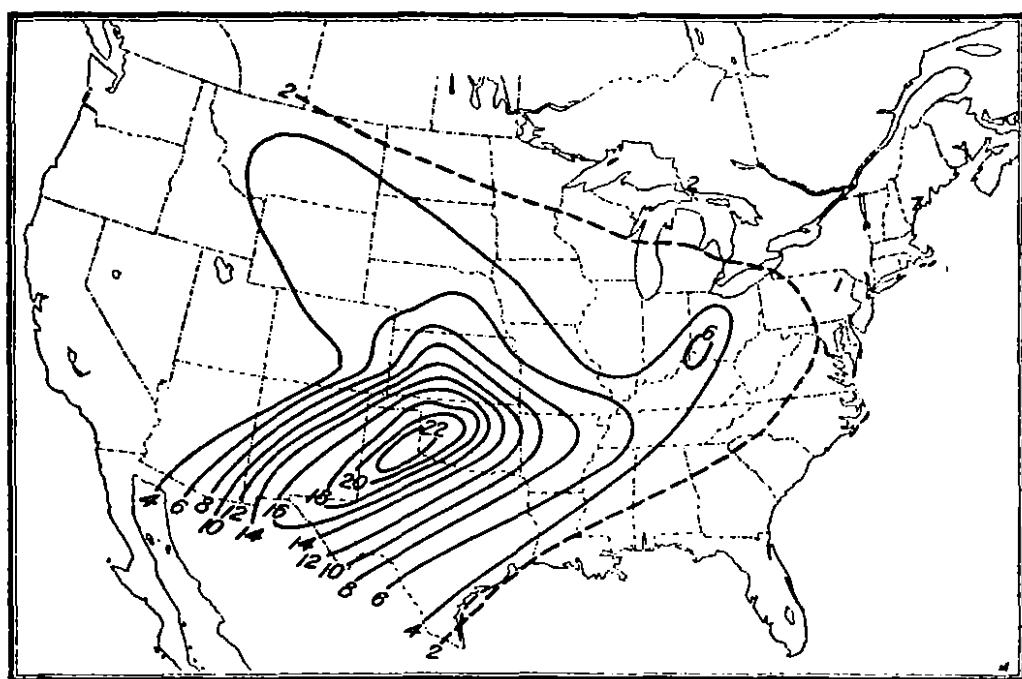


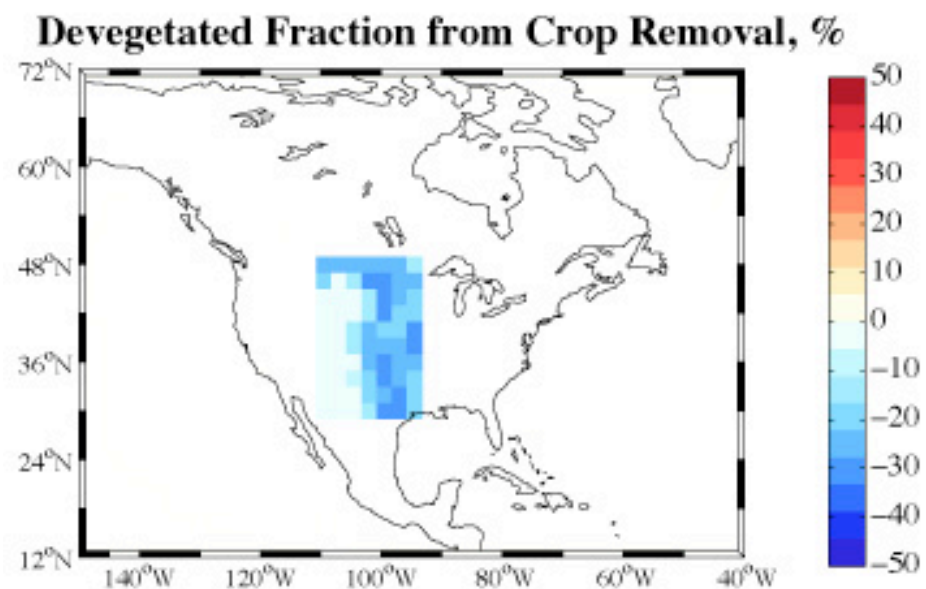
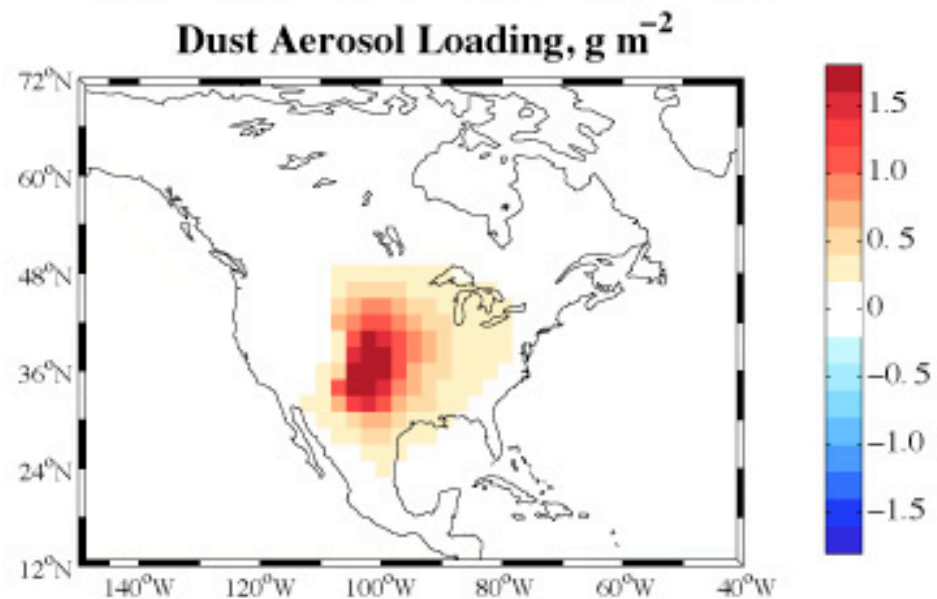
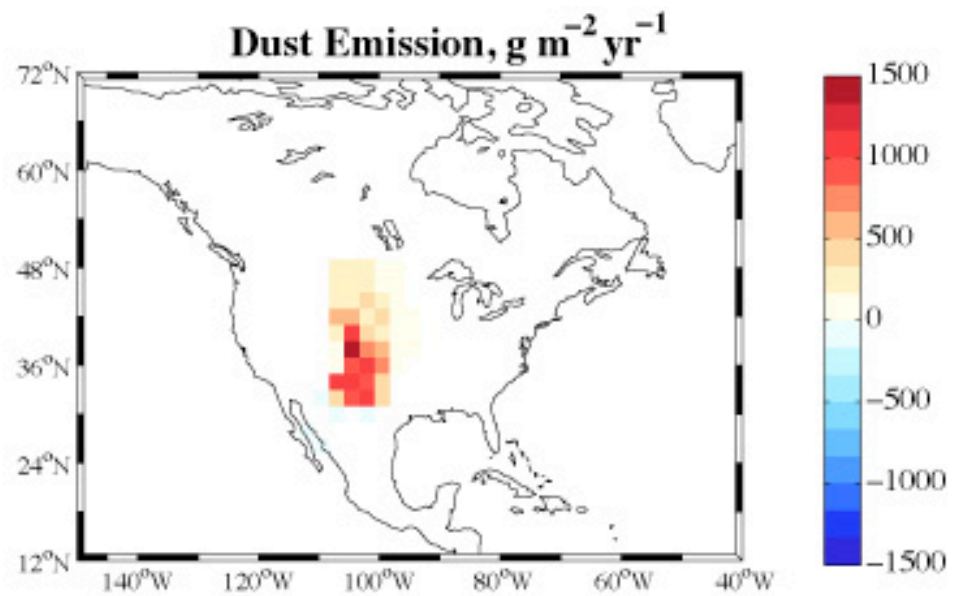
FIG. 1.—Wind erosion in the Great Plains in the 1930s. An irregular line bounds the Great Plains region as delimited by the Great Plains Committee. Source: Adapted from “General Distribution of Erosion” (U.S. Dept. Agriculture, Soil Conservation Service, August 1936).

contemporary observations of dust storms and modeled dust storms
 (GISS model, Cook et al. (2008, 2009))



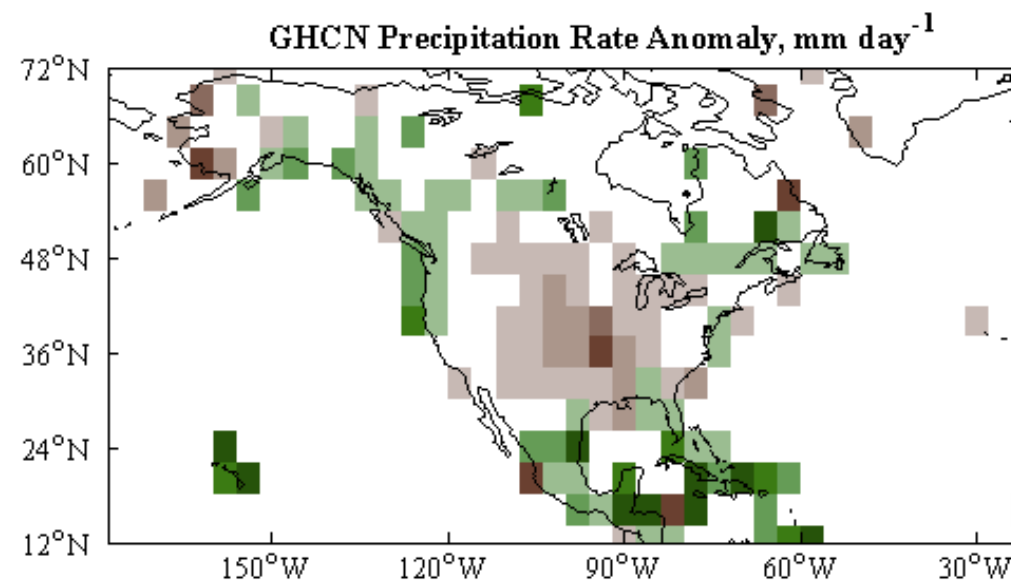
Number of days with duststorms, or dusty conditions, March 1936.—W. A. M

Martin, 1936

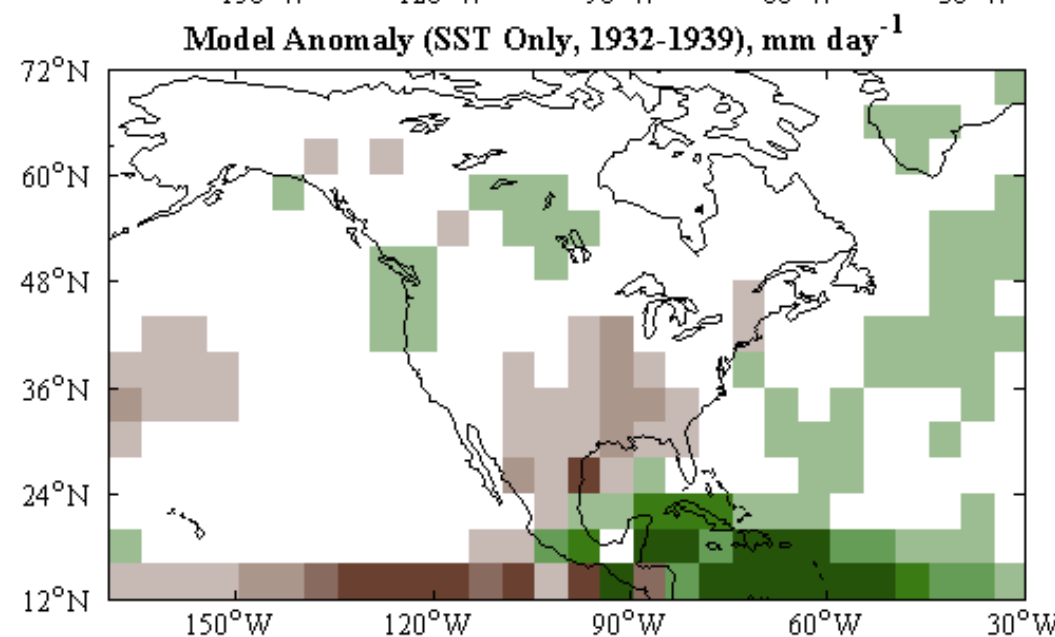


Based on wind erosion maps convert portions of model grid boxes to bare soil

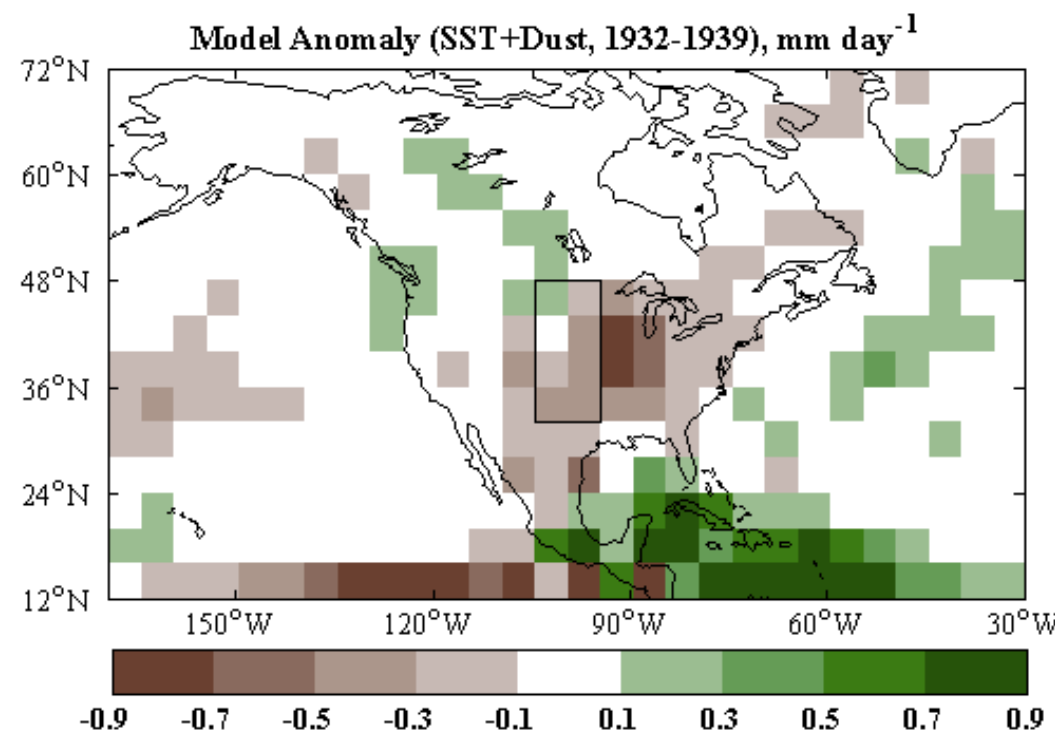
Model created dust storms, the dust interacted with solar and longwave radiation intensifying the drought and moving it north



Observed
1930s
precipitation
anomaly



Modeled with
SST forcing only

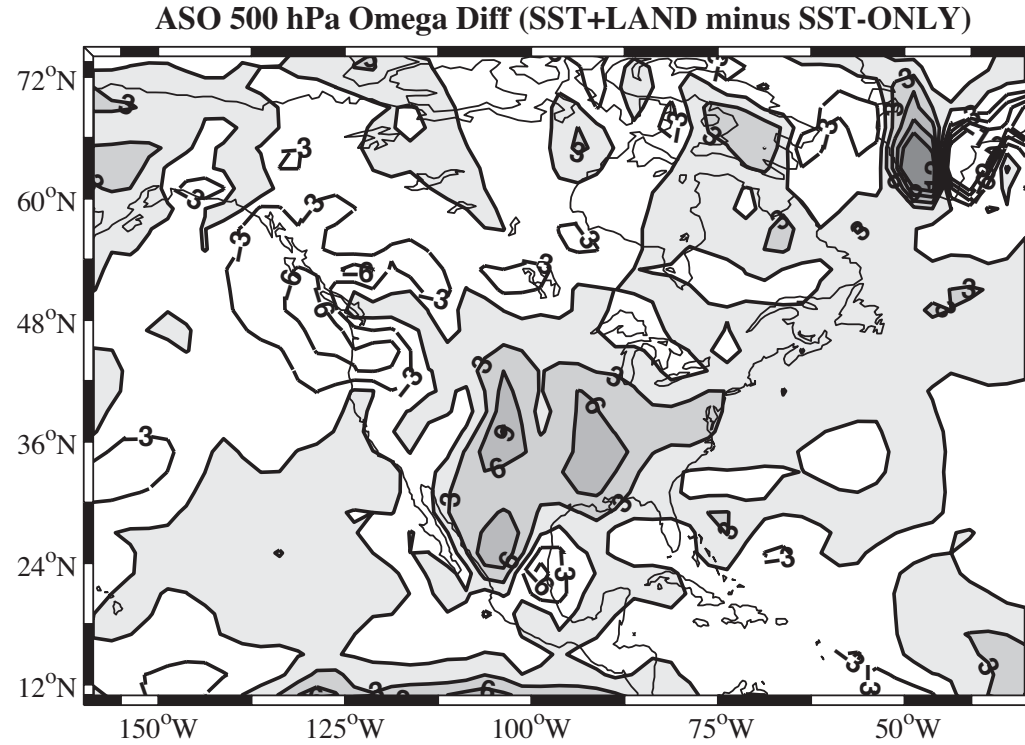
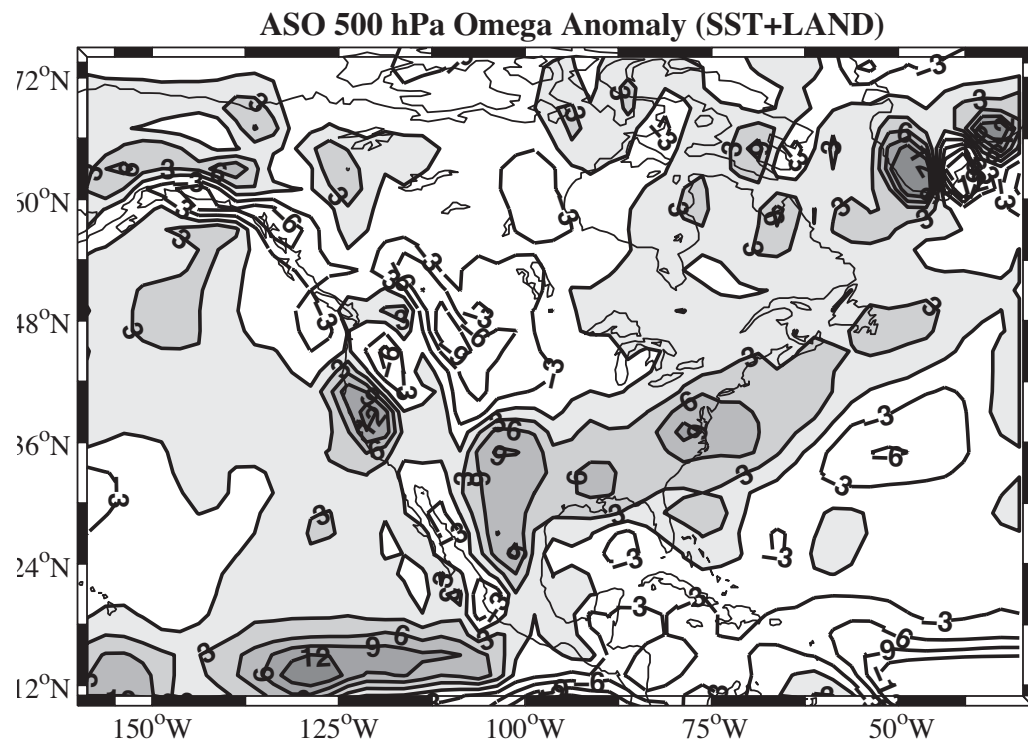
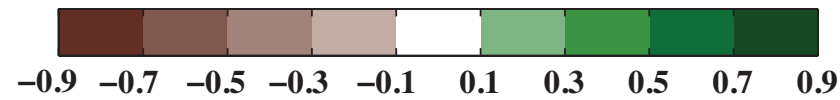
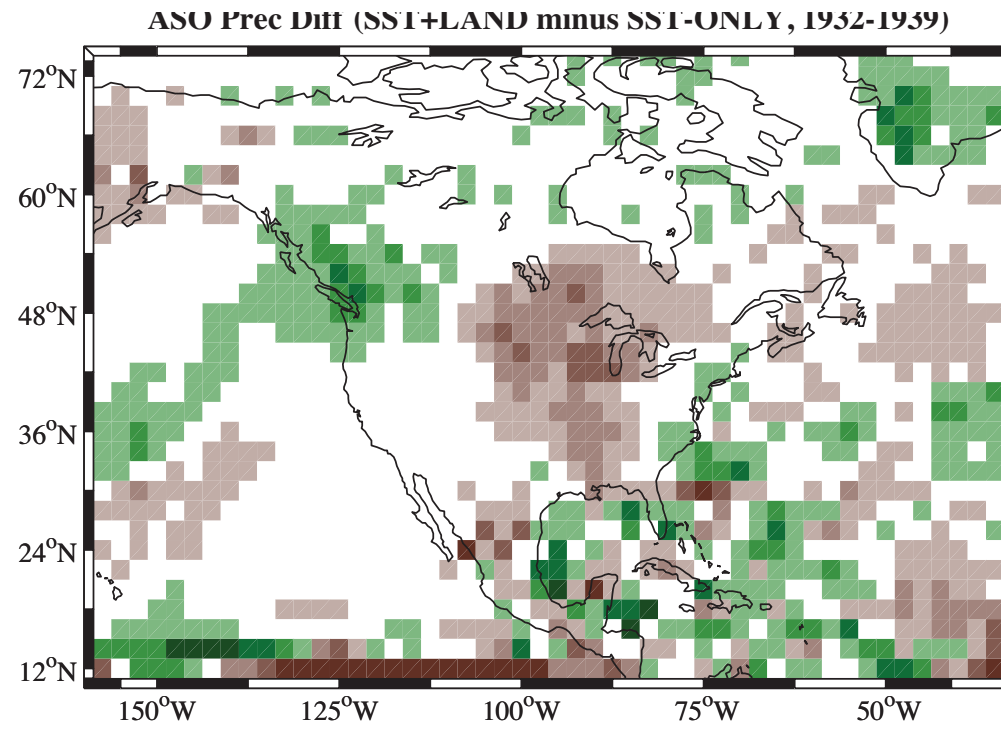
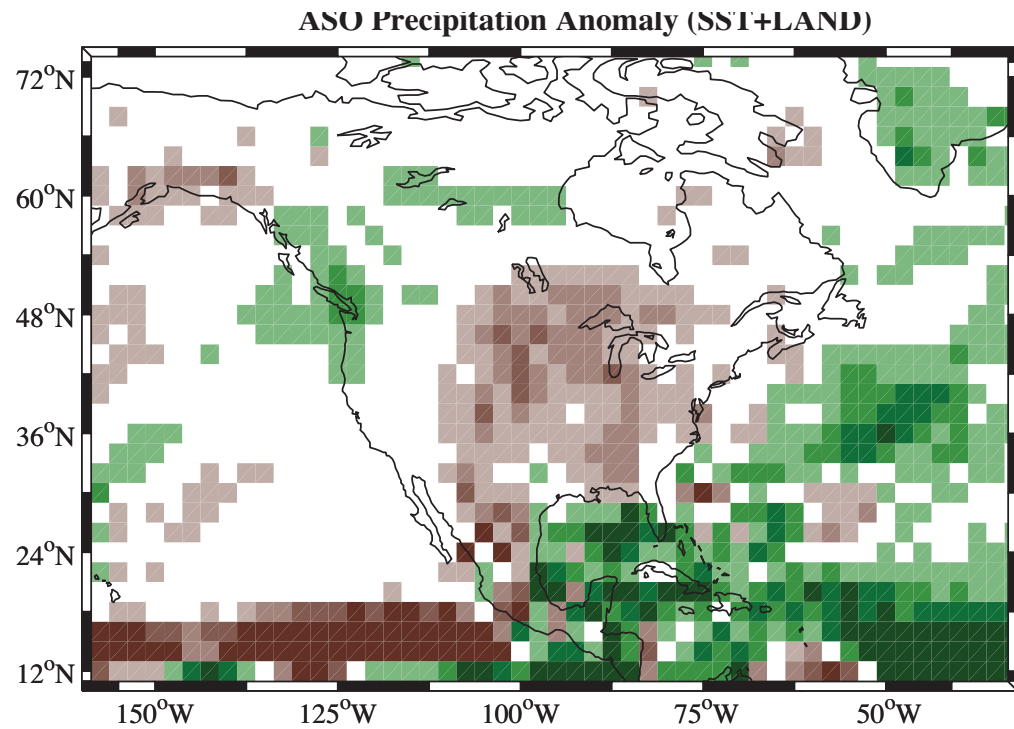


Modeled with
SST forcing
and
interactive
dust

GISS model simulation of Dust Bowl drought:

SST + dust + crop failure

Impact of dust and deveg



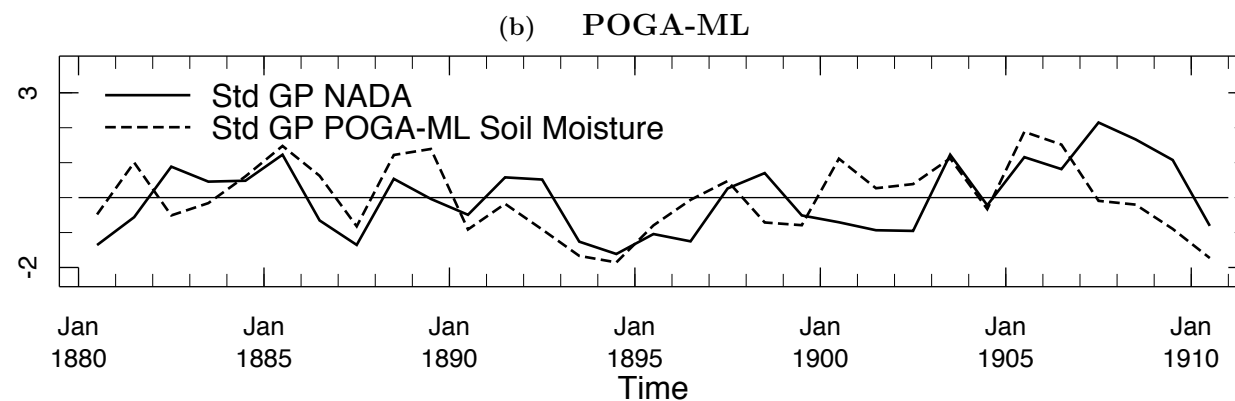
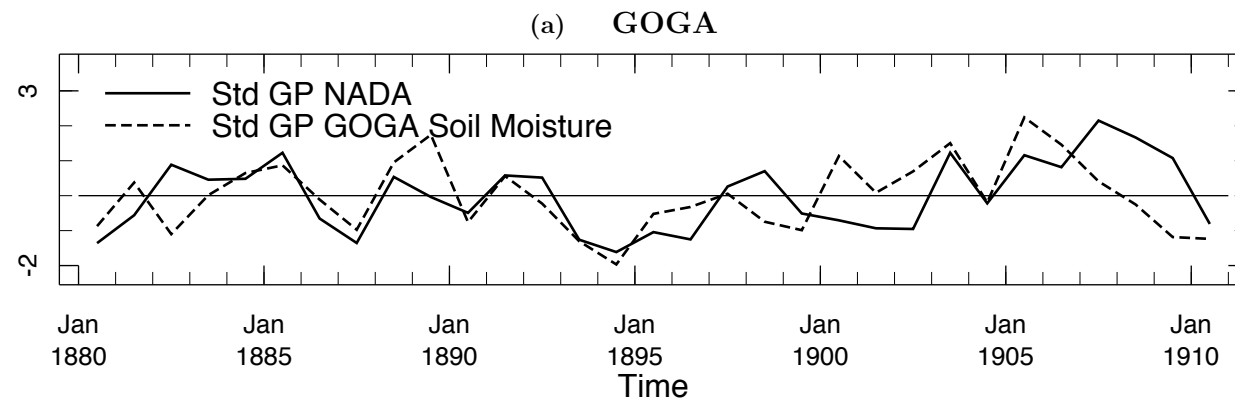
Precip

Dust and crop failure create observed drought pattern and matching continent-centered subsidence

500mb
omega

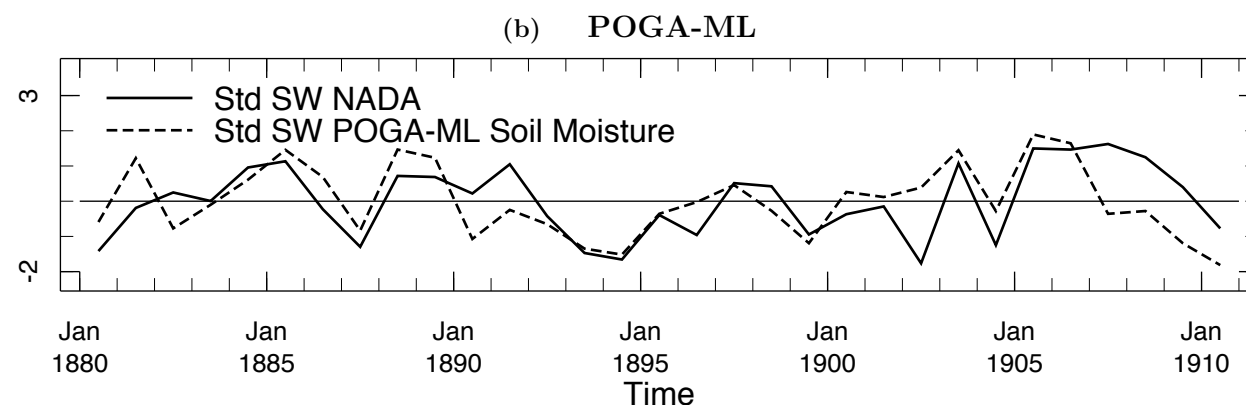
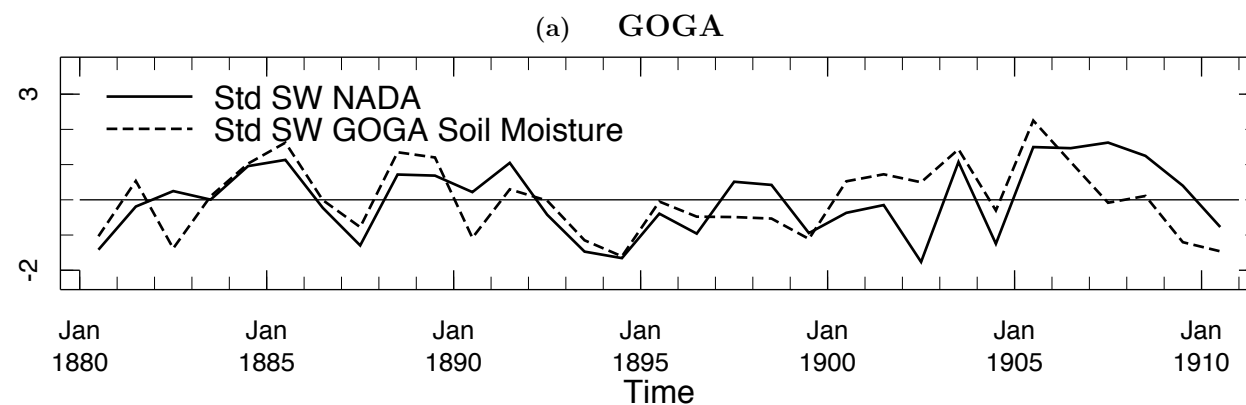
The 1890s drought, relatively brief, is quite well modeled with tropical Pacific forcing dominant.

Annual GP Indices, Std NADA V2A (solid) Model Mean (dashed)



Great Plains, modeled soil moisture and tree ring PDSI

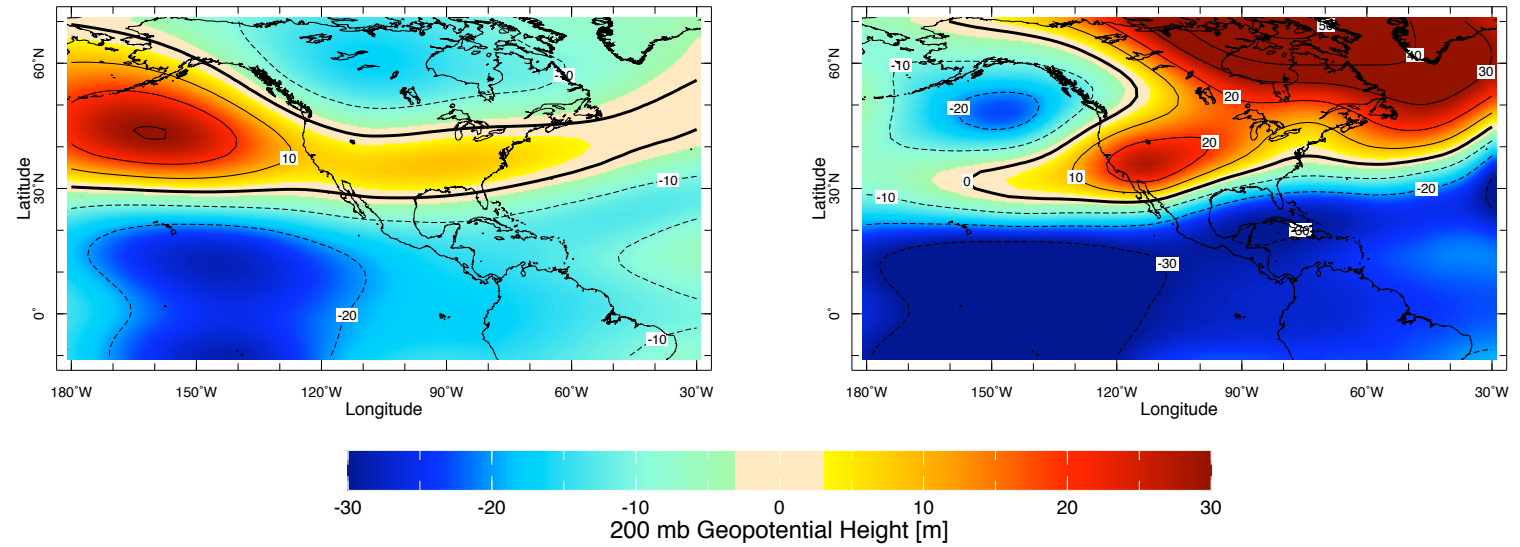
Annual SW Indices, Std NADA V2A (solid) Model Mean (dashed)



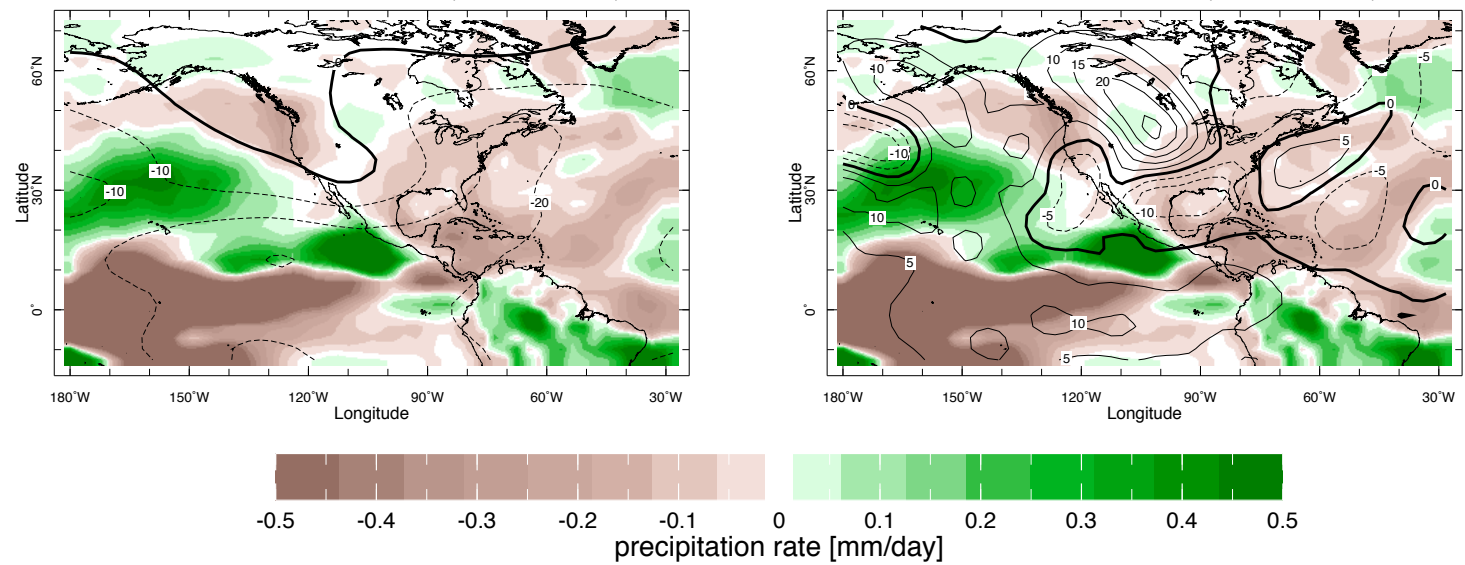
As above but for Southwest

1890s drought was associated with a La Nina but the circulation anomalies were not very typical in 20CR or model.

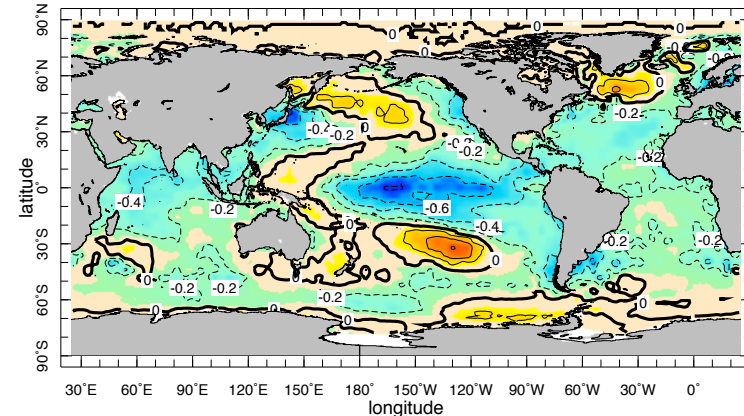
Winter COMPO 200 mb Height
Regression on TP * -1 Average Oct 1889 to Mar 1896



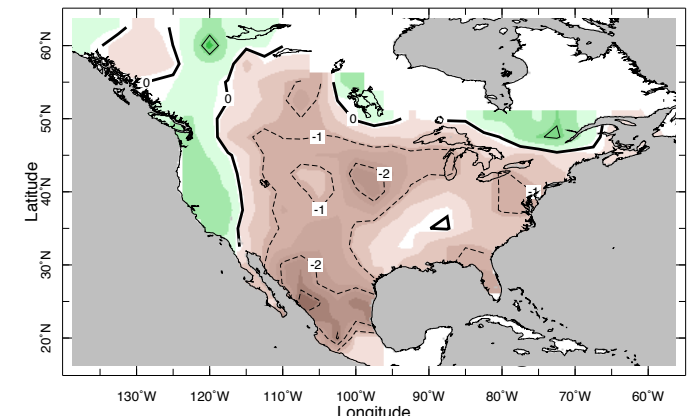
Winter GOGA Mean Average Oct 1889 to Mar 1896, Precipitation (colors) and 200 mb height (contours) 300 mb V^2 (contours)



Winter Observed SSTA

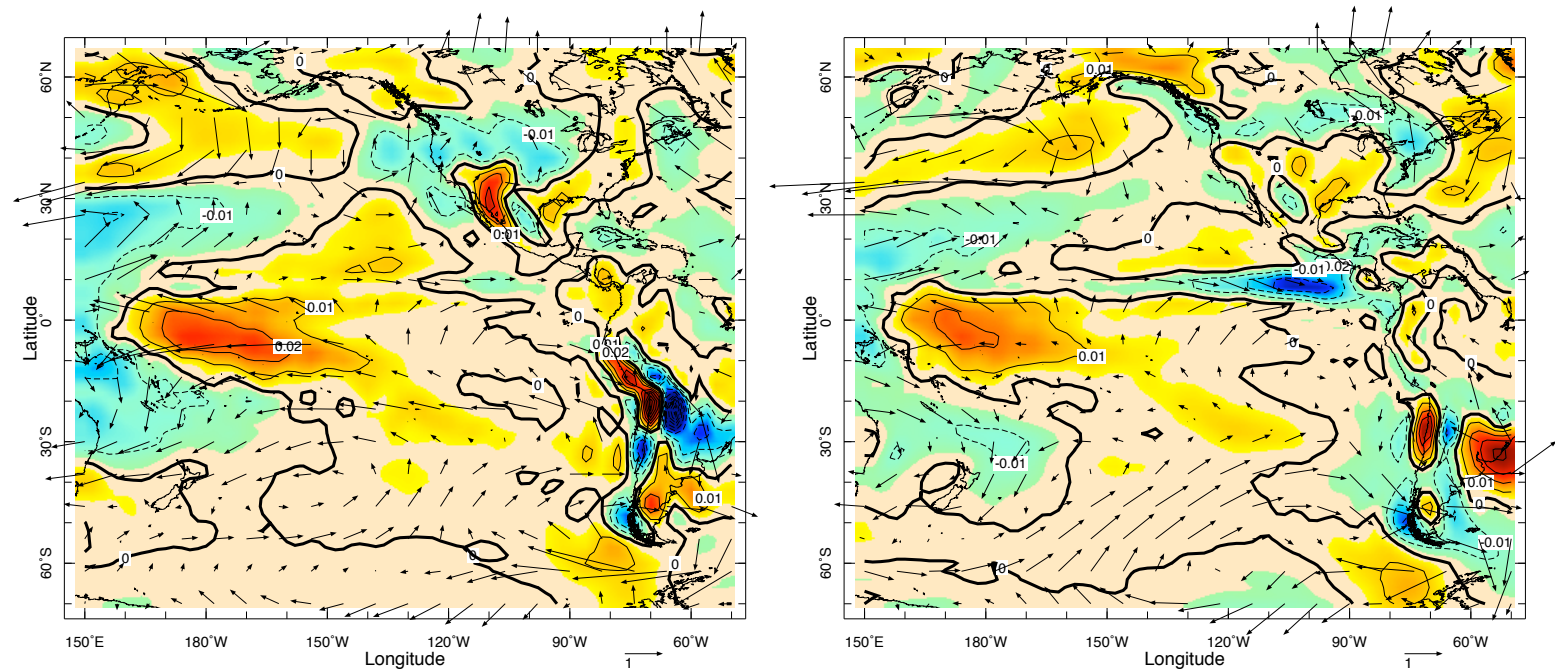


NADA V2A PDSI 1890-1896

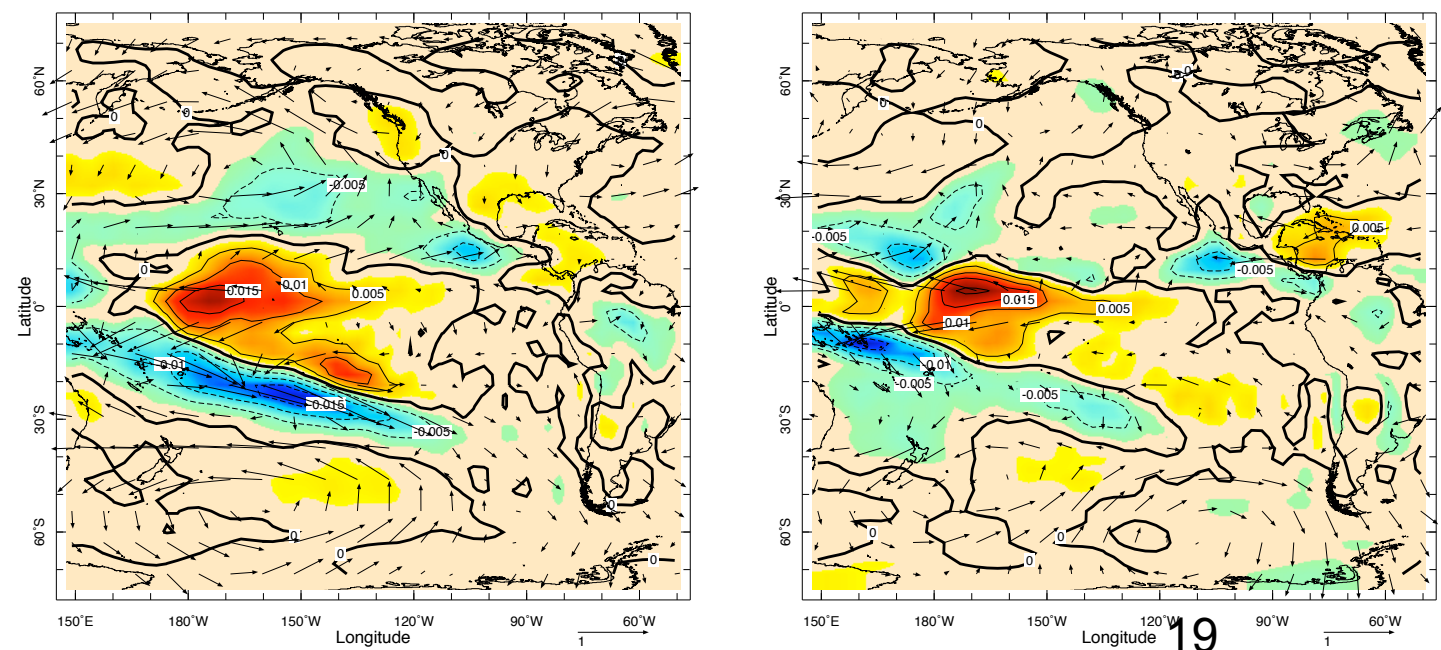


500 mb Vertical Pressure Velocity (colors, contours), 850 mb Winds (vectors)
Oct-Mar 1889-1896
Apr-Sep 1890-1896

COMPO Average



GOGA Mean Average



20CR and the model circulation match well in the tropics but not so well over N. America. In the model there is general southerly flow and descent forcing the drought.

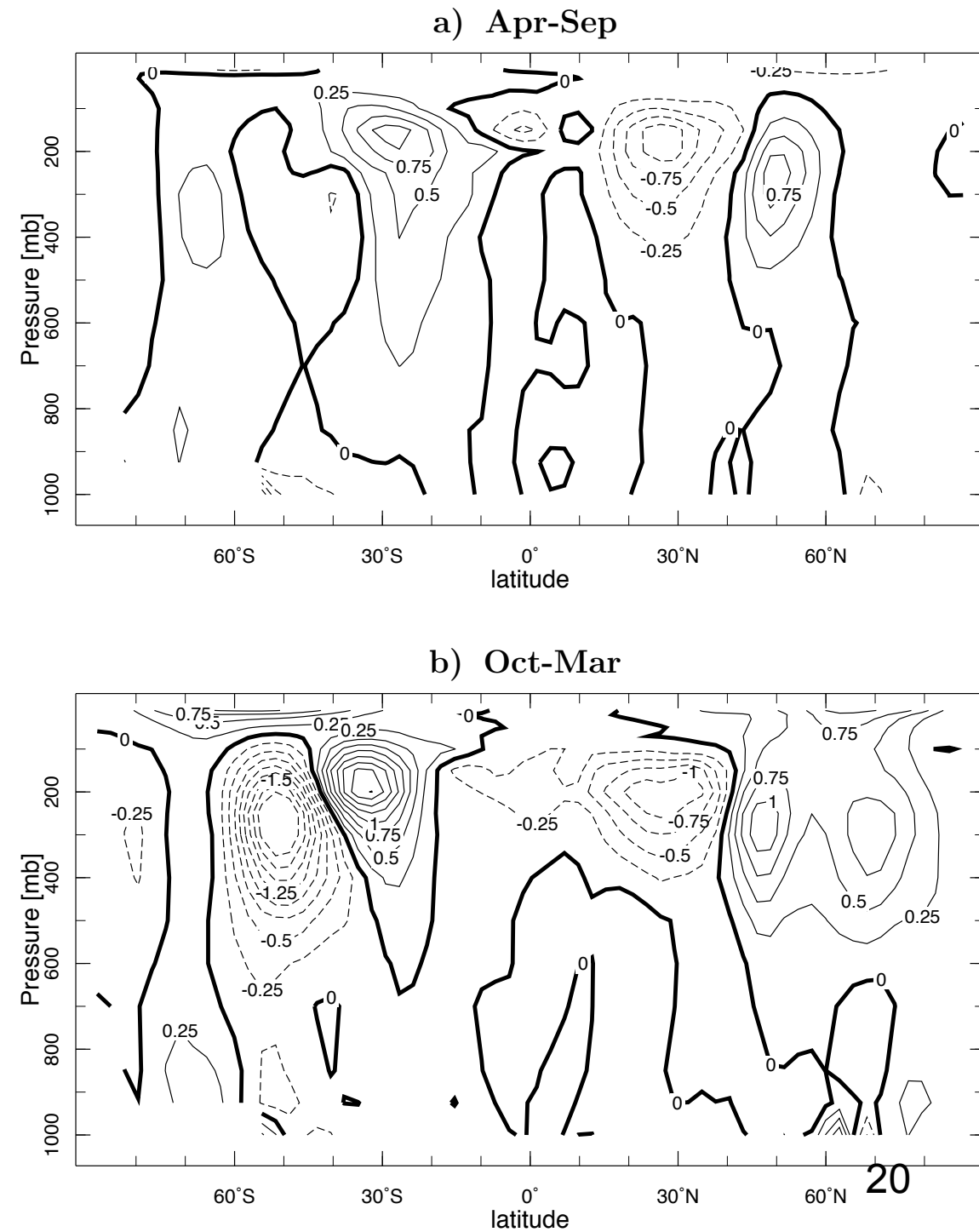
At least in the tropical Pacific SST-forced model the familiar pattern of tropically organized transient eddy-driven descent is clear in the 1890s.

For drought regimes:

1. Tropical tropospheric cooling
⇒ poleward shift of subtropical jet
2. jet shift ⇒ shift in the pattern of eddy momentum transport
3. Balancing Coriolis torque
⇒ eddy-induced upper tropospheric flow subtropics to mid-latitudes.
4. Mass convergence in the mid-latitudes ⇒ descent
⇒ suppressed precipitation.

Need to see ensemble members in 20CR to check this

POGA-ML 1890-1896 Zonal Averaged U'V'



Conclusions

1. 20CR allows examination of the similarities and differences of the 20th Century droughts and moving beyond SST-forced model simulations to validate proposed mechanisms.
2. 1950s drought akin to that expected from cold tropical Pacific-warm subtropical North Atlantic SST forcing.
3. 1930s Dust Bowl drought has, unlike SST forced droughts, massive continental-centered subsidence.
4. The Dust Bowl subsidence is consistent with dust aerosol forcing following crop failure. Dust forcing worsened drought and moved it northwards explaining its unique spatial pattern.
5. 1890s drought - more work needed!